

METEOROLOGICAL MODELS FOR DETERMINATION OF SURFACE RADIO REFRACTIVITY OVER NIGERIA

Abstract

The problem of quantifying surface refractivity in terms of atmospheric variables in different climatic scenarios is very complex because the meteorological variables required are neither perfectly correlated nor totally uncorrelated. The previous studies showed that linear meteorological expressions obtained were localized and could not be generally applied. It is therefore required that more locations should be investigated to deduce new linear meteorological models best suitable for estimation of surface refractivity.

Surface meteorological data, including pressure, temperature and relative humidity, was downloaded from Modern-Era Retrospective analysis for Research and Application (MERRA - 2) for six locations defined by different climatic conditions over Nigeria, namely Yenagoa, Abakaliki, Ibadan, Anyigba, Jalingo and Sokoto for a period of 40 years partitioned into two periods of 20 years each. The 1979 – 1998 dataset of atmospheric temperature, T (K), atmospheric pressure, P (hPa) and relative humidity, R_h (%) were extracted using text import wizard to calculate surface refractivity using existing model (N_{itu-r}). Regression analysis was carried out to obtain new linear meteorological expressions as function of temperature (N_t), relative humidity (N_{Rh}), and combination of relative humidity and temperature ($N_{Rh,t}$). The new expressions were tested using 1999 – 2018 meteorological dataset and the results of surface refractivity from the new linear expressions were compared with values from existing equations.

It was observed that high relative humidity and low temperature values prevalent in the three southern stations – Yenagoa, Abakaliki and Ibadan – significantly enhanced the suitability of $N_{Rh,t}$ linear meteorological model for estimation of surface refractivity values comparable to N_{itu-r} . Contrary to the observations in the southern stations, in the northern stations, relative humidity contributions (N_{Rh}) had the best linear correlation of 0.96 at Anyigba, 0.98 at Jalingo and Sokoto. Estimation of surface refractivity from the new linear meteorological models was found to be best as a function of temperature and relative humidity in the south and a function of relative humidity in the north. Linear meteorological models as a function of relative humidity and temperature were best suitable for stations in the south with minimum correlation of 0.98 while linear expressions as a function of relative humidity only were best suitable for the stations in the north with minimum correlation of 0.96.

1.0 Introduction

Atmospheric refractivity has been well studied over the years and some mathematical expressions have been developed to estimate the value of refractivity (Amajama and Eshiet, 2016). However, most of the expressions are complex and involve some nonlinear expressions. The complexity makes it more difficult to integrate the refractivity expressions into other formula to develop simple closed-form mathematical expressions needed in wireless link design (Joseph, 2015). In recent years, researchers have tried to develop simple mathematical equations that can be used to estimate the radio refractivity from the atmospheric parameters (Jari and Ismo, 2015). Prediction of radio refractivity has been required since the early days of microwave radio links. Although one of the earliest prediction models is still in use, it is only valid for continental temperate climate (ITU – R, 2013).

In a study conducted by Stephen, *et al.*, (2014), radiosonde data from 65 stations covering parts of Australia, India, South Africa, Europe and North America, most of them with around 16 years of data, typically 1997–2012, were processed. A new prediction model for the distribution of refractivity gradient in the atmospheric surface layer, having better prediction accuracy than existing models, and using only

48 data that can be obtained from surface weather stations, was obtained with rms error of 17 N-units per
49 km, and correlation coefficient of 0.79

50 In the recent research conducted by Enyenihi, *et al.*, (2019), radiosonde meteorological data for
51 Cross River state, Nigeria was obtained from Nigerian Meteorological Agency (NIMET) for the twelve
52 months in 2013. The study examined the correlation among radio refractivity and the meteorological
53 parameters, namely atmospheric temperature (T), atmospheric pressure (P) and relative humidity (H).
54 The correlation values from the results showed that there is a positive correlation between surface
55 refractivity (N) and temperature (T) and pressure (P). However, there is a negative correlation between
56 N and relative humidity (H) which indicated inverse proportionality between N and H. In all, T and
57 products of TPH showed highest correlation values for the model training datasets; as such the two
58 parameters were used in the development of the linear regression model that can be used to estimate
59 the radio refractivity for the available meteorological dataset.

60 The model was developed using XURU online regression tool where the values of T was X1, TPH
61 was X2 and N was Y. The dataset of T, TPH and N for the 12 months were pasted into the text box
62 provided in Xuru Multiple Linear Regression (MLR) webpage and the model obtained from the Xuru MLR
63 was given as;

$$64 \quad N = 2.38T + 0.000051[(T)(P)(H)] + 198.38 \quad 1.0$$

66 Two sample meteorological dataset from published articles were also used by Enyenihi, *et al.*, (2019) to
67 validate the model. The model gave a maximum absolute percentage error of 2.46 % for the first test
68 meteorological dataset while it gave a maximum absolute percentage error of 1.25 % for the second test
69 meteorological dataset. The results that was obtained from the new model showed that the model can
70 estimate refractivity with a maximum prediction error of about $\pm 3.35\%$.

71 In this study, a set of meteorological data measured over 20 years was used to deduce useful
72 linear expressions for the estimation of surface refractivity adaptable to specific climatic features in six
73 locations across Nigeria. The linear expressions were tested using meteorological data from another
74 period of 20 years. The values of surface refractivity from the new linear expressions were compared
75 with values from existing equations.

77 **2.0 Theoretical Background**

78 The refractive index of the troposphere is an important factor in predicting the performance of
79 terrestrial radio links. Refractive index variations of the atmosphere affect radio frequencies above 30
80 MHz, although these effects become significant only at frequencies greater than about 100MHz
81 especially in the lower atmosphere. The radio refractive index n of the troposphere deviates slightly
82 from unity due to the polarisability of the constituent molecules by the incident electromagnetic field,
83 and the quantum mechanical resonances at certain frequency bands (Ayantunji and Okeke, 2011). While
84 molecular polarisability is independent of frequency up to millimeter waves, molecular resonance is
85 totally frequency dependent, and n tends to be dispersive above approximately 50 GHz.

86 Radio refractivity N is a measure of deviation of refractive index n of air from unity which is
87 scaled-up in parts per million to obtain more amenable figures. Thus, N is a dimensionless quantity
88 defined as measured in N units (ITU-R, 2003).

$$89 \quad N = (n - 1) \times 10^6 \quad 2.0$$

91 N depends on meteorological parameters of pressure P (hPa), temperature T (K) and water vapour
92 pressure e (hPa), as given by the relation (ITU-R, 2013):
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$$N_{itu-r} = \frac{77.6P}{T} + 3 \times 10^5 \frac{e}{T^2} \quad 3.0$$

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97 The vapour pressure is also related to the relative humidity H (%):

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$$e = \frac{He_s}{100} \quad 4.0$$

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101 e_s is the maximum (or saturated) vapour pressure at the given air temperature t °C, and may be
 102 obtained from:

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$$e_s = 6.11 \exp \left[\frac{17.502t}{(t+240.97)} \right] \quad 5.0$$

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106 Generally P and e decrease rapidly with height while T decreases slowly with height (ITU-R, 2013).

107 Horizontal variation of refractive index is generally negligible in the lower troposphere
 108 compared to the large-scale vertical variation which has a median gradient of about –40 N/km near the
 109 surface in mid-latitude and most temperate regions. However, significant deviations can arise from local
 110 or mesoscale meteorological factors, especially in the tropics. This horizontal variation of refractive
 111 index is very significant over Nigeria because of the significant change in climatic condition from the
 112 coastal region in the extreme south to the semi-arid region in the extreme North.

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114 3.0 Climate of the Study Area

115 The study area in this research spread across six geopolitical zone across Nigeria. Each locations
 116 has a diverse climatic pattern to ensure adequate representation of surface refractivity patterns. The
 117 locations are shown in figure 1 namely Yenagoa in south-south zone, Abakaliki in south-east zone,
 118 Ibadan in south-west zone, Anyigba in north-central zone, Jalingo in north-east zone and Sokoto in
 119 north-west zone. The climate at Yenagoa is tropical. Most months of the year are marked by significant
 120 rainfall. The short dry season has little impact. The average annual temperature is 26.7 °C in Yenagoa. In
 121 a year, the average rainfall is 2899 mm (Odjugo, 2013). The driest month is January, with 40 mm of rain.
 122 The greatest amount of precipitation occurs in September, with an average of 472 mm. March is the
 123 warmest month of the year. The temperature in March averages 28.0 °C (Abu, 2007). The lowest
 124 average temperatures in the year occur in July, when it is around 25.4 °C. There is a difference of 432
 125 mm of precipitation between the driest and wettest months (Nwagbara, et al., 2009). The variation in
 126 temperatures throughout the year is 2.6 °C.

127 Abakaliki climate is classified as tropical. The average temperature in Abakaliki is 27.7 °C.
 128 Precipitation here averages 1918 mm (Maria, 2013). Precipitation is the lowest in December, with an
 129 average of 7 mm. Most precipitation falls in September, with an average of 291 mm. At an average
 130 temperature of 29.6 °C, March is the hottest month of the year (James, 2009). In August, the average
 131 temperature is 26.0 °C (Jagtap, 2007). It is the lowest average temperature of the whole year. Between
 132 the driest and wettest months, the difference in precipitation is 284 mm. The average temperatures
 133 vary during the year by 3.6 °C.

134 The climate is tropical in Ibadan. The average annual temperature is 26.5 °C in Ibadan. About
 135 1311 mm of precipitation falls annually. The driest month is January. There is 6 mm of precipitation in
 136 January. In June, the precipitation reaches its peak, with an average of 190 mm (Ayoade, 2004). With an
 137 average of 28.6 °C, March is the warmest month. At 24.1 °C on average, August is the coldest month of
 138 the year. The precipitation varies 184 mm between the driest month and the wettest month. The
 139 variation in annual temperature is around 4.5 °C.

140 Anyigba has a tropical climate. The average temperature in Anyigba is 25.4 °C. The average
141 annual rainfall is 1321 mm. The least amount of rainfall occurs in January. The average in this month is 7
142 mm. The greatest amount of precipitation occurs in September, with an average of 265 mm. The
143 temperatures are highest on average in March, at around 27.9 °C (Nwafor, 2007). The lowest average
144 temperatures in the year occur in August, when it is around 23.8 °C. The variation in the precipitation
145 between the driest and wettest months is 258 mm. The variation in temperatures throughout the year is
146 4.1 °C.

147 The climate is tropical in Jalingo. The temperature here averages 27.3 °C. About 1053 mm of
148 precipitation falls annually. The driest month is January. There is 0 mm of precipitation in January. In
149 August, the precipitation reaches its peak, with an average of 224 mm. With an average of 30.6 °C, April
150 is the warmest month (IPCC, 2007). At 25.2 °C on average, December is the coldest month of the year.
151 The precipitation varies 224 mm between the driest month and the wettest month. The variation in
152 annual temperature is around 5.4 °C.

153 The climate in Sokoto is referred to as a local steppe climate. There is little rainfall throughout
154 the year. In Sokoto, the average annual temperature is 28.4 °C (Ayuba et al., 2007). The average annual
155 rainfall is 629 mm. The least amount of rainfall occurs in January. The average in this month is 0 mm.
156 The greatest amount of precipitation occurs in August, with an average of 211 mm. The temperatures
157 are highest on average in April, at around 33.2 °C. The lowest average temperatures in the year occur in
158 January, when it is around 24.5 °C. The variation in the precipitation between the driest and wettest
159 months is 211 mm. The variation in temperatures throughout the year is 8.7 °C.

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162 **4.0 Data Source and Analysis**

163 Surface meteorological data, including pressure, temperature and relative humidity, was
164 downloaded from Modern-Era Retrospective analysis for Research and Application (MERRA - 2). The
165 data project was stimulated by the recognition that various aspects of the hydrologic cycle represented
166 in previous generations of re-analyses were not adequate for climate and weather studies. MERRA – 2
167 offered an improvement upon the water cycle as a contribution to the science community and
168 reanalysis research (Colarco, *et al.*, 2010).

169 MERRA uses a three-dimensional variation (3D-Var) analysis algorithm based on the Grid-point
170 Statistical Interpolation scheme. Like other current re-analyses, it makes extensive use of satellite
171 radiance information, including data from hyper-spectral instruments such as the Atmospheric Infrared
172 Sounder (AIRS) on Aqua (Bacmeister, 2006). MERRA was processed in three separate streams, each
173 spun-up in two stages: Stream 1 for 1 January 1979 to 31 December 1992, followed by Stream 2 for 1
174 January 1993 to 31 December 2000, and then continues with Stream 3 for 1 January 2001 to the
175 present.

176 The MERRA meteorological data for six locations defined by different climatic conditions over
177 Nigeria, namely Yenagoa, Abakaliki, Ibadan, Anyigba, Jalingo and Sokoto, were downloaded for two
178 periods of 20 years, 1979 – 1998 and 1999 – 2018, on monthly averages. The dataset of atmospheric
179 temperature, T (K), atmospheric pressure, P (hPa) and relative humidity, Rh (%) were extracted using
180 text import wizard and appropriate delimited options to prepare excel spreadsheet format for data
181 manipulations. Radio refractivity (N) was computed from a set of T, Rh and P on monthly basis within
182 the period of 1979 – 1998 using ITU-R equation 3.0 with a view to deduce monthly averages for 20
183 years. Monthly values of t (°C), Rh and P were also separately estimated for the period covering 1979 –
184 1998 as shown in tables 1 - 3.

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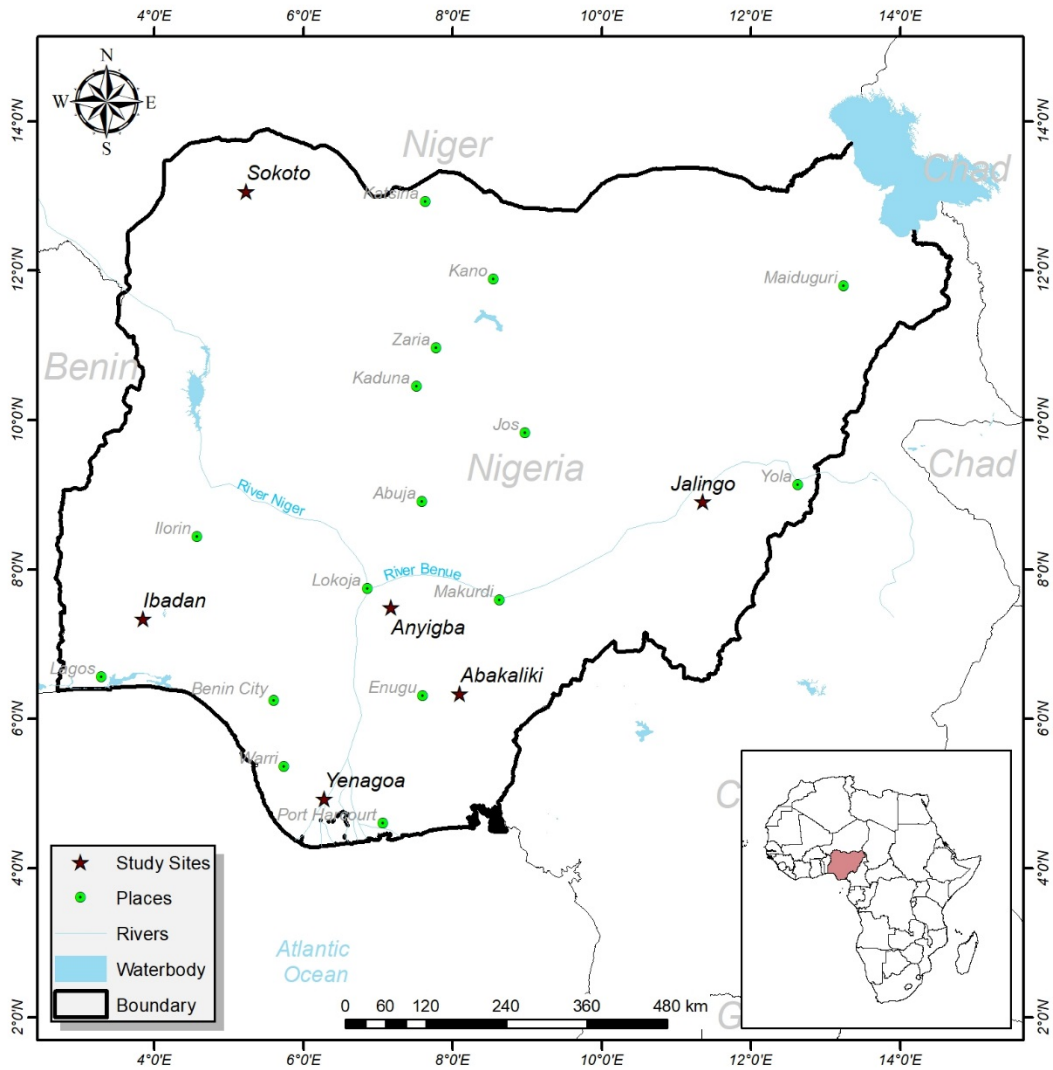


Figure 1: Map showing the Six Locations of the Study Areas over Nigeria

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Multiple linear regression was developed for each location to estimate radio refractivity directly from meteorological parameters. This was carried out using scattered plots of N and t ($^{\circ}\text{C}$); N and R_h ; N and R_h , t ($^{\circ}\text{C}$) with the aid of Origin graphing software to obtain a linear relationship between N and other meteorological parameters for different climatic zones across Nigeria. The linear expressions obtained for N from different regression analysis in the six climatic locations (as a function of t , R_h and t , R_h) and the ITU-R equation were applied on new set of monthly averages of meteorological data within a period of 1999 – 2018. Values of N obtained from linear regressions from each location were compared with values of N estimated from ITU-R equations. Correlation between radio refractivity values evaluated from linear expressions and ITU-R equation were computed to determine the degree of agreement and suitability of the new empirical relationships.

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Table 1: Temperature (K) at 2 m above ground averaged monthly over a period of 1979 – 1998)

Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	299.46	297.01	298.45	299.28	299.00	297.92
2	300.06	300.32	299.12	300.44	299.87	298.30
3	300.19	300.16	299.82	302.05	303.53	302.55
4	300.13	300.05	299.62	300.75	303.06	306.18
5	299.85	299.83	299.28	299.88	300.32	306.32
6	299.03	298.79	298.58	299.12	299.86	303.22
7	298.13	298.42	297.63	298.10	298.18	300.68
8	297.84	298.34	297.38	297.93	297.75	299.38
9	298.21	298.73	297.91	298.15	298.46	300.22
10	299.04	298.90	298.63	298.94	299.49	299.95
11	299.38	299.87	298.58	298.78	299.53	298.77
12	299.67	298.65	298.16	298.47	299.01	297.25
Monthly Mean	299.25	299.09	298.59	299.32	299.84	300.89

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Table 2: Atmospheric Pressure (hPa) at ground level averaged monthly over a period of 1979 – 1998)

Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	1008.29	1001.57	989.87	985.05	985.05	979.13
2	1007.86	999.28	989.53	984.69	984.69	979.52
3	1007.43	1000.41	988.88	983.80	983.8	976.51
4	1007.77	1000.42	989.25	984.28	984.28	975.40
5	1009.13	1001.31	990.66	985.68	985.68	976.39
6	1010.04	1003.09	991.51	986.48	986.48	977.88
7	1011.85	1003.61	993.03	988.04	988.04	979.54
8	1010.93	1003.60	992.10	987.08	987.08	978.78
9	1010.58	1001.75	991.83	986.91	986.91	978.88
10	1009.39	1001.49	990.83	985.81	985.81	978.54
11	1008.50	1000.43	990.10	985.19	985.19	978.68
12	1008.30	1001.80	990.07	985.28	985.28	979.60
Monthly Mean	1009.17	1001.56	990.64	985.69	976.01	978.24

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Table 3: Relative humidity (%) at 2 m above ground averaged monthly over a period of 1979 – 1998)

Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	82.31	54.61	76.42	58.85	38.77	18.94
2	80.13	75.93	76.77	45.73	21.27	11.03
3	86.59	80.79	82.67	64.66	26.1	13.15
4	86.85	81.80	85.16	73.63	55.25	20.31
5	87.69	82.89	87.26	81.23	76.07	42.54
6	89.43	86.18	87.35	82.80	76.1	56.97
7	89.20	86.25	88.52	84.86	83.91	70.96
8	89.67	85.61	88.49	85.79	85.97	79.72
9	89.67	85.76	87.97	85.81	83.24	74.31
10	88.04	85.96	87.02	82.62	70.67	46.15
11	85.66	79.24	83.09	74.25	49.28	24.15
12	81.55	64.03	76.31	62.47	35.25	24.78
Monthly Mean	86.40	79.09	83.92	73.56	58.49	40.25

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5.0 Results and Discussion

5.1 Derivation of Linear Meteorological Models for Surface Refractivity Estimation

The surface radio refractivity values estimated from existing ITU-R equation (quoted in equation 3) for six locations across Nigeria were presented in table 4. The results showed monthly averages of surface refractivity a period of 1979 – 1998. The spatial trend showed a significant decrease in monthly average of surface refractivity from southern location to the northern locations with the highest value of 383.41 N-Unit in Yenagoa and the least value of 314.81 N-Unit in Sokoto. Monthly distributions of refractivity values showed that high values were attained from March to October for Yenagoa, Abakaliki and Ibadan while high values were attained from May to October for Anyigba and Jalingo. Sokoto attained its high values from June to September.

The distributions were greatly influenced by wide variations in meteorological parameters most especially temperature (T) and relative humidity (Rh) along the latitudes across Nigeria. Temperature values reduced on monthly average in the south from 299.25 K in Yenagoa to 298.59 K at Ibadan while it increased in the north from 299.32 K at Anyigba to 300.89 K at Sokoto. Atmospheric pressure reduced south – north on monthly average from 1009.17 hPa at Yenagoa to 978.24 hPa at Sokoto. Also, relative humidity on monthly average reduced south – north from 86.40 % at Yenagoa to 40.25 % at Sokoto.

Linear regression analysis between the N values obtained from ITU-R in table 4 and meteorological parameters measured within the same period were carried out to obtain new expressions in terms of t ($^{\circ}$ C) and Rh for the six locations identified over Nigeria. Linear expressions 6a – 6c were obtained for Yenagoa as possible empirical models for estimating surface refractivity expressed in terms of relative humidity and temperature only (Eq. 6a); in terms of temperature only (Eq. 6b) and in terms of relative humidity only (Eq. 6c).

The regression analysis was repeated for other locations namely Abakaliki, Ibadan, Anyigba, Jalingo and Sokoto. The resulting linear expressions are presented subsequently. Linear expressions for Abakaliki were presented in equations 7a – 7c; expressions for Ibadan were presented in equations 8a –

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258 8c; expressions for Anyigba were presented in equations 9a – 9c; expressions for Jalingo were presented
 259 in expressions 10a – 10b; expressions for Sokoto were presented in 11a – 11c respectively.

260 Table 4: Surface Refractivity (N-Units) estimated from ITU-R Equation (N_{itu-r}) averaged over a period of
 261 1979- 1998
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Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	378.55	360.38	360.62	338.47	307.18	279.92
2	378.44	351.45	364.15	322.90	283.39	269.60
3	388.58	379.52	375.96	358.10	294.80	272.44
4	388.71	380.79	378.56	366.16	344.21	287.97
5	388.66	383.41	380.02	373.36	365.51	333.31
6	386.68	380.93	376.51	371.80	363.46	348.79
7	381.86	377.73	373.42	369.75	366.35	360.53
8	380.69	377.50	371.84	369.86	366.65	366.80
9	382.59	379.08	373.85	370.97	366.61	363.17
10	384.63	379.98	376.16	370.44	353.74	320.61
11	382.94	370.70	370.35	357.95	323.21	287.38
12	378.57	355.28	359.21	340.65	302.34	287.15
Monthly Mean	383.41	373.06	371.72	359.20	336.45	314.81

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The expressions showed linear dependence of surface refractivity on temperature and relative humidity. It was observed that the expressions had a pattern of highly varied numerical constants which makes the expressions significantly dependent on weather characteristics at specific location. The linear meteorological models are listed as follows:

$$N_{Rh,t}(Yenagoa) = 335.32 + 0.33Rh + 0.75t \quad 6a$$

$$N_t(Yenagoa) = 344.07 + 1.50t \quad 6b$$

$$N_{Rh}(Yenagoa) = 326.57 + 0.66Rh \quad 6c$$

$$N_{Rh,t}(Abakaliki) = 251.0 + 0.69Rh + 1.78t \quad 7a$$

$$N_t(Abakaliki) = 280.27 + 3.55t \quad 7b$$

$$N_{Rh}(Abakaliki) = 261.73 + 1.37Rh \quad 7c$$

$$N_{Rh,t}(Ibadan) = 285.68 + 0.6Rh + 1.39t \quad 8a$$

$$N_t(Ibadan) = 300.25 + 2.79t \quad 8b$$

$$N_{Rh}(Ibadan) = 271.1 + 1.2Rh \quad 8c$$

$$N_{Rh,t}(Anyigba) = 358.18 + 0.66Rh + 1.67t \quad 9a$$

$$N_t(Anyigba) = 446.8 - 3.33t \quad 9b$$

$$N_{Rh}(Anyigba) = 269.56 + 1.22Rh \quad 9c$$

290 $N_{Rh,t(Jalingo)} = 381.84 + 0.67Rh - 3.15t$ 10a
 291 $N_{t(Jalingo)} = 505.25 - 6.29t$ 10b
 292 $N_{Rh(Jalingo)} = 258.43 + 1.33Rh$ 10c
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 294 $N_{Rh,t(Sokoto)} = 381.84 + 0.67Rh - 3.15t$ 11a
 295 $N_{t(Sokoto)} = 505.25 - 6.29t$ 11b
 296 $N_{Rh(Sokoto)} = 258.43 + 1.33Rh$ 11c

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 298 **5.2 Performance of Linear Meteorological Model in Estimation of Surface Refractivity**

299 The results of monthly estimates of surface refractivity from new set of meteorological data for
 300 a period of 1999 – 2018 using equations 3, 6, 7, 8, 9, 10 and 11 are shown in tables 5 – 10. The values of
 301 surface refractivity averaged on monthly basis within the period of 20 years showed that surface
 302 refractivity had highest linear dependence on relative humidity in five out of the six locations. The five
 303 locations include Yenagoa, Abakaliki, Anyigba, Jalingo and Sokoto. Surface refractivity estimated from
 304 ITU-R equation was compared with values from linear meteorological models in equations 6 – 11 to
 305 determine the suitability of the expressions.

306 The mean values of surface refractivity from different linear meteorological models for the six
 307 locations are presented in table 11. In Yenagoa, ITU-R had a monthly average value of 381.87 N-Units
 308 while N_{Rh} had the closest value of 383.16 N-Units. In Abakaliki, ITU-R had a monthly average value of
 309 370.75 N-Units while N_{Rh} had the closest value of 369.35 N-Units. In Ibadan, ITU-R had a monthly
 310 average value of 370.02 N-Units while N_t had the closest value of 370.89 N-Units. In Anyigba, ITU-R had
 311 a monthly value of 358.74 N-Units while N_{Rh} had the closest value of 359.58 N-Units. In Jalingo, ITU-R
 312 had a monthly value of 338.77 N-Units while N_{Rh} had the closest value of 337.09 N-Units. In Sokoto,
 313 ITU-R had a monthly value of 319.09 N-Units while N_{Rh} had the closest value of 319.61 N-Units. The
 314 temporal and spatial difference in surface refractivity values estimated from linear meteorological
 315 models with highest dependence is given as ± 1.12 N-Units when compared with values from ITU-R.

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 321 Table 5: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Yenagoa for a
 322 period of 1999 – 2018
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Months	N_{itu-r}	$N_{Rh,t(Yenagoa)}$	$N_{t(Yenagoa)}$	$N_{Rh(Yenagoa)}$
1	363.77	379.11	382.64	375.58
2	383.28	383.13	384.86	381.41
3	386.42	383.99	384.23	383.75
4	386.54	384.03	384.17	383.90
5	386.89	384.23	383.72	384.75
6	385.46	384.21	382.62	385.81
7	382.87	383.80	381.99	385.60
8	380.17	383.30	381.54	385.05
9	383.83	384.07	382.19	385.96
10	385.49	384.17	383.01	385.32
11	386.98	384.06	384.39	383.73
12	370.75	380.45	383.85	377.05

Monthly Mean	381.87	383.21	383.27	383.16
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Table 6: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Abakaliki for a period of 1999 – 2018

Months	N_{itu-r}	N_{Rh,t}(Abakaliki)	N_t(Abakaliki)	N_{Rh}(Abakaliki)
1	330.03	331.42	365.51	335.82
2	371.35	352.02	377.26	365.02
3	378.02	355.09	376.69	371.68
4	378.92	355.59	376.30	373.07
5	379.55	355.95	375.52	374.56
6	379.05	356.37	371.82	379.07
7	377.31	355.76	370.51	379.16
8	376.03	355.18	370.23	378.29
9	377.81	355.97	371.61	378.49
10	378.93	356.41	372.22	378.77
11	374.23	353.50	375.66	369.56
12	347.73	340.84	371.33	348.72
Monthly Mean	370.75	352.01	372.89	369.35

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Table 7: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Ibadan for a period of 1999 – 2018

Months	N_{itu-r}	N_{Rh,t}(Ibadan)	N_t(Ibadan)	N_{Rh}(Ibadan)
1	342.73	359.37	366.68	352.28
2	371.04	371.00	373.52	368.73
3	375.60	372.78	373.65	372.15
4	377.65	373.71	373.49	374.18
5	377.85	374.16	372.12	376.45
6	375.43	373.55	370.33	377.01
7	373.73	373.06	369.41	376.94
8	371.97	372.69	368.21	377.40
9	375.22	373.81	369.80	378.06
10	376.35	373.69	371.59	376.03
11	371.36	371.24	372.71	370.03
12	351.26	362.97	369.16	357.01
Monthly Mean	370.02	371.00	370.89	371.36

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342 Table 8: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Anyigba for a
343 period of 1999 – 2018
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Months	N_{itu-r}	$N_{Rh,t(Anyigba)}$	$N_{t(Anyigba)}$	$N_{Rh(Anyigba)}$
1	310.60	346.72	367.68	321.72
2	354.19	355.95	354.96	350.57
3	363.88	359.09	354.09	357.18
4	364.82	359.42	354.03	357.86
5	368.71	364.15	356.72	364.10
6	370.20	370.81	361.79	371.71
7	369.46	372.34	363.35	373.10
8	369.19	372.73	363.75	373.44
9	370.60	372.72	363.02	374.10
10	371.87	370.76	360.95	372.39
11	361.19	363.48	359.02	360.73
12	330.12	353.96	364.45	338.10
Monthly Mean	358.74	363.51	360.32	359.58

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350 Table 9: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Jalingo for a
351 period of 1999 – 2018
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Months	N_{itu-r}	$N_{Rh,t(Jalingo)}$	$N_{t(Jalingo)}$	$N_{Rh(Jalingo)}$
1	283.17	321.98	355.74	288.24
2	307.86	312.42	319.95	304.85
3	322.58	312.11	310.83	313.29
4	341.74	320.35	312.27	328.20
5	359.84	341.17	329.70	352.21
6	362.29	348.49	337.31	359.18
7	365.60	357.57	345.92	368.63
8	364.25	358.98	348.25	369.13
9	366.15	357.76	345.61	369.34
10	362.14	348.04	336.68	358.91
11	327.98	330.85	334.67	326.79
12	301.62	328.06	349.64	306.36
Monthly Mean	338.77	336.48	335.55	337.09

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355 Table 10: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Sokoto for a
 356 period of 1999 – 2018

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Months	N_{itu-r}	$N_{Rh,t(Sokoto)}$	$N_t(Sokoto)$	$N_{Rh(Sokoto)}$
1	276.13	289.67	300.32	278.81
2	283.63	300.08	314.14	285.76
3	273.52	297.91	320.64	274.88
4	290.79	306.81	326.21	287.09
5	331.74	321.98	326.31	317.33
6	347.40	329.10	320.51	337.40
7	355.37	334.67	315.33	353.74
8	363.85	341.26	311.48	370.79
9	365.93	342.13	311.97	372.04
10	353.01	334.84	313.03	356.39
11	298.71	306.43	307.20	305.44
12	289.03	299.13	302.38	295.66
Monthly Mean	319.09	317.00	314.13	319.61

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364 Table 11: Mean Surface Refractivity Values from ITU-R and Linear Meteorological Models for a period of
 365 1999 - 2018

Stations	N_{itu-r}	$N_{Rh,t}$	N_t	N_{Rh}
Yenagoa	381.87	383.21	383.27	383.16
Abakaliki	370.75	352.01	372.89	369.35
Ibadan	370.02	371.00	370.89	371.36
Anyigba	358.74	363.51	360.32	359.58
Jalingo	338.77	336.48	335.55	337.09
Sokoto	319.09	317.00	314.13	319.61

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368 5.3 Validation and Comparison of Linear Meteorological Models with ITU-R Equation

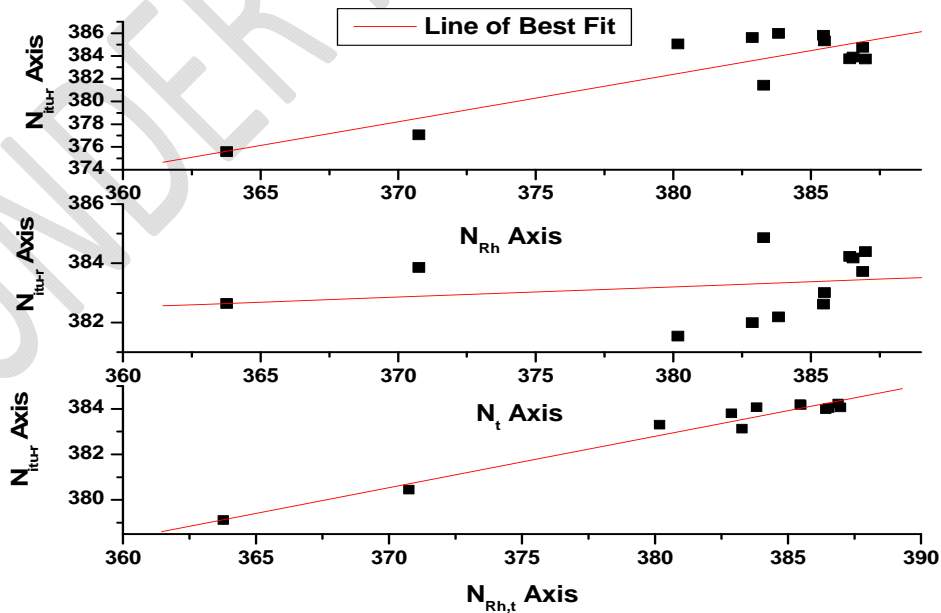
369 The results of surface refractivity estimated from different linear meteorological models were
 370 compared in scattered plots to establish levels of disparity in the values obtained. The graphs are
 371 presented in figures 2 – 7 showing the line of best fit. Surface refractivity values from N_{itu-r} was
 372 compared with values obtained from N_{Rh} , N_t and $N_{Rh,t}$ being the contributions of relative humidity,
 373 temperature and combination of relative humidity and temperature to linear meteorological models.
 374 This was to determine the correlation coefficient and the standard deviation of the surface refractivity

375 from ITU-R results at six locations – 3 from the south and 3 from the north – over Nigeria and the results
 376 are shown in table 12.

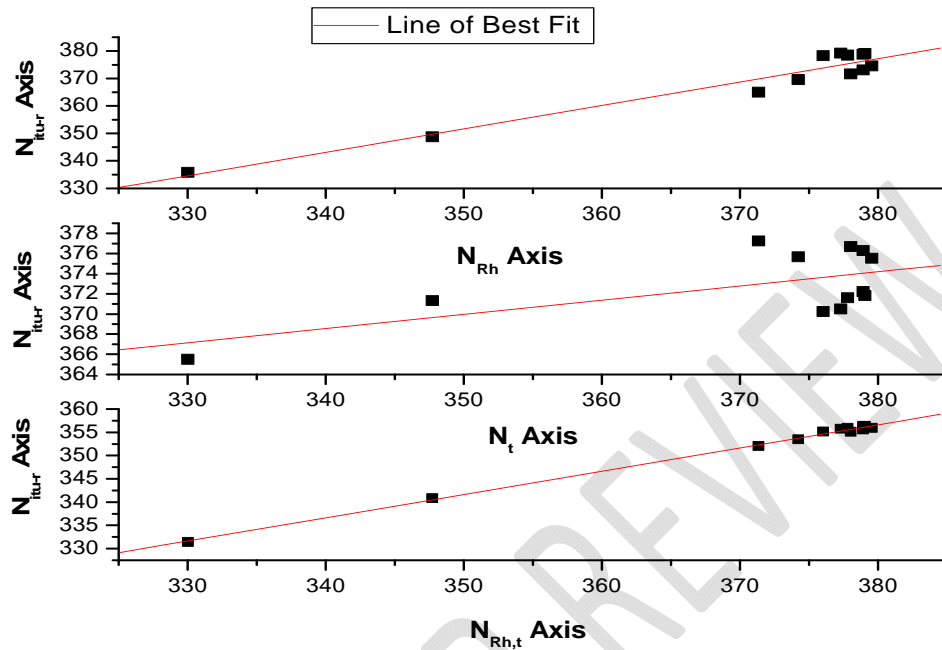
377 In Yenagoa, contributions of Rh and t ($N_{Rh,t}$) had the best linear correlation of 0.98 and standard
 378 deviation of ± 0.32 while temperature contribution (N_t) had the least correlation of 0.23 and standard
 379 deviation of ± 1.09 . The contributions of relative humidity (N_{Rh}) had a dominance on temperature with
 380 correlation of 0.88 and standard deviation of ± 1.73 . Results from Abakaliki and Ibadan were similar to
 381 that of Yenagoa, correlation of $N_{Rh,t}$ with N_{ITU-R} was 0.99 at Abakaliki and Ibadan while the standard
 382 deviation was ± 0.44 and ± 0.49 respectively. Similarly, N_{Rh} had a dominance over N_t having correlation
 383 of 0.97 and 0.95 at Abakaliki and Ibadan while N_t had correlation of 0.62 and 0.65 respectively. High
 384 relative humidity and low temperature values observed in the 3 southern stations – Yenagoa, Abakaliki
 385 and Ibadan – significantly enhanced the suitability of $N_{Rh,t}$ linear meteorological model for estimation of
 386 surface refractivity values comparable to N_{ITU-R} .

387 The remaining 3 stations in the north – Anyigba, Jalingo and Sokoto – had divergent trend from
 388 the southern stations. Surface refractivity values obtained from the contributions of temperature (N_t)
 389 was extremely low and mostly negative. Results from N_t at Anyigba, Jalingo and Sokoto had correlation -
 390 0.46, -2.99 and 0.17 while the standard deviation was ± 4.21 , ± 15.53 and ± 8.64 respectively. Contrary to
 391 the observations in the southern stations, in the northern stations, relative humidity contributions (N_{Rh})
 392 had the best linear correlation having 0.96 at Anyigba, 0.98 at Jalingo and Sokoto. Standard deviation
 393 was given as ± 4.71 at Anyigba, ± 6.39 at Jalingo and ± 7.17 at Sokoto. The trend in the north was as result
 394 of low relative humidity and high temperature.

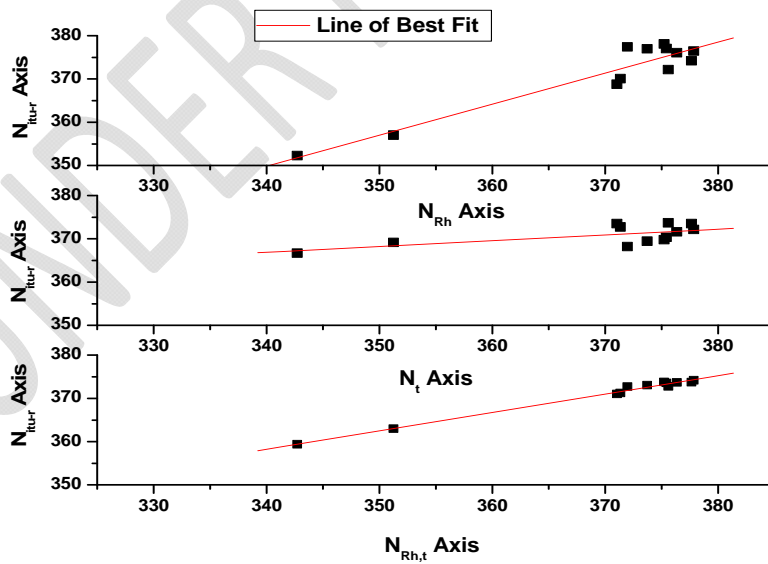
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401 Figure 2: Correlation between ITU-R and Linear Meteorological Models Values for Yenagoa
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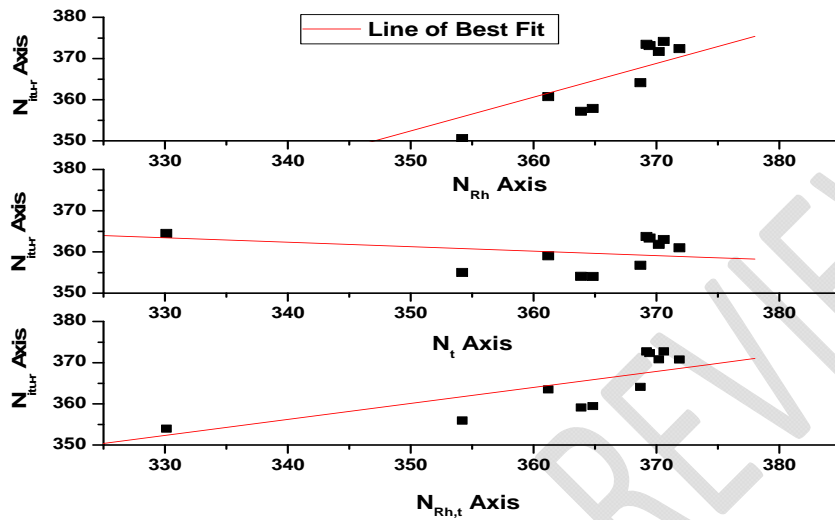


403 Figure 3: Correlation between ITU-R and Linear Meteorological Models Values for Abakaliki
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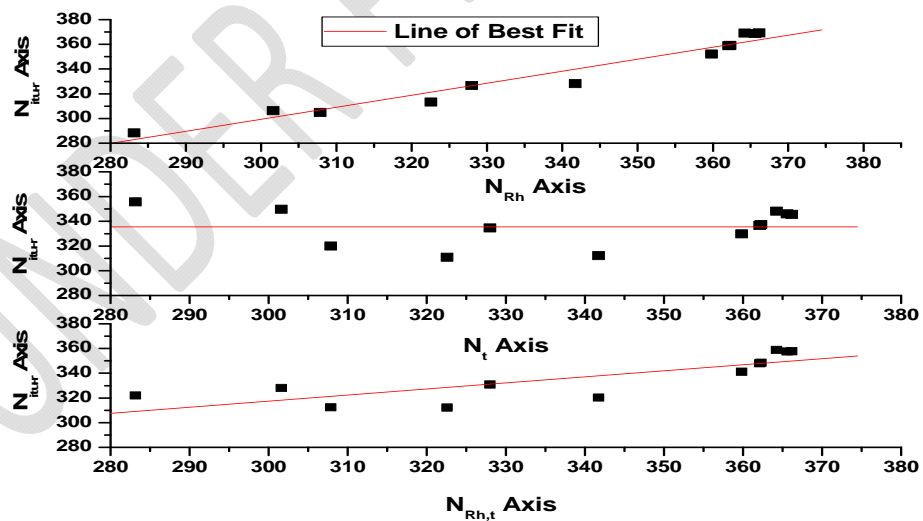
406 Figure 4: Correlation between ITU-R and Linear Meteorological Models Values for Ibadan
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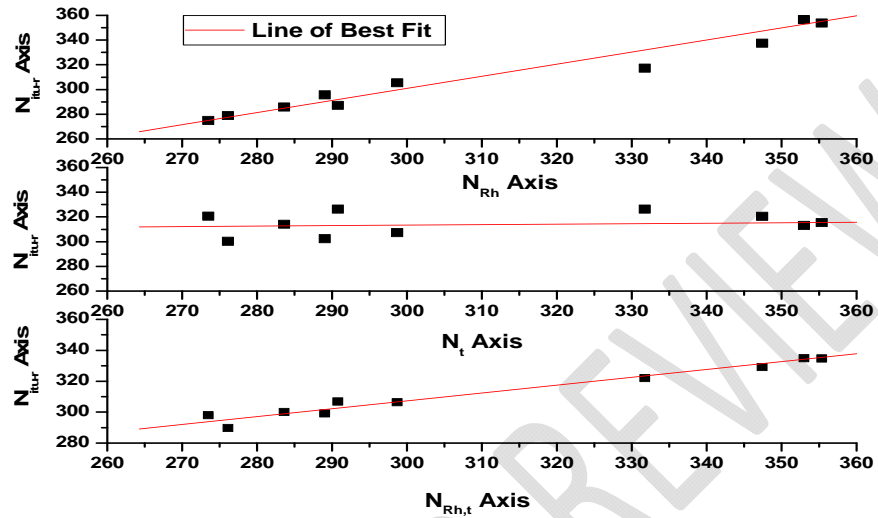
Figure 5: Correlation between ITU-R and Linear Meteorological Models Values for Anyigba



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Figure 6: Correlation between ITU-R and Linear Meteorological Models Values for Jalingo

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Figure 7: Correlation between ITU-R and Linear Meteorological Models Values for Sokoto

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Table 12: Correlation Coefficient (r) and Standard Deviation (σ) between Surface Refractivity Values estimated from ITU-R and Linear Meteorological Models

Stations	$N_{Rh,t}$		N_t		N_{Rh}	
	r	σ	r	σ	r	σ
Yenagoa	0.98	0.32	0.23	1.09	0.88	1.73
Abakaliki	0.99	0.44	0.63	2.83	0.97	3.52
Ibadan	0.99	0.49	0.65	1.84	0.95	2.78
Anyigba	0.86	4.64	-0.46	4.21	0.96	4.71
Jalingo	0.81	10.83	-2.99	15.53	0.98	6.39
Sokoto	0.99	2.91	0.17	8.64	0.98	7.17

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6.0 Conclusion

The problem of quantifying surface refractivity as a function of atmospheric variables in various climatic situations is very complex because the three meteorological variables involved are neither perfectly correlated nor totally uncorrelated. In this study, new sets of linear expressions were obtained in terms of meteorological parameters from six locations across Nigeria as alternate means of estimating surface refractivity. The linear expressions were validated using data obtained for a period of 20 years and compared with results from existing equations. Linear expressions as a function of relative humidity and temperature were best suitable for stations in the south with minimum correlation of 0.98 while linear expressions as a function of relative humidity only were best suitable for the stations in the north with minimum correlation of 0.96.

The sensitivity of the parameters showed that surface refractivity is much more likely to vary due to moisture changes in the atmosphere than temperature changes. In the northern Nigeria, relative humidity changes near the surface have more than twice the impact on refractivity as temperature changes. This is partly a result of the high variability of humidity and partly due to the inherent sensitivity of refractivity to moisture in the additional wet term, which is also a maximum in wet months.

At low temperatures, surface refractivity varies very little with relative humidity because the saturation vapour pressure is low: as air temperatures increase, the saturation vapour pressure rises, expanding the range of possible refractivity values. Hence, estimation of surface refractivity from linear meteorological models is mainly a function of low temperature and high relative humidity in the south and a function of relative humidity in the north.

7.0 References

1. Amajama, J., and Eshiet, M. (2016). Impact of weather components on (UHF) radio signal. *International Journal of Engineering Research and General Science (IJERGS)*, 4(3), 474-480.
2. Joseph, A. (2015) Mathematical Relationships between Radio Refractivity and Its Meteorological Components with A New Linear Mathematical Equation to Determine Radio Refractivity. *International Journal of Innovative Science, Engineering & Technology (IJSET)*, Vol. 2 Issue 12, December, 2015
3. Jari, L., and Ismo, H. (2015). Effect of temperature and humidity on radio signal strength in outdoor wireless sensor. In *Proceedings of the federated conference on computer science and information systems (Vol. 5, pp. 1247-1255)*.
4. Enyenihi H. J., Simeon Ozuomba, Kalu Constance (2019): Development of Model for Estimation of Radio Refractivity from Meteorological Parameters. *Universal Journal of Engineering Science* 7(1): 20-26, DOI: 10.13189/ujes.2019.070103
5. ITU-R (2013): "Propagation data and prediction methods required for the design of terrestrial line-of-sight systems," Geneva, July 2013, pp. 530
6. Stephen J. S., Hedley Hansen and Derek Abbott (2014): Prediction of Surface Refractivity Gradient Distributions, from Weather Station Surface Data. Conference Paper, <https://www.researchgate.net/publication/272093235>; DOI: 10.1109/EuCAP.2014.6901750

- 484 7. Ayantunji, B. G. and Okeke, P. N. (2011): Diurnal and Seasonal Variation of Surface Refractivity
485 over Nigeria. Progress in Electromagnetics Research B, Vol. 30, 201 – 222, 2011
486
- 487 8. ITU-R (2003): The radio refractive index: Its formula and refractivity data," pp. 453 – 459
488
- 489 9. Odjugo, P. A. O. (2013): Analysis of climate change awareness in Nigeria. Scientific Research and
490 Essays, Vol. 8 (26), 1203 – 1211.
491
- 492 10. Nwagbara, M. O., Ijeoma, M. A., Chima, G. N. (2009): Climate change and flood in Northern
493 Nigeria. An examination of Rainfall Trends over the region". In R. N. C. Anyadike, I. M. Madu and
494 C. K Ajaero, Eds. Climate change and the Nigeria Environment, pp. 525 –538.
- 495 11. Maria D. S., Diego, B. and Giovanni, S. (2013): Measuring the effect of climate change on
496 agriculture: A literature review of analytical models. Journal of Development and Agricultural
497 Economic Vol. 5 (2), pp. 499 - 509.
498
- 499 12. James, F. (2009): Climate change: How o report the story of the century.
500 [http://www.scidev.met/ent. Practical guide/ climate. Change –how-to-report-the- storyof–the](http://www.scidev.met/ent. Practical guide/ climate. Change –how-to-report-the- storyof–the cent.html)
501 [cent.html](http://www.scidev.met/ent. Practical guide/ climate. Change –how-to-report-the- storyof–the cent.html).
502
- 503 13. Abu B (2007): Sea level rise and the Niger Delta of Nigeria. Journal of Wetland, 3(1): 44-52
504
- 505 14. Intergovernmental Panel on Climate (IPCC) (2007): Climate change 2007. The fourth assessment
506 report (AR4). Synthesis report for policy makers. [Http://www.ipcc.ch/pdf/assessment-](Http://www.ipcc.ch/pdf/assessment-report/ar4/syr/ar4_syr_spm.pdf)
507 [report/ar4/syr/ar4_syr_spm.pdf](Http://www.ipcc.ch/pdf/assessment-report/ar4/syr/ar4_syr_spm.pdf). (Access 10th August, 2009).
508
- 509 15. Jagtap, S (2007): Managing vulnerability to extreme weather and climate events:
510 Implications for agriculture and food security in Africa. Proceedings of the Inter-national
511 Conference on Climate Change and Economic Sustainability held at Nnamdi Azikiwe
512 University, Enugu, Nigeria. 12-14, June 2007.
513
- 514 16. Ayoade, J. O (2004): Climate Change. Ibadan. Vantage Publishers, pp. 45-66
515
- 516 17. Nwafor J. C., (2007): Global climate change: The driver of multiple causes of flood intensity
517 in Sub-Saharan Africa. Paper presented at the International Conference on Climate Change
518 and Economic Sustain-ability held at Nnamdi Azikiwe University, Enugu, Nigeria, 12-14
519 June 2007.
520
- 521 18. Ayuba H. K, Maryah U. M, and Gwary D. M (2007): Climate change impact on plant species
522 composition in six semi-arid rangelands of Northern Nigeria. Nigerian Geographical Journal
523 5(1): 35-42
524
- 525 19. Colarco, P. R., A. da Silva, M. Chin, and T. Diehl, (2010): Global Aerosol Distributions in the NASA
526 GEOS-4 Model and comparisons to Satellite and Ground-based Aerosol Optical Depth. Journal of
527 Geophysical Research, 115, DOI: 10.1029/2009JD012820
528

529 20. Bacmeister, J. T., Suarez, M. J. and Robertson, F. R. (2006): Rain Re-evaporation, Boundary-
530 Layer/Convection Interactions and Pacific Rainfall Patterns in an AGCM, Journal of Atmospheric
531 Science, 63, pp. 3383-3403.
532

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