

**DETERMINING TEMPERATURE EXTREME IN WARRI CITY, NIGER-DELTA  
REGION, NIGERIA**

**ABSTRACT**

Climate change and global warming which is also known as a change in Earth's overall climate or rising temperature have taken centre stage in international concerns, several fora and treaties have been observed with a view of stemming trend, in rising temperatures. This study evaluated ten years of maximum and minimum annual temperature of Warri (Fig. 1) in Nigeria between (2005 and 2015) to determine trends and identified extreme fluctuation in temperature. Data used for this study were sourced from the Nigerian Meteorological Agency's zonal office, Warri. An objective method for determining temperature extreme is used. Least square linear regression equation have been used to estimate temperature that would be equalled or surpassed 1%, 5% and 10% of the hours at any given location during the warmest and coldest months of the year. These equations are based on an index calculated from the three readily available parameters; the mean monthly temperature, the mean daily maximum temperature for the month and the mean daily minimum temperature for the month. The warmest month in Warri was March with a mean monthly temperature of 33.9 °C while the coldest month was July with mean monthly of 25.8°C.

**Keywords:** Mean temperature, Least square, linear regression equation, Warri, fluctuation.

**1.0 INTRODUCTION**

During the second half of the 20th century, the globally averaged air temperature increased by 0.6°C (Folland *et al.*, 2001). However, this warming was not spatially or temporally uniform. Typically, climate change detection is associated more often with the analysis of changes in extreme events than with changes in the mean (Katz and Brown,1992). Extreme

temperature events can impact many aspects of human life including: mortality, comfort, ecology, agriculture, and hydrology (Ciais *et al.*, 2005; Patzet *et al.*, 2005). Accordingly, the characterization of climate extremes can provide invaluable information for impact assessment studies, particularly those related to hydrological and environmental modeling. Recently, substantial efforts have been made to estimate not only changes in mean temperature series, but also changes in the frequency, intensity, and duration of extreme events (Easterling *et al.*, 2000; Jones *et al.*, 2001; Frichet *et al.*, 2002; Kostopoulou and Jones, 2005; Moberg and Jones, 2006; Moberg *et al.*, 2006; Brown *et al.*, 2008). These studies have analyzed temperature extremes at different spatial scales, ranging from the regional to the global. In general, most of the findings revealed a significant upward (downward) trend in the duration and frequency of hot (cold) extremes (Hunsen *et al.*, 2008). For instance, Alexander *et al.* (2006) noted a global significant decrease in cold temperature extremes throughout the second half of the 20<sup>th</sup> century. Also, Parmesan *et al.* (2000) showed that extreme weather and climate events has severely influenced ecosystems and human society. High temperatures are among the most frequently investigated extreme events; the domains in which they affect society include agriculture, water resources, energy demand and human mortality.

Changes in temperature can stimulate other components of the environment capable of arousing human health problem which link weather parameter such as solar radiation, temperature wind etc to human health; that exposure to high air temperature accompanied by intense radiation may result in heat stroke; and other health problems (Afangiden *et al.*, 2005), (Garcia-Herrera *et al.*, 2005). It is also expected that the geographical range of vector will be expanded as temperature rises. Susan *et al.* (2008) stated that mosquitoes, tick, rodents and other vectors are expanding their geographic range altering long established patterns of diseases as a result of global warming. Temperature affects pathogenic replication, maturation and period of infectivity, hence indirectly, human health. West Nile encephalitis,

lung cancer, heart diseases, asthma and allergies and other health problem are linked to global warming (Susan *et al.*, 2008). Hence, the trends of an extreme fluctuation in temperature in Warri are needed to be determined and identified so as to avoid any health hazard.



Fig 1: Map of Delta State Showing location of the Study Area (Warri)

## 2.0 Derivation

### 2.1 Warm Temperature

The ideal method for determining frequency distribution of temperature would be to obtain actual distributions of hourly temperature for  $n$  long period. These data are readily available for a large number of stations in Nigeria but on a world-wide basis, there is an insufficient number of stations with complete, long term (at least 10 years) records to permit an accurate analysis. This difficulty was overcome in an earlier study by (Tattelman *et al* 1976 a&b) on the frequency of high temperatures. In that report he determined that high temperatures corresponding to low probabilities are found where the monthly mean temperatures are highest and the mean daily range is greatest. A simple index of these values is expressed by (Tattelman *et al* 1977)

$$I_w = T + (T_x - T_n) \dots\dots\dots (1)$$

Where  $I_w$  is the warm temperature index

$T$  is the mean,

$T_x$  is the mean daily maximum, and

$T_n$  is the mean daily minimum temperature for the warmest month.

Since good climatic records are readily available for these parameters, it was decided to determine if equation (1) is applicable on a more general basis than just the very hot locations for which it was originally used. The index was correlated with each of the observed 1%, 5% and 10% warm temperatures during the warmest month.

The following regression lines for the 1%, 5% and 10% temperatures were found by the method of least squares:

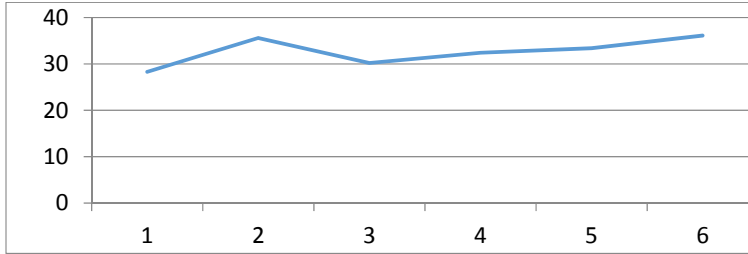
$$T_{1\%} = 0.676 I_w + 10.657$$

$$T_{5\%} = 0.733 I_w + 5.682$$

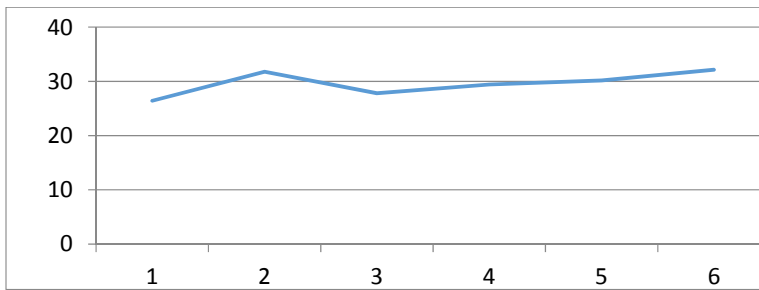
$$T_{10\%} = 0.762 I_w + 2.902.$$

**Table 1: The Mean Daily Maximum and Minimum Warm Temperature**

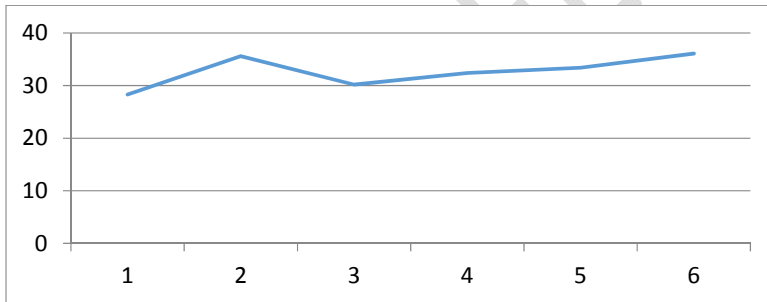
Month	The mean Daily maximum $T_x$	The mean Daily minimum $T_n$	Monthly mean $T$	Temperature Index $I_w$ $T + (T_x - T_n)$	$T_{1\%}$ ( $^{\circ}\text{C}$ )	$T_{5\%}$ ( $^{\circ}\text{C}$ )	$T_{10\%}$ ( $^{\circ}\text{C}$ )
OCT	29.0	26.0	25.3	28.3	29.788	26.422	24.467
NOV	29.8	20.8	26.6	35.6	34.723	31.776	30.029
DEC	29.6	26.8	27.4	30.2	31.072	27.819	26.004
JAN	30.1	25.4	27.7	32.4	32.559	29.431	27.591
FEB	30.8	27.4	30.0	33.4	33.235	30.164	28.353
MAR	30.2	28.0	33.9	36.1	35.061	32.143	30.410



**Fig 2: 1% Warm Temperature(°C)**



**Fig 3: 5% Warm Temperature(°C)**



**Fig 4: 10% Warm Temperature (°C)**

## 2.2 Cold Temperatures

Since Equation (1) proved successful for describing warm temperature extremes, the same principle was used to estimate cold temperature extremes. A cold temperature index,  $I_c$ , is expressed by:

$$I_c = T - (T_X - T_N) \dots \dots \dots (2)$$

Where  $T$  is the mean

$T_X$  is the mean daily maximum for the coldest month, and

$T_N$  is the mean daily minimum temperature for the coldest month.

The index was correlated with the 1%, 5% and 10% cold temperatures during the coldest month.

The following regression lines for the 1%, 5% and 10% temperatures were found by the method of least squares:

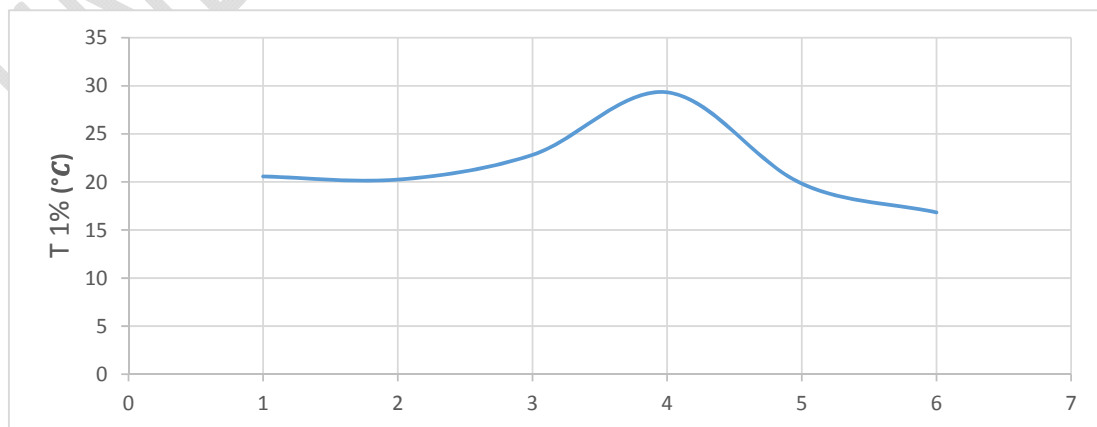
$$T_{1\%} = 1.069I_c - 7.013$$

$$T_{5\%} = 1.084 I_c - 3.050$$

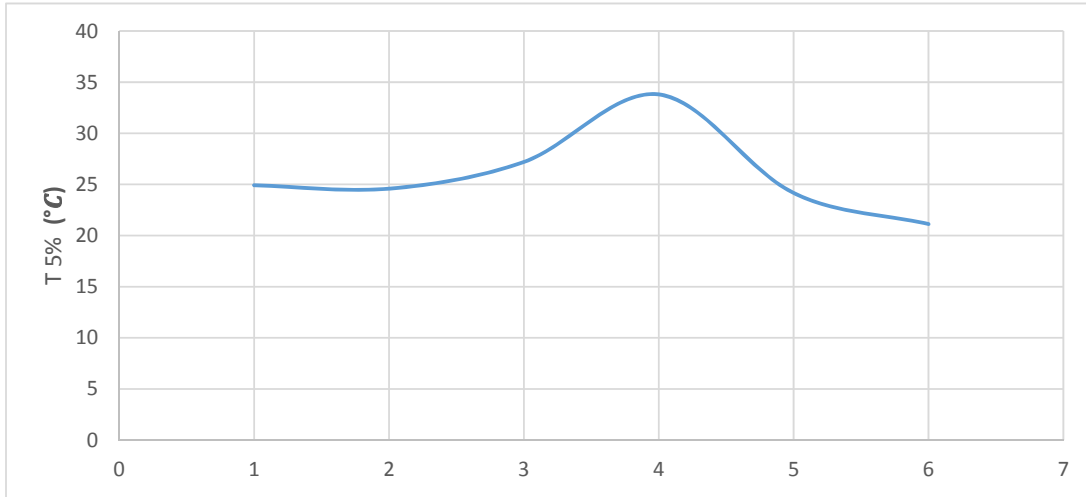
$$T_{10\%} = 1.082I_c - 0.704$$

**Table 2: The Mean Daily Maximum and Minimum Cold Temperature**

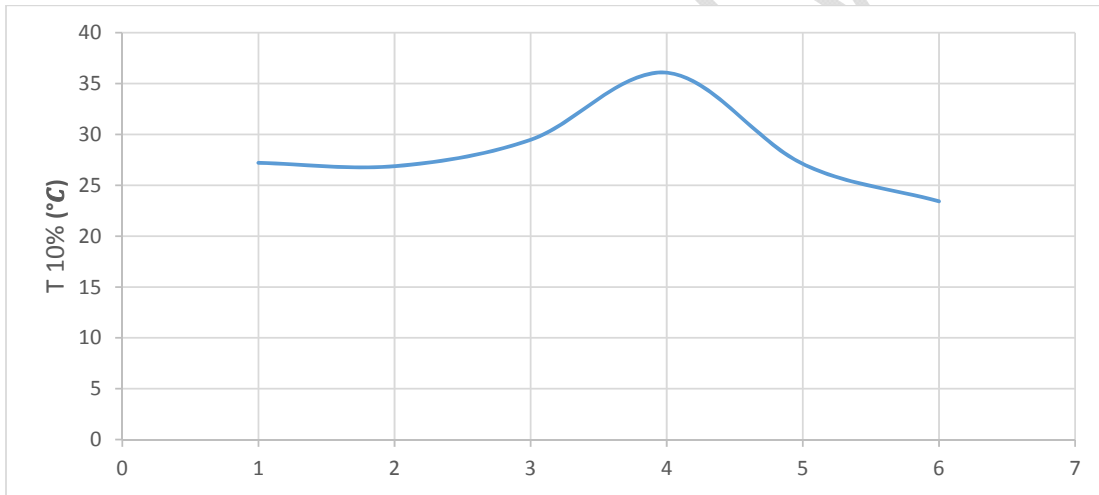
	The mean Daily maximum $T_X$	The mean Daily maximum $T_N$	Monthly mean $T$	Temperature Index $I_c$ $T-(T_X-T_N)$	$T_{1\%}$ (°C)	$T_{5\%}$ (°C)	$T_{10\%}$ (°C)
APR.	30.6	27.2	29.2	25.8	20.567	24.917	27.216
MAY.	30.4	27.8	28.1	25.5	20.246	24.586	26.887
JUN.	28.8	26.8	29.9	27.9	22.812	27.194	29.484
JUL.	28.4	26.6	25.8	34.0	29.333	33.806	36.084
AUG.	26.6	26.6	26.9	25.7	19.819	24.158	27.103
SEP.	23.4	23.4	27.3	22.3	16.826	21.123	23.425



**Fig. 5: 1% Cold Temperature (°C)**



**Fig. 6: 5% Cold Temperature (°C)**



**Fig 7: 10% Cold Temperature (°C)**

### 3.0 DISCUSSIONS

The warmest month, between the year (2005-2015) in warri, Nigeria is March with an average monthly temperature (mean) of  $33.9^{\circ}\text{C}$  (*Table 1*). For warm temperature, the estimated temperature is at extreme at  $35.0^{\circ}\text{C}$  while the observed temperature is at  $31.8^{\circ}\text{C}$ . (*Fig. 2, 3&4*)

The coldest month in Warri, Nigeria was JULY, with monthly mean temperatures averaging  $25.8^{\circ}\text{C}$  from the year (2005-2015) (Table 2). Warri thus experienced extreme coldness at estimated temperature ranging from  $16.0^{\circ}\text{C}$  to  $24.0^{\circ}\text{C}$  and observed temperature of  $26.6^{\circ}\text{C}$  within the months of April, May, June, July, August, and September within the year 2005-2015. (Fig. 5, 6 & 7)

Warri is a typical Niger Delta area and part of the global environment. Warri has the potential for sustained anthropogenic greenhouse gases blamed for global temperature rise. From the results of data analysis, it is apparent that; trends in temperature of Warri conform with global trend and mean annual temperature has varied remarkably in Positive and negative extreme fluctuations which have influenced human health condition in Warri. The relationship between temperature and health condition in Warri is inverse, hence as temperature increases cases of health condition decrease. There is a gradual rising (upward) fluctuation in temperature trends of Warri.

#### **4.0 CONCLUSION**

This method represents a unique tool for estimating warm and cold temperature extremes. Least squares linear regression equation have been used to calculate temperatures that would equalled or exceeded 1, 5 and 10% of the hours at Warri during the warmest month of the year while Analogous regression have been used to calculate temperatures equalled or less than 1, 5 and 10% of the hours at Warri during the coldest month of the year. The high temperatures described herein normally will be encountered during periods of strong sunshine and fairly light winds. Similarly, low temperatures generally will be encountered during nights with clear skies and little or no wind. The ground can attain temperatures from  $15^{\circ}\text{C}$  to  $30^{\circ}\text{C}$  higher and  $50^{\circ}\text{C}$  to  $10^{\circ}\text{C}$  lower than that of the free air, depending upon radiation, conduction, wind, and turbulence. Since the design philosophy for temperature



extremes, as adopted for this report, is based on the probability of being exceeded during the warmest (coldest) month of the year, the number of hours this temperature is encountered during all other months will be smaller than in the warmest (coldest) month. Also, the annual risk will be roughly one tenth of that shown for the warmest (coldest) month. It should be noted that the warmest (coldest) month is not necessarily the same for each station. This fact, however, does not alter the desired concept of percentage of time (risk) of inoperability for design. Hence, as temperature increases, cases of health condition decrease. Therefore, there is a gradual rising (upward) fluctuation in temperature trends of Warri.

## **5.0 RECOMMENDATION**

From the mitigation measure a sustainable, efficient and affordable health care delivery system should be put in place to meet the health need of the people since they cannot run from the environment. Also a deliberate plan should be made to monitor trends in temperature and other climatic parameters, so as to avoid any implications on health, agriculture, wildlife and the economy.

## **6.0 REFERENCES**

1. Afangideh, A: Okpiliya, F., and Ekanem E. (2005)The changing annual rainfall and temperature average in the humid tropical city of Uyo, Southern Nigeria. African Journal of Environmental Pollution and Health 4(2) Pp 54-61.
2. Alexander, L. V., Zhang, X., Peterson, T. C., Caesar, J., Gleason, B., Klein Tank, A. M. G., Haylock, M., Collins, D., Trewin, B., Rahimzadeh, F., Tagipour, A., Ambenje, P., Rupa Kumar, K., Revadekar, J., and Griffiths, G. (2006) Global observed changes

- in daily climate extremes of temperature and precipitation, *J. Geophys. Res. Atmos.*, 111, D05109, doi:10.1029/2005JD006290.
3. Brown, S. J., Caesar J., and Ferro, C. A. T. (2008). Global changes in extreme daily temperature since 1950, *Journal Geophys. Res. Atmos.*, 113, D05115, doi:10.1029/2006JD008091.
  4. Ciais, Ph., Reichstein, M., Viovy, N., Granier, A., Ogée, J., Allard, V., Aubinet, M., Buchmann, N., Bernhofer, Chr., Carrara, A., Chevallier, F., De Noblet, N., Friend, A. D., Friedlingstein, P., Grunwald, T., Heinesch, B., Keronen, P., Knohl, A., Krinner, G., Loustau, D., Manca, G., Matteucci, G., Miglietta, F., Ourcival, J. M., Papale, D., Pilegaard, K., Rambal, S., Seufert, G., Soussana, J. F., Sanz, M. J., Schulze, E. D., Vesala, T., and Valentini, R. (2005). Europe-wide reduction in primary productivity caused by the heat and drought in 2003, *Nature*, 437(7058), 529–534.
  5. Easterling, D. R., Meehl, G. A., Parmesan, C., Changnon, S. A., Karl, T. R., and Mearns, L. O. (2000) Climate extremes: observations, modeling, and impacts, *Science*, 289, 2068–2074.
  6. Folland, C. K., Karl, T. R., Christy, J. R., Clarke, R. A., Gruza, G. V., Jouzel, J., Mann, M. E., Oerlemans, J., Salinger M. J. and Wang, S. W. (2001). Observed Climate Variability and Change in Climate Change.
  7. Frich, P., Alexander, L. V., Della-Marta, P., Gleason, B., Haylock, M., Klein Tank, A. M. G., and Peterson, T (2002). Observed coherent changes in climatic extremes during the second half of the twentieth century, *Clim. Res.*, 19, 193–212.
  8. Garcia-Herrera, R., Diaz, J., Trigo, R. M., and Hernandez, E. (2005). Extreme summer temperatures in Iberia: health impacts and associated synoptic conditions, *Ann. Geophys.*, 23, 239–251, doi: 10.5194/angeo-23-239.

9. Hansen, J. (2005). Global temperature trend: 2005 summation GISS surface Temperature Analysis, NASGoddard institute for space studies (GISS) and Columbia University Earth institute. New York <http://data.giss.nasa.gov/gis/temp/2005>. Retrieved 4/10/2008.
10. Jones, P. D., Jones, R. N., Nicholls, N., and Sexton, D. H. M. (2001). Global temperature change and its uncertainties since 1861, *Geophys. Res. Lett.*, 28, 2621–2624.
11. Katz, R. W. and Brown, B. G. (1992) Extreme events in a changing climate: variability is more important than averages, *Climatic Change*, 21, 289–302.
12. Kostopoulou, E. and Jones, P. (2005) Assessment of climate extremes in Eastern Mediterranean, *Meteorol. Atmos. Phys.*, 89, 69–85.
13. Moberg, A., Jones, P. D., Lister, D., Walther, A., Brunet, M., Jacobeit, J., Alexander, L. V., Della-Marta, P. M., Luterbacher, J., Yiou, P., Chen, D., Klein Tank, A. M. G. (2006). Indices for daily temperature and precipitation extremes in Europe analyzed for the period 1901–2000, *J. Geophys. Res.*, 111, D22106, doi:10.1029/2006JD007103.
14. Parmesan, C., Changnon, S. A., Karl, T. R., and Mearns, L. O. (2000). Climate extremes: observations, modeling, and impacts, *Science*, 289, 2068–2074.
15. Patz, J. A., Campbell-Lendrum, D., Holloway, T., and Foley, J. A. (2005). Impact of regional climate change on human health, *Nature*, 438, 310–317.
16. Paul I. Tattelman and Arthur J. Kantor (1976a). Atlas of Probabilities of Surface temperature extremes: Part I, Northern Hemisphere, *Environ. Res. Pap.*, No. 557, AFGL-TR-76-0084.

17. Paul I. Tattelman and Arthur J. Kantor (1976b). Atlas of Probabilities of Surface temperature extremes: Part II, Southern Hemisphere, Environ, Res, Pap., No. 558, AFGL-TR-77-0001.
18. Paul I. Tattelman and Arthur J. Kantor (1977). A Method for Determining Probabilities of Surface Temperature Extremes. *Air force Geophysics Laboratory, Hunscom AFB, Mass, 01731*
19. Susan B.M.D, Yi-An Ko, Stephanie S., Beth H. and Julie C. (2008). World Health Day: The Hazards of Global Warming to Your Health; Human health, influenced by a complex system of biological, social, economic, political and geographic factors, is particularly vulnerable to the effects of global warming.