

1 **EDAPHIC FACTORS AND FLOODING PERIODICITY DETERMINING**
2 **FOREST TYPES IN A TOPOGRAPHIC GRADIENT IN THE NORTHERN**
3 **BRAZILIAN AMAZONIA**

4
5
6 **Abstract**
7

8 The Brazilian Amazonia is a region covered by an extensive mosaic of tropical forests
9 conditioned by different topographical and hydro-edaphic features. Although studies
10 relating environmental determinants of structure and floristic composition are
11 systematically evolving in the region, there is no doubt that there are still information
12 gaps due to the lack of research in peripheric areas of the Amazonia. The seasonally
13 flooded areas of the state of Roraima situated on rio Branco-rio Negro basin, northern
14 Brazilian Amazonia, still are deprived of such information. In this way, this work had as
15 objective determine the physical and soil chemical attributes, and the flooding
16 periodicity that characterize different forest types dispersed in a topographic gradient
17 located in a area on the north of rio Branco-rio Negro basin. Soil samples (0-60 cm)
18 were collected along a 2.7 km transect (31.1-64.8 m a.s.l.) crossing three different forest
19 types: (i) mosaic between treed and forested shade-loving (La+Ld), (ii) area of
20 ecological tension between forested shade-loving and open ombrophilous forest (LO)
21 and (iii) open ombrophilous forest (Ab+As). The results indicated different soil classes
22 and flooding periodicity for each forest type observed: Entisols Fluvents (La+Ld, 3-4
23 months flooded), Entisols Quartzipsamments (LO, 1-2 months) and Yellow Ultisols
24 (Ab+As, no flooding). All analyzed soils were defined as nutrient-poor areas, especially
25 those located on low altitude, characterized for higher hydrological restrictions
26 (seasonal flooding) aggregating forest types of lower structural patterns (e.g. La+Ld).
27 Soils on low altitude were also characterized as those with the highest percentage of fine
28 sand and silt, while soil free of seasonal flooding (Yellow Ultisols) presented the
29 highest levels of clay and coarse sand, always associated with the ombrophilous forests
30 (higher structural patterns). These results improve our understanding of the
31 environmental factors conditioning different forest types in this peripheral region of
32 Amazonia, suggesting that ecosystems with higher hydro-edaphic restrictions are a
33 strong indicator of forest types with lower structural patterns.

34 Keywords: oligotrophic ecosystems, water table, ecotone, phytophysiognomy.
35

36 **INTRODUCTION**

37 The Amazon basin occupies ~40% of the surface of South America and about 60%
38 is inserted within the Brazilian territory (Carvalho & Domingues, 2016). Its
39 predominant covering is defined as dense and open ombrophilous forests (Veloso et al
40 1991, Fearnside, 2018), but throughout the Amazonian biome there are many different
41 forest types that may be distinguished by their floristic composition and structure due to
42 the large environmental heterogeneity (Fearnside & Ferraz, 1995; Nogueira et al. 2015).
43 In general, the factors modeling the different forest types are attributed mainly to the

44 climatic variations, hydro-edaphic conditions, topography, and anthropogenic
45 interferences, all interacting and acting at different spatial scales (Castilho et al., 2006;
46 Phillips et al., 2003; Laurance et al., 1999; Franco-Moraes et al., 2019). This
47 congregation of factors generates different structural and floristic shades with different
48 ecosystem values, but generally information on the weight of each one, in the local and
49 regional context, is little known due to the gigantism of the Amazonia, which makes
50 sampling in peripheral areas a difficult process (Philips et al, 1998; Fearnside. 2008).
51 However, this kind of information is the basis for improving our knowledge on
52 specialization of the various Amazonian phytophysiognomies, with different structural
53 patterns and species diversity, directly influencing the estimates of biomass/carbon
54 stocks and fluxes in this region considered to be the largest and most important "natural
55 environment" mitigating the harmful effects of global climate change (ter Steege et al.
56 2016; Lewis et al., 2004; IPCC. 2006)

57 Although studies involving edaphic and hydrological factors in association with
58 topographic gradients related to the dynamics of ecosystems have evolved rapidly in the
59 Amazon (e.g. Silva et al., 2016; Tuomisto et al. 2003; Tuomisto et al 2014), there is still
60 a great lack of information due to huge regional gaps with rare scientific investigations.
61 In this context, ecotone forests (transition areas, contact zones or areas of ecological
62 tension) are ecoregions representing about 15% of the biome (Santos et al. 2007), but
63 still have a lack of knowledge about the processes of their formation and maintenance
64 (Santos et al., 2013). These ecoregions are characterized by mosaics of different forest
65 types condensed into distinct spatial scales that hamper their floristic and structural
66 characterization and, above all, biomass/carbon estimates (Nascimento et al., 2014;
67 Barni et al., 2016). This scarcity is mainly detected in the northern of the Brazilian
68 Amazonia, especially in the seasonally flooded areas of the rio Branco-rio Negro basin
69 located in the state of Roraima (Silva et al., 2016; Mendonça et al., 2013, Damasco et
70 al., 2013).

71 Recent works have demonstrated that the ecosystems which form the mosaic of
72 landscapes in this region present a direct integration between the plant cover and the
73 physical, chemical and biological attributes of the soil (Mendonça et al., 2017), due to
74 essential processes related to the biogeochemical cycles, water table outcrops,
75 accumulation and decomposition of organic matter (Cordeiro et al. 2016; Silva et al.,
76 2016). This is a strong indication that environmental conditions associated with

77 temporal flooding, sediment drag, and nutrient leaching may have high importance in
78 the formatting of different forest types in this peripheral region of the Amazonia (Junk
79 et al., 2015; Suwa et al. 2013; Luizão et al. 2007). In the same context, the physical and
80 chemical attributes of the soil, altitude, flooding periodicity, drainage and microclimate
81 also ca considered as important determinants in the formation of different natural
82 environments with specific structural patterns (Khorramdel et al, 2013; Scaranello et al.
83 2012)

84 The objective of this study was to determine the flooding periodicity and the
85 physical and soil chemical attributes that characterize a topographic gradient established
86 between the Água Boa do Univini River and the Cumarú Mountain, a peripheral area of
87 the Rio Branco-Rio Negro basin. This region belongs to the Serra da Mocidade National
88 Park, a federal protected area located in the state of Roraima, northern Brazilian
89 Amazonia. This region is formed by a large mosaic of forest and non-forest ecosystems
90 without rare scientific investigations about the role of hydro-edaphic factors as
91 determinants of different phytophysiological formations. Our results aim to improve
92 the understanding of the environmental factors that determine different forest types in
93 this region of Amazonia, indicating the association between environments with
94 higher/lower hydro-edaphic restrictions and their respective forest structural pattern
95 taking into account horizontal (stem diameter) and vertical (total height) parameters.

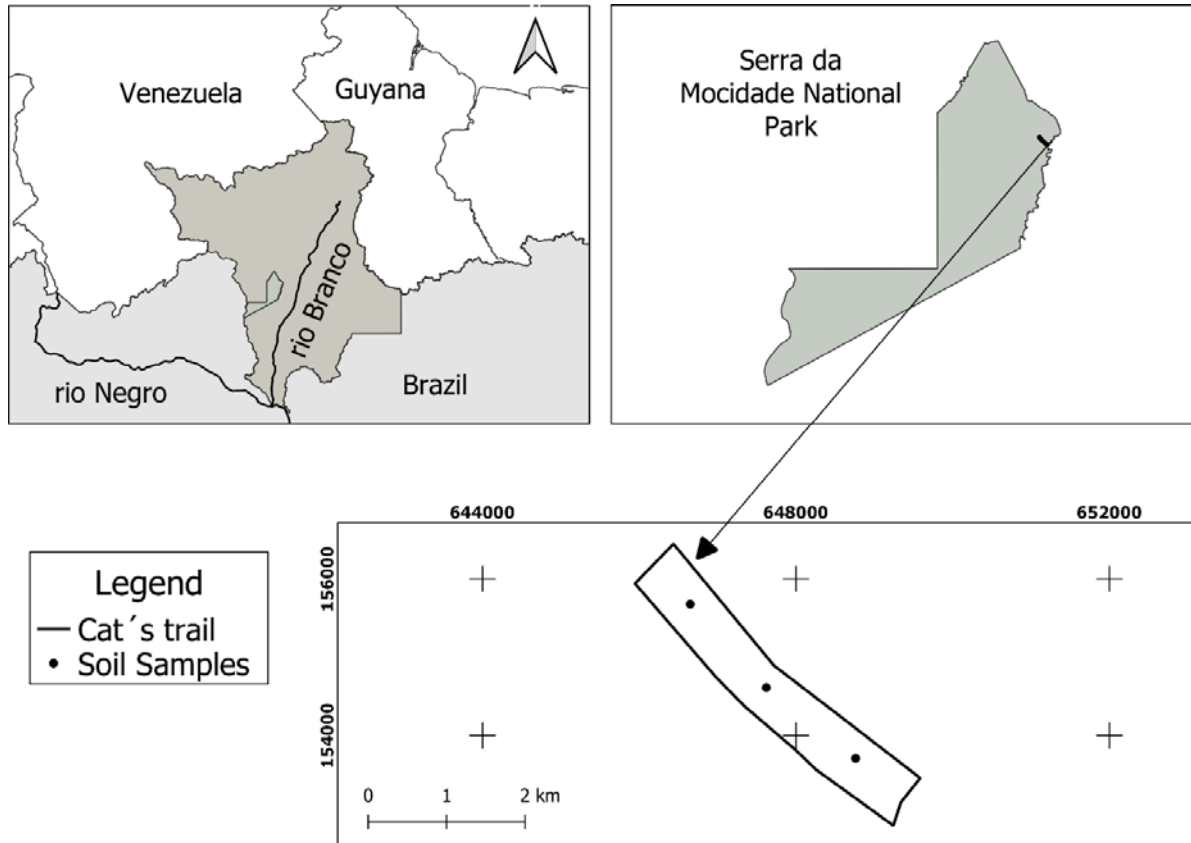
96 MATERIALS AND METHODS

97 Study area

98 This study was carried out in the Serra da Mocidade National Park (1.405° N -
99 61.648° W and 1.382° N - 61.673° W), a federal protected area managed by ICMBio
100 (Chico Mendes Institute for Biodiversity and Conservation), located in the municipality
101 of Caracaraí, ~ 290 km south of Boa Vista, capital of Roraima. The sampling area is
102 characterized by ecotone zones (area of ecological tension or transition area or contact
103 zone) between seasonally flooded forests under strong influence of the Água Nova do
104 Univini river (black water) in association with open ombrophilous forests that reach the
105 first steps of elevation of the Cumarú Mountain (Figure 1). The phytophysiological
106 characterization, edaphic and flood periodicity were performed in a topographic
107 gradient (31.1-64.8 m a.s.l.) inserted in an irregular transect (2.7 km long) known as
108 "Cat's Trail", which begins on the right bank of the Água Boa do Univini river to the

109 first steps of the northeast sector of the Cumaru Mountain, located at the northeastern
110 end of the National Park (Figure 1).

111 **Figure 1** - Study area indicating the geographical location of the Serra da Mocidade
112 National Park and the soil profiles sampled along the Cat's Trail.



113 Characterization of the study area

114 The National Park is totally inserted in the rainy tropical monsoon type (Am)
115 following Köppen classification, with annual rainfall of 1700-2000 mm and May-
116 August representing the rainiest period (~40% of annual rainfall) (Barbosa, 1997). This
117 region is marked by a chain of mountains that lends its name (Serra da Mocidade),
118 resulting from the erosion of the Guyanese Craton, a large continental block formed by
119 magmatic and metamorphic rocks dated between 1.8-2.5 billion years, in the Lower Pre-
120 Cambrian period (BRAZIL-MME, 1975). The characteristics of the main soil types in
121 the National Park region are defined from lithological residues of the same geological
122 constitution of the rock formation complex of the Serra da Mocidade, being a large
123 residual mass, with an altitude reaching ~1.800 m, characterized by sharp crests and
124 ravine slopes, covered by high altitude forests that lose this characteristic when reaching
125 the zones of flooded forests of low altitude and smaller biometric structural patterns.

126 Along this rocky complex, eight different soil classes can be found: Neossolo Flúvico
127 (Entisols Fluvents), Neossolo Quartzarênico (Entisols Quartzipsamments), Espodossolo
128 Humilúvico (Spodosols), Latossolo Amarelo (Yellow Oxisols), Gleissolo Háplico
129 (Entsols), Latossolo Vermelho (Red Oxisols), Neossolo Litólico (Entsols Lithic) and
130 Argissolo Amarelo (Yellow Ultisols) (BRAZIL-MME, 1975; USDA, 1999).

131 **Phyto-characterization and periodicity of flooding**

132 In order to carry out the phytophysionomic classification of the forest types in
133 the sample area, we adopted the criteria proposed by the Brazilian Vegetation
134 Classification System (BRAZIL-IBGE, 2012), based on a forest inventory
135 carried out along the entire transect at Cat's Trail (R. I.
136 Barbosa, personal communication). In this survey, the
137 structure (horizontal and vertical) and arboreal groups (trees
138 and palms) with stem diameter ≥ 10 cm were defined in each
139 forest type arranged along the topographic gradient (31,1–64, 8
140 m a.s.l.). The flood periodicity data were obtained from
141 observations performed in two consecutive rainy periods (2016
142 and 2017), where the sampling transect was coursed from start
143 to finish in both periods, estimating a mean time interval
144 (months) of flooding for each of them. All information was
145 aggregated into an ecosystem conceptual model that faithfully
146 followed the observed topographic gradient. Later, this model
147 was adopted as an associative basis of the chemical and soil
148 physical characteristics under each defined forest type.

149 **Soil sampling and physical/chemical analyzes**

150 In order to analyze and describe the physical and chemical attributes of the soil,
151 three profiles (1m wide, 1m long, 80cm deep) were opened for each forest type
152 considered. Soil samples were collected at 0-20cm, 20-40cm and 40-60cm depths. After
153 that, the samples were deposited in plastic bags and identified by forest type and depth.
154 All samples were air dried (TFSA), sieved (2 mm) and sent to the Soil Laboratory for
155 physical (% sand,% silt and% clay) and chemical (pH, organic matter, exchangeable

156 acidity, potential acidity, Ca, Mg, K, P, Cu, Zn, Fe, Mn and B) analysis following the
157 methodology specified by [EMBRAPA \(2009\)](#). The descriptive classification of the soils
158 sampled was performed by the Brazilian Soil Classification System ([Santos et al., 2018](#))
159 up to the third categorical level and correlated to the Soil Taxonomy ([USDA, 1999](#)).

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161 RESULTS AND DISCUSSION

162 The three profiles open along the sampled transect of the Serra da Mocidade
163 National Park delimited three different types of soils: (i) Entisols Fluvents, (ii) Entisols
164 Quartzipsamments and (iii) Yellow Ultisols (Figure 2). All of them were characterized
165 as chemically poor (Table 1) and with high sand contents (Table 2). The results found in
166 the forest type classified as Treed and Forested shade-loving (La + Ld), 3-4 month
167 flooded in the year, are associated with the Entisols Fluvents, while Entisols
168 Quartzipsamments and Yellow Ultisols are associated, respectively, to the Area of
169 Ecological Tension (LO) and Open Ombrophylous Forest (Ab+As) (Table 3). These
170 hydro-edaphic characteristics are similar to those reported in the Viruá National Park by
171 [Mendonça et al. \(2013\)](#) and [Damasco et al. \(2013\)](#), a region with similar ecological
172 characteristics to the Serra da Mocidade, strongly indicating that soils with higher levels
173 of sand, low nutrient contents and higher flooding periodicity can define oligotrophic
174 forest types with lower structural patterns (vertical and horizontal), in counterpoint to
175 areas free of flooding processes, such as ombrophilous forests (Figure 3).

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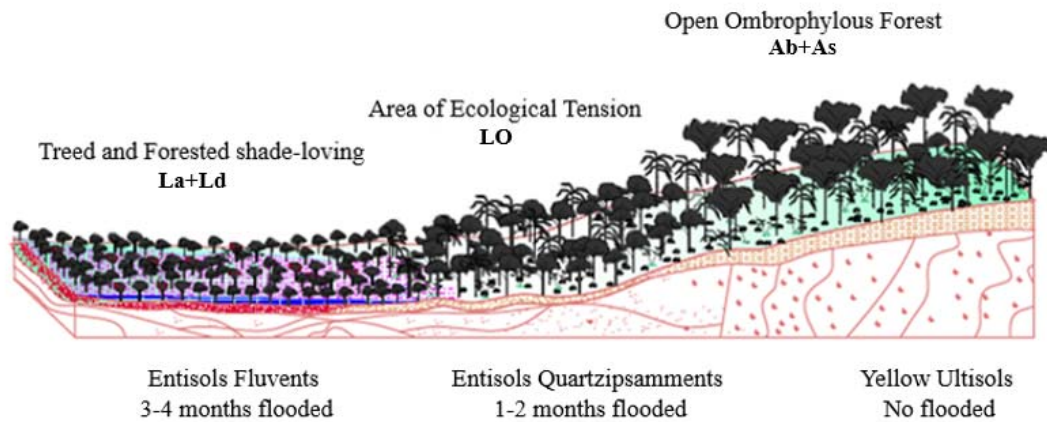


177 **Figure 2- Soil profiles opened in the Serra da Mocidade**
178 **National Park.**

179 Where the profiles area: **A**= Entisols Fluvents (Treed and Forested shade-loving La+Ld) **B**= Entisols
180 Quartzipsamments (Area of Ecological Tension LO) and **C**= Yellow Ultisols (Open Ombrophylous
181 Forest Ab+As).

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183 **Figure 3-** Conceptual model of the three forest types characterized by the soil classes
 184 and flooding periodicity in a transect in the Serra da Mocidade National Park,
 185 northern Brazilian Amazonia.



186 Entisols Fluvents can be considered as a low developed soil formed by
 187 quaternary sediment deposits, where drainage varies from moderate to imperfect, and it
 188 is very influenced by the outflow of the water table (Vale Jr. and Schaefer, 2010). In the
 189 study area, this soil class was characterized by the predominance of high levels of fine
 190 sand followed by the silt (Table 2), associated with mosaics of treed and forested shade-
 191 loving, which are oligotrophic forest environments with low structural expression that
 192 naturally dominate part of the rio Branc-rio Negro basin (Prance & Schubart, 1977;
 193 Barbosa & Ferreira, 2004; Mendonça et al., 2014). The pH in this environment showed
 194 a tendency to reduce as a function of soil depth, due to the high levels of fine sand,
 195 which facilitates the processes of water percolation and nutrient leaching (Quesada et
 196 al., 2009). The available phosphorus also presented a decrease with depth, being this a
 197 process associated to alluvial sediment entrainment condition of the high parts of the
 198 topographic gradient, considering that the parts of lower altitude are characterized by
 199 more restrictive soils (3-4 months flooded) that may eventually accumulate higher
 200 levels of fertility in the early layers in contrast with deeper layers (Suwa et al., 2013).

201 Our sampling in the profile of this soil class indicated that the organic matter had
 202 the highest values in the first sampling layer (0-20 cm). This result is in agreement with
 203 that presented by Ferraz et al (1998) and Luizão et al. (2007), where the authors suggest
 204 that soils with higher elevations tend to have higher clay contents compared to lowland
 205 soils, but the organic matter contents are higher in the first layers of lowland soils due to
 206 the strong drag provided by the topography of terrain. Concerning to the physical
 207 properties of this soil class situated in the La+Ld forest type, it was verified higher

208 concentrations of silt and fine sand, which are textural particles fully compatible with
209 natural lowland environment (Magalhães and Gomez, 2013).

210 Entisols Quartzipsamments is characterized as a soil under Area of Ecological
211 Tension (LO), 1-2 months flooded in the year, and it can be described as a soil class
212 extremely weathered, much quartzous, where almost all the clay is destroyed by
213 acidolysis, or sandy deposits formed by wind phenomena, occurring in flat reliefs or
214 basin reliefs or even in soft undulating reliefs (Vale Jr. & Schaefer, 2010; Santos et al.,
215 2018). This class of soil presented high levels of fine and coarse sand (Table 2), where
216 these physical characteristics are associated to the formation processes of this
217 environment, especially by the presence of small streams that precede the first
218 undulating steps of the Cumarú Mountain, with undulations which aid the accumulation
219 of alluvial material derived from the highest part. Thus, this soil type is related to the
220 hydrological processes of sediment trapping, as also observed by Mendonça et al (2013)
221 in the region of the National Park of Viruá (Roraima), in similar environments to those
222 found in Serra da Mocidade.

223 The pH values determined for this soil class in the study area are in line with the
224 standards reported by EMBRAPA (2009). An analysis along the soil profile in this
225 forest type allowed to understand that there is a slight increase in the pH values from the
226 most superficial layer (0-20cm) to the deepest ones (20-40cm, 40-60cm) and,
227 consequently, a reduction in exchangeable and potential acidity (Table 1). This is a
228 process fed by the infiltration of exchangeable bases or increase of organic matter, as
229 previously established by Santos et al. (2011). Likewise, the values of available
230 phosphorus presented a decrease between the second (20-40 cm) and the third layer (40-
231 60 cm), being a strong indicative of the reduction of this element along the vertical soil
232 profile, as observed by Duivenvoorden. (1996), and it can act as a limiting element in
233 the larger / smaller vertical and horizontal structuring of the forest.

234 Organic matter also declines from the superficial to the deeper layers of this soil
235 class in the study area. This same observation was reported in the Viruá National Park,
236 with the authors suggesting that the topography of the terrain, especially those with soft
237 ripples, may retain organic matter in the superficial layers of the soil due to the sediment
238 trapping or the temporal outcropping of the water table (Mendonça et al., 2013), which
239 are the same environmental characteristics observed in the ecotone (LO) of the Cat's
240 Trail. In the same sense of the organic matter, the CEC also has a reduction pattern from

241 the most superficial to the deepest layers, presenting a CEC saturation problem with the
242 exchangeable Al, being common in this soil type, which can be aggravated by depth due
243 to the decrease of organic matter in the soil and exchangeable bases such as Ca^{2+} , Mg^{2+}
244 and K^+ , which would limit the development of the roots of the plants and affect the
245 structure of the forest (Sacramento et al., 2008).

246 The Yellow Ultisols is characterized as a soil type flooding free, with strong
247 presence of Open Ombrophylous Forest (Ab+As), with high contents of coarse sand and
248 clay (Table 2). According to Vale Jr. and Schaefer (2010), the formation factors of this
249 soil class are similar to that of the Oxisols, with the same geomorphological
250 characteristics and natural vegetation, but with a textural gradient. The pH values found
251 in this environment are slightly acidic and there is not much difference between the
252 layers (Table 1). This process may be happening due to the absence of both water table
253 outcrops and temporal flooding in this area (Quesada et al., 2009; Scopel et al., 2005).
254 The content of organic matter found in this soil type was high in the first layer,
255 decreasing drastically towards the deeper layers. This result indicates a direct
256 relationship with the CEC values found in this soil type, within the relational
257 congruence suggested by Ostertag (2001), where organic matter and clays are the
258 edaphic parameters with the greatest contribution to the formation of CEC values.

259 In this same analytical path, the concentrations of the micronutrients observed in
260 the profile of this soil highlight the higher concentration of Fe in the first two layers (0-
261 20 cm and 20-40 cm), being this element the main constituent of the structure of the
262 clays (Tsui et al., 2004). This chemical characteristic was also observed in the
263 ombrophilous forests of the Viruá National Park by Mendonça et al. (2017), suggesting
264 that higher Fe contents in the ombrophilous forests of regions under the influence of
265 treed and forested shade-loving mosaics may be due to the high presence of mineral
266 particles (oxides of iron) derived from the organic matter deposited on the soil in litter
267 form. This indicates that both Fe and the other micronutrients (Zn, Mn, B, Cu) have an
268 important role in the nutrient cycling of this forest environment, but without a clearly
269 defined role as a characterizer of forest types. This evidence the large range of
270 uncertainties that still persist in the evaluations on the relationships between hydro-
271 edaphic conditions and their role in the construction of Amazonian ecosystems. In this
272 way, it is inferred that a better spatial distribution of the pedo-phytosociological studies,
273 addressing peripheral regional gaps, can help us to generate environmental standards
274 (topographic gradients, periodicity of flooding, soil classes) that more accurately

275 configure the particular structural and floristic characteristics of different forest types,
 276 allowing better accuracy in biomass/carbon estimates, and giving the real importance of
 277 the Amazonia in the context of the global warming mitigation.

Depth	pH	OM	P	Cu	Fe	Zn	Mn	B	K	Ca	Mg	H+Al	Al	SB	CEC	Sat.	Sat.
(cm)	H ₂ O	g/kg				mg/kg						cmolc/Kg				Bases	Al
																V%	m%
Entisols Fluvents – Treed and Forested shade-loving (La+Ld)																	
0-20	5.1	14	3	6.80	92.00	8.20	3.00	0.43	0.10	0.20	0.10	3.40	0.20	0.40	3.80	11	33
20-40	5.2	7	2	6.70	37.20	8.05	1.40	0.25	0.06	0.10	0.10	2.80	0.20	0.26	3.06	9	43
40-60	4.8	5	2	4.60	14.20	5.80	1.10	0.24	0.06	0.10	0.10	2.50	0.20	0.26	2.76	9	43
Entisols Quartzipsamments – Area of Ecological Tension (LO)																	
0-20	4.6	23	3	1.60	9.80	5.65	15.40	0.26	0.20	0.30	0.10	6.40	0.40	0.60	7.00	9	40
20-40	4.9	9	3	2.10	18.60	3.80	2.40	0.38	0.15	0.30	0.10	4.20	0.30	0.55	4.75	12	35
40-60	4.9	5	2	2.00	19.00	3.60	1.20	0.35	0.13	0.30	0.10	3.80	0.20	0.53	4.33	12	27
Yellow Ultisols – Open Ombrophylous Forest (Ab+As)																	
0-20	4.6	11	3	2.90	106.00	5.95	6.50	0.34	0.15	0.60	0.20	4.70	0.20	0.95	5.65	17	17
20-40	4.7	5	2	2.10	102.00	4.30	1.90	0.32	0.12	0.20	0.10	3.40	0.30	0.42	3.82	11	42
40-60	4.6	5	1	0.70	80.00	3.45	2.00	0.27	0.06	0.10	0.10	3.40	0.20	0.26	3.66	7	43

278 **Table 1** - Chemical attributes determined in three profiles along a topographic gradient
 279 located in the Serra da Mocidade National Park, northern Brazilian Amazonia.
 280 Where: OM = organic matter, P = Phosphorus; Cu = Copper, Fe = Iron, Zn = Zinc, Mn = Manganese, B =
 281 Boron, K = Potassium, Ca = Calcium, Mg = Magnesium, H + Al = Acidable exchangeable, Al =
 282 Aluminium exchangeable, SB = Sum of bases, CEC =Cationic Exchange Capacity, V = Saturation by
 283 Bases, m =Saturation by Aluminium.
 284

285 **Table 2** - Physical attributes determined in three profiles along a topographic gradient
 286 located in the Serra da Mocidade National Park, northern Brazilian Amazonia.

Depth	Arg	Sil	ArT	ArG	ArF
(cm)		%		
Entisols Fluvents – Treed and Forested shade-loving (La+Ld)					
0-20	13.9	19.1	67.0	7.0	60.0
20-40	14.2	16.8	69.0	4.0	65.0
40-60	14.2	16.8	69.0	4.0	65.0
Entisols Quartzipsamments – Area of Ecological Tension (LO)					
0-20	15.7	11.3	73.0	20.0	53.0
20-40	17.4	9.6	73.0	21.0	52.0
40-60	16.9	8.1	75.0	23.0	52.0
Yellow Ultisols – Open Ombrophylous Forest (Ab+As)					
0-20	15.4	3.6	81.0	51.0	30.0
20-40	20.3	6.7	73.0	42.0	31.0
40-60	25.0	6.0	69.0	37.0	32.0

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Where: Arg Clay (<0.002mm), Sil Silt (0.053-0.002mm), Total ArT_Area, ArG_ Thick Sand (2.00-0.210mm), ArF_ Fine Sand (0.210-0.053mm).

303 **Table 3** - Soil classes, flood periodicity and structural parameters (DBH and total
 304 height) of the forest types observed in the sampling transect formed by the Cat's Trail,
 305 Serra da Mocidade National Park. Where: La + Ld = Mosaic of treed and forested
 306 shade-loving; LO = Area of ecological tension between forested shade-loving and
 307 ombrophylous forest; Ab + As = open ombrophylous forest associated with the first
 308 steps of the Cumaru Mountain. Different uppercase (Trees, ANOVA followed by Tukey
 309 test) and lowercase (Palms; Test t) letters in the columns indicate discrepancies ($\alpha =$
 310 0.05) between the values of the taxonomic groups.

Forest Type	Density (ind ha ⁻¹)		DBH (cm)		Ht (m)	
	Trees	Palms	Trees	Palms	Trees	Palms
La+Ld	940	0	13.7± 3.3 A	-	12.2±2.3 A	-
LO	710	45	17.3± 7.4 B	14.5±3.5 a	17.7±2.7 B	17.5±1.9 a
Ab+As	423	85	20.9±12.9 C	18.5±6.5 b	18.3±4.0 B	16.6±3.2 a

311 (*) DBH = diameter at breast height (cm) and Ht = total height (m)

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322 CONCLUSION

323 We conclude that each soil class determined in this study has a strong
324 association with the topographic gradient sampled in the study area situated on Serra da
325 Mocidade National Park, where lower altitude environments with larger flooding
326 periods are related to forest types of lower structural pattern (e.g. treed and forested
327 shade-loving) preferentially on oligotrophic soils (poor and sandy). These
328 characteristics indicate the formation of environments influenced by continuous hydro-
329 edaphic and geological processes, where seasonal flooding and sediment trawling are
330 part of the process of formation of the main forest types in the study area. Therefore,
331 edaphic factors and flooding periodicity are environmental characteristics that act as
332 environmental filters which are important in the formation of the landscape in this region
333 of rio Branco-rio Negro basin. These results improve our understanding of the
334 environmental factors that determine different forest types in this region of the
335 Amazonia, where environments with higher hydro-edaphic restrictions are a strong
336 indicator of forest types with lower vertical and horizontal structure.

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339 REFERENCES

340

341 BARBOSA, R.I. Distribuição das chuvas em Roraima. p. 325-335, 1997. In Barbosa,
342 R.I. Ferreira, E.J.G.; Castellón, E.G (Eds.). **Homem, Ambiente e Ecologia no Estado**
343 **de Roraima**. INPA-AM, Manaus (AM).

344

345 Barbosa, R.I.; Ferreira, C.A.C. Biomassa acima do solo de um ecossistema de
346 “campina” em Roraima, norte da Amazônia Brasileira. **Acta Amazonica**, Boa Vista, v.
347 34, n.4, p. 577- 586, 2004.

348

349 Barni, P. E. et al. Spatial distribution of forest biomass in Brazil’s state of Roraima,
350 northern Amazonia. **Forest Ecology and Management**, v. 377, p. 170-181, [s.m],
351 2016.

352

353 Brazil-MME. 1975. *Projeto RADAMBRASIL - Levantamento de Recursos Naturais*.
354 Ministério das Minas e Energia, Departamento Nacional de Produção Mineral, Rio de
355 Janeiro. 475 p. (<http://biblioteca.ibge.gov.br/visualizacao/livros/liv24025.pdf>.)

356

357 Carvalho, S,T & Domingues, P, E. Economic and deforestation scenario for the
358 Brazilian Amazon between 2006 and 2030, *Nova Economia*, v.26, n2, p. 585-621, 2016

359

360 Castilho, C.V.; Magnusson, W.E.; Araújo, R.N.O.; Luizão, R.C.C.; Luizão, F.J.; Lima,
361 A.P.; Higuchi, N. 2006. Variation in aboveground tree live biomass in a central
362 Amazonian Forest: Effects of soil and topography. *Forest Ecology and Management*,
363 234(1-3): 85-96.
364
365 Cordeiro, C. et al. Impact of sedimentary processes on white-sand vegetation in an
366 Amazonian megafan. **Journal of Tropical Ecology**, [s.l], v.32, n. 6, p. 498-509, [s.m],
367 2016.
368 Costa F, R. Magnusson, W, E. Luizao, R, C. Mesoscale distribution patterns of
369 Amazonian understorey herbs in relation to topography, soil and watersheds, *Journal of*
370 *Ecology*, 2005, v.93, p. 863-878.
371
372 Damasco, G.; Vicentini, A.; Castilho, C.V.; Pimentel, T.P.; Nascimento, H.E.M. 2013.
373 Disentangling the role of edaphic variability, flooding regime and topography of
374 Amazonian white-sand vegetation. *Journal of Vegetation Science*, 24(2): 384–394
375
376 Duivenvoorden, J. F. Patterns of tree species richness in rain forests of the middle
377 Caqueta area, Colombia, NW Amazonia. **Biotropica**, [s.l], v. 28, [s.n], p. 142-158,
378 [s.m], 1996
379 EMBRAPA. Manual de Análises Químicas de Solos, Plantas e Fertilizantes. 2009.
380 624p.
381 Fearnside, P. M. Brazil's Amazonian forest carbon: the key to Southern Amazonia's
382 significance for global climate. *Regional Environmental Change*, 2018; 18: 47-61.
383 Fearnside, P.M. Quantificação do serviço ambiental do carbono nas florestas
384 amazônicas brasileiras. **Oecologia Brasiliensis**, Manaus, v. 12, n. 4, p. 743-756, [s.m],
385 2008.
386
387 Fearnside, P.M.; Ferraz, J. 1995. A conservation gap analysis of Brazil's Amazonian
388 vegetation. *Conservation Biology*, 9(5): 1134-1147.
389 Ferraz J. et al. Distribuição dos solos ao longo de dois transectos em floresta primária ao
390 Norte de Manaus (AM). In: Higuchi N. et al. Pesquisas florestais para a conservação da
391 floresta e reabilitação de áreas degradadas da Amazônia. Instituto Nacional de Pesquisas
392 da Amazônia. 1998; 111-143.
393 IBGE. Manual técnico da vegetação brasileira: Sistema fitogeográfico, inventário das
394 formações florestais e campestres, técnicas e manejo de coleções botânicas,
395 procedimentos para mapeamentos. Instituto Brasileiro de Geografia e Estatística. 2012.
396
397 IPCC (INTERGOVERNMENTAL PANEL ON CLIMATE CHANGE). **Guidelines for**
398 **National Greenhouse Gas Inventories**. National greenhouse gas inventories
399 programme, H. S. Eggleston, L. Miwa, T. Ngara and K. Tanabe (eds).
400 Intergovernmental Panel on Climate Change (IPCC), Institute for Global Environmental
401 Strategies (IGES), JAPAN, 2006.
402

403 Juliano Franco-Moraesa, Armindo F.M.B. Baniwa, Flávia R.C. Costa, Helena P. Lima,
404 Charles R. Clement, Glenn H. Shepard Jr. 2019. Historical landscape domestication in
405 ancestral forests with nutrient-poor soils in northwestern Amazonia. *FORECO*.
406

407 Junk, J, W, Wittmann, F, Schongart, J, Piedade, M. A classification of the major
408 habitats of Amazonian black-water river floodplains and a comparison with their white-
409 water counterparts, *Wetlands Ecol Manage*, v. 23, p. 677-693, 2015.
410

411 Khorramdel S. et al. Evaluation of carbon sequestration potential in corn fiels with
412 different management systems. *Soil & Tillage Research*. 2013; 133: 25-31.

413 Laurance, W.F.; Fearnside, P.M.; Laurance, S.G.; Delamonica, P.; Lovejoy, T.E.;
414 Rankin-de-Merona, J.M.; Chambers, J.Q.; Gascon, C. 1999. Relationship between soils
415 and Amazon forest biomass: a landscape-scale study. *Forest Ecology and Management*,
416 18: 127-138.
417

418 Lewis, S.L.; Phillips, O.L.; Baker, T.R.; Lloyd, J.; Malhi, Y.; Almeida, S., *et al.* 2004.
419 Concerted changes in tropical forest structure and dynamics: evidence from 50 South
420 American long-term plots. *Philosophical Transactions of the Royal Society of London*,
421 *Series B: Biological Sciences*, 359(1443): 421-436.

422 Luizão, F. J. et al. Soil acidity and nutrient deficiency in central Amazonian heath forest
423 soils. *Plant Ecol*, [s.l], v.192, [s.n], p. 209-224, [s.m], 2007.
424

425 Magalhães RC, Gomes RCM. Mineralogy and Chemistry of the lowland soil and its
426 sensibilities in the process of Lands Falls in community Divino Espírito Santo
427 (Amazonas, Brazil). *Soc. & Nat.* 2013; 25: 609-621.

428 Marcelo T. Nascimento, Lidiany C. da Silva Carvalho, Reinaldo I. Barbosa & Dora M.
429 Villela (2014). Variation in floristic composition, demography and above-ground
430 biomass over a 20-year period in an Amazonian monodominant forest, *Plant Ecology &*
431 *Diversity*, 7:1-2, 293-303.
432

433 Marcelo Trindade Nascimento, Reinaldo Imbrozio Barbosa, Kyle G. Dexter, Carolina
434 Volkmer de Castilho, Lidiany Camila da Silva Carvalho, Dora Maria Villela. 2017. Is
435 the *Peltogyne gracilipes* monodominant forest characterised by distinct soils? *Acta*
436 *Oecologica*, v, 85, 104-107
437

438 Mendonça BA. et al. Solos e Geoambientes do Parque Nacional do Viruá e entorno,
439 Roraima: visão integrada da paisagem e serviço ambiental. *Ciência Florestal*. 2013; 23:
440 427-442.

441 Mendonça, B.A.F.; Fernandes Filho, E.I.; Schaefer, C.E.G.R.; Carvalho, A.F.; Vale Jr,
442 J.F.; Corrêa, G.R. 2014. Use of geophysical methods for the study of sandy soils under
443 Campinarana at the National Park of Viruá, Roraima state, Brazilian Amazonia. *Journal*
444 *of Soils and Sediments*, 14(3): 525-537.

445 Mendonça, B.A.F.; Fernandes, E.I.F.; Schaefer, C.; Mendonca, J.G.F.; Vasconcelos,
446 B.N.F. 2017. Soil-vegetation relationships and community structure in a "terra-firme"-
447 white-sand vegetation gradient in Virua National Park, northern Amazon, Brazil. *Anais*
448 *da Academia Brasileira de Ciências*.n, 89, v, 2, p. 1269-1293

449 Nogueira, E.M.; Yanai, A.M.; Fonseca, F.O.; Fearnside, P.M. Carbon stock loss from
450 deforestation through 2013 in Brazilian Amazonia. *Global Change Biology*, 2015; 21:
451 1271–1292.

452 Ostertag R. effects of nitrogen and phosphorus availability on fine root dynamics in
453 Hawaiian montane forests. *Ecology*. 2001; 82: 485-499.

454 Phillips, O.P; Vargas, O.; Monteagudo, A.L, Cruz, A.P, Zans, M., Sánchez, W, Yli-
455 halla, M and Rose, R. Habitat association among Amazonian tree species: a landscape-
456 scale approach, *Journal of Ecology*, v. 91, p. 757-775, 2003
457

458 Phillips, O. L. et al. Changes in the carbon balance of tropical forest: evidence from
459 long-term plots. **Science**, [s.l], v. 282, n. 5388, p. 439-442, [s.m], 1998.
460

461 Prance, G.T.; Schubart, H.O.R. 1977. Notes on the vegetation of Amazonia I. A
462 preliminary note on the origin of the open white sand Campinas of the Lower Rio
463 Negro. *Brittonia*, 30(1): 60-63.

464 Quesada CA. et al. Soils of Amazonia with particular reference to the Rain for sites.
465 *Biogeosciences Discussion*. 2009; 6: 3851-3921.

466 Sacramento et al. Atributos químicos e físicos de um argissolo cultivado com *Panicum*
467 *maximum* Jacq. Cv. IPR-86 Milênio, sob lotação rotacionada e adubado com
468 nitrogênio. *R. Bras. Ci. Solo*. 2008; 32; 1: 183-193.

469 Santos HG. et al. Sistema Brasileiro de Classificação de Solos. Embrapa Solos. Brazil.
470 2018. 590p.

471 Santos HG. et al. O novo mapa de solos do Brasil: legenda atualizada. Rio de Janeiro:
472 Embrapa Solos. 2011; 67p.

473 Santos, C.P.F. et al. 2007. Mapeamento dos remanescentes e ocupação antrópica no
474 Bioma Amazônia.
475 <<http://marte.dpi.inpe.br/rep/dpi.inpe.br/sbsr@80/2006/11.18.01.25?mirror=dpi.inpe.br>
476 >. Acesso em 02 dez. 2017.
477

478 Santos, N.M.C.; Vale Júnior, J.F.; Barbosa, R.I. 2013. Florística e estrutura arbórea de
479 ilhas de mata em áreas de savana do norte da Amazônia brasileira. *Boletim do Museu*
480 *Paraense Emílio Goeldi (Ciências Naturais)*, 8(2): 205-221.

481 Scaranello, M. et al. Height-diameter relationships of tropical Atlantic moist forest trees
482 in southeastern Brazil. **Scientia Agricola**, São Paulo, v. 69, p. 26-37, jan/feb, 2012.
483

484 Scopel I. et al. Formação de areais e perspectivas de uso e manejo de Neossolos
485 Quartzarênicos em Serranópolis (GO). *Boletim Goiano de Geografia*. 2005: 25: 11-27.

486 Silva, L.F.S.G.; Castilho, C.V.; Cavalcante, C.O.; Pimentel, T.P.; Fearnside, P.M.;
487 Barbosa, R.I. 2016. Production and stock of coarse woody debris across a hydro-
488 edaphic gradient of oligotrophic forests in the northern Brazilian Amazon. *Forest*
489 *Ecology and Management*, 364: 1-9.

- 490 Suwa R. et al. Meaning of the topographic gradient in stem diameter-Height allometry
491 for precise biomass Estimation of a tropical humid forest in the central Amazon. JARQ.
492 2013; 47: 109-114.
- 493 ter Steege, H.; Vaessen, R.W.; Cárdenas-López, D.; Sabatier, D.; Antonelli, A.;
494 Oliveira, S.M.; Pitman, N.C.A.; Jørgensen, P.M.; Salomão, R.P. 2016. The discovery of
495 the Amazonian tree flora with an updated checklist of all known tree taxa. *Scientific*
496 *reports*, 6: 29549
- 497 Toledo J de, Magnusson W (2012) Tree mode of death in Central Amazonia: Effects of
498 soil and topography on tree mortality associated with storm disturbances. *Forest*
499 *Ecology and Management*, 263: 253–261.
- 500 Tsui CC, Chen ZS, Hsieh CF. Relationships between soil properties and slope position
501 in a lowland rain forest of Southern Taiwan. *Geoderma*. 2004; 123: 131-142.
- 502 Tuomisto, H, Zuquim, G, Cardenas, G. Species richness and diversity along edaphic
503 and climatic gradients in Amazonia, *Ecography*, v. 37, n 11, p. 1034-1046, 2014
- 504 Tuomisto, H., Ruokolainen, K., Aguilar, M. & Sarmiento, A. (2003) Floristic patterns
505 along a 43-km long transect in an Amazonian rain forest. *Journal of Ecology*, 91, 743–
506 756
- 507 Vale Jr. JF, Schaefer CEGR. Solos sob Savanas de Roraima. Gênese, classificação e
508 relações ambientes. Gráfica Ioris: Boa Vista-Roraima. Brazil. 2010, 219p.
- 509 Veloso, H.P., Rangel Filho, A.L.R. & Lima, J.C.A. 1991. Classificação da vegetação
510 brasileira adaptada a um sistema universal. Instituto Brasileiro de Geografia e
511 Estatística, Rio de Janeiro.
512
- 513 USDA (1999). Soil Taxonomy: A Basic System of Soil Classification for Making and
514 Interpreting Soil Surveys
515