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METEOROLOGICAL MODELS FOR DETERMINATION OF SURFACE RADIO REFRACTIVITY OVER NIGERIA Ajileye, O. O¹. Popoola O. S². Kayode, F. F¹ and Rabiu, A. B¹

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11 Abstract

Previous studies showed that linear meteorological expressions obtained were localized and 12 could not be generally applied. It is therefore required that more locations should be investigated to 13 deduce new linear meteorological models best suitable for estimation of surface refractivity. Surface 14 15 meteorological data, including pressure, temperature and relative humidity, was downloaded from 16 Modern-Era Retrospective analysis for Research and Application (MERRA - 2) for six locations defined by 17 different climatic conditions over Nigeria, namely Yenagoa, Abakaliki, Ibadan, Anyigba, Jalingo and 18 Sokoto for a period of 40 years partitioned into two periods of 20 years each. The 1979 – 1998 dataset 19 of atmospheric temperature, T (K), atmospheric pressure, P (hPa) and relative humidity, Rh (%) were 20 extracted using text import wizard to calculate surface refractivity using existing model (N_{itu-r}) . Regression analysis was carried out to obtain new linear meteorological expressions as function of 21 22 temperature (N_t), relative humidity (N_{Rh}), and combination of relative humidity and temperature 23 (N_{Rht}). The new expressions were tested using 1999 – 2018 meteorological dataset and the results of 24 surface refractivity from the new linear expressions were compared with values from existing equations.

25 It was observed that high relative humidity and low temperature values prevalent in the three 26 southern stations – Yenagoa, Abakaliki and Ibadan – significantly enhanced the suitability of N_{Rh.t} linear meteorological model for estimation of surface refractivity values comparable to N_{itu-r}. Contrary to the 27 28 observations in the southern stations, in the northern stations, relative humidity contributions (N_{Bh}) 29 had the best linear correlation of 0.96 at Anyigba, 0.98 at Jalingo and Sokoto. Estimation of surface 30 refractivity from the new linear meteorological models was found to be best as a function of 31 temperature and relative humidity in the south and a function of relative humidity in the north. Linear 32 meteorological models as a function of relative humidity and temperature were best suitable for 33 stations in the south with minimum correlation of 0.98 while linear expressions as a function of relative 34 humidity only were best suitable for the stations in the north with minimum correlation of 0.96.

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1.0 Introduction

38 Atmospheric refractivity has been well studied over the years and some mathematical 39 expressions have been developed to estimate the value of refractivity [1]. However, most of the expressions are complex and involve some nonlinear expressions. The complexity makes it more difficult 40 41 to integrate the refractivity expressions into other formula to develop simple closed-form mathematical 42 expressions needed in wireless link design [2]. In recent years, researchers have tried to develop simple 43 mathematical equations that can be used to estimate the radio refractivity from the atmospheric 44 parameters [3]. Prediction of radio refractivity has been required since the early days of microwave 45 radio links. Although one of the earliest prediction models is still in use, it is only valid for continental 46 temperate climate [5].

In a study conducted by [6], radiosonde data from 65 stations covering parts of Australia, India,
 South Africa, Europe and North America, most of them with around 16 years of data, typically 1997–

49 2012, were processed. A new prediction model for the distribution of refractivity gradient in the 50 atmospheric surface layer, having better prediction accuracy than existing models, and using only data 51 that can be obtained from surface weather stations, was obtained with rms error of 17 N-units per km, 52 and correlation coefficient of 0.79

53 In the recent research conducted by [4], radiosonde meteorological data for Cross River state, 54 Nigeria was obtained from Nigerian Meteorological Agency (NIMET) for the twelve months in 2013. The 55 study examined the correlation among radio refractivity and the meteorological parameters, namely 56 atmospheric temperature (T), atmospheric pressure (P) and relative humidity (H). The correlation values 57 from the results showed that there is a positive correlation between surface refractivity (N) and 58 temperature (T) and pressure (P). However, there is a negative correlation between N and relative 59 humidity (Rh) which indicated inverse proportionality between N and Rh. In all, T and products of TPRh 60 showed highest correlation values for the model training datasets; as such the two parameters were 61 used in the development of the linear regression model that can be used to estimate the radio 62 refractivity for the available meteorological dataset.

The model was developed using XURU online regression tool where the values of T was X1, TPRh was X2 and N was Y. The dataset of T, TPRh and N for the 12 months were pasted into the text box provided in Xuru Multiple Linear Regression (MLR) webpage and the model obtained from the Xuru MLR was given as;

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$$N = 2.38T + 0.000051[(T)(P)(Rh)] + 198.38$$
 1.0

Two sample meteorological dataset from published articles were also used by [4] to validate the model. The model gave a maximum absolute percentage error of 2.46 % for the first test meteorological dataset while it gave a maximum absolute percentage error of 1.25 % for the second test meteorological dataset. The results that was obtained from the new model showed that the model can estimate refractivity with a maximum prediction error of about ±3.35%.

In this study, a set of meteorological data measured over 20 years was used to deduce useful linear expressions for the estimation of surface refractivity adaptable to specific climatic features in six locations across Nigeria. The linear expressions were tested using meteorological data from another period of 20 years. The values of surface refractivity from the new linear expressions were compared with values from existing equations.

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80 2.0 Theoretical Background

The refractive index of the troposphere is an important factor in predicting the performance of 81 82 terrestrial radio links. Refractive index variations of the atmosphere affect radio frequencies above 30 83 MHz, although these effects become significant only at frequencies greater than about 100MHz 84 especially in the lower atmosphere. The radio refractive index n of the troposphere deviates slightly 85 from unity due to the polarisability of the constituent molecules by the incident electromagnetic field, 86 and the quantum mechanical resonances at certain frequency bands [7]. While molecular polarisability 87 is independent of frequency up to millimeter waves, molecular resonance is totally frequency 88 dependent, and n tends to be dispersive above approximately 50 GHz.

Radio refractivity N is a measure of deviation of refractive index n of air from unity which is
scaled-up in parts per million to obtain more amenable figures. Thus, N is a dimensionless quantity
defined as measured in N units [8].

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$$N = (n-1) \times 10^6$$

N depends on meteorological parameters of pressure P (hPa), temperature T (K) and water vapour
 pressure e (hPa), as given by the relation [5]:

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$$N_{itu-r} = \frac{7}{2}$$

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$$N_{itu-r} = \frac{77.6P}{T} + 3 \times 10^5 \frac{e}{T^2}$$

100 The vapour pressure is also related to the relative humidity Rh (%):

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$$e = \frac{Rhe_s}{100}$$

 e_s is the maximum (or saturated) vapour pressure at the given air temperature t ${}^{0}C$, and may be 104 105 obtained from:

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 $e_s = 6.11 \exp\left[\frac{17.502t}{(t+240.97)}\right]$ 5.0

3.0

4.0

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109 Generally P and e decrease rapidly with height while T decreases slowly with height [5]. Horizontal 110 variation of refractive index is generally negligible in the lower troposphere compared to the large-scale 111 vertical variation which has a median gradient of about -40 N/km near the surface in mid-latitude and 112 most temperate regions. However, significant deviations can arise from local or mesoscale 113 meteorological factors, especially in the tropics. This horizontal variation of refractive index is very 114 significant over Nigeria because of the significant change in climatic condition from the coastal region in 115 the extreme south to the semi-arid region in the extreme North.

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117 3.0 **Climate of the Study Area**

118 The study area in this research spread across six geopolitical zone across Nigeria. Each locations 119 has a diverse climatic pattern to ensure adequate representation of surface refractivity patterns. The 120 locations are shown in figure 1 namely Yenagoa in south-south zone, Abakaliki in south-east zone, 121 Ibadan in south-west zone, Anyigba in north-central zone, Jalingo in north-east zone and Sokoto in 122 north-west zone. The climate at Yenagoa is tropical. Most months of the year are marked by significant 123 rainfall. The short dry season has little impact. The average annual temperature is 26.7 °C in Yenagoa. In 124 a year, the average rainfall is 2899 mm [9]. The driest month is January, with 40 mm of rain. The 125 greatest amount of precipitation occurs in September, with an average of 472 mm. March is the 126 warmest month of the year. The temperature in March averages 28.0 °C [15]. The lowest average 127 temperatures in the year occur in July, when it is around 25.4 °C. There is a difference of 432 mm of 128 precipitation between the driest and wettest months [10]. The variation in temperatures throughout the 129 year is 2.6 °C.

130 Abakaliki climate is classified as tropical. The average temperature in Abakaliki is 27.7 °C. 131 Precipitation here averages 1918 mm [11]. Precipitation is the lowest in December, with an average of 7 132 mm. Most precipitation falls in September, with an average of 291 mm. At an average temperature of 133 29.6 °C, March is the hottest month of the year [12]. In August, the average temperature is 26.0 °C [15]. 134 It is the lowest average temperature of the whole year. Between the driest and wettest months, the 135 difference in precipitation is 284 mm. The average temperatures vary during the year by 3.6 °C.

136 The climate is tropical in Ibadan. The average annual temperature is 26.5 °C in Ibadan. About 137 1311 mm of precipitation falls annually. The driest month is January. There is 6 mm of precipitation in 138 January. In June, the precipitation reaches its peak, with an average of 190 mm [16]. With an average of 139 28.6 °C, March is the warmest month. At 24.1 °C on average, August is the coldest month of the year. 140 The precipitation varies 184 mm between the driest month and the wettest month. The variation in 141 annual temperature is around 4.5 °C.

Anyigba has a tropical climate. The average temperature in Anyigba is 25.4 °C. The average annual rainfall is 1321 mm. The least amount of rainfall occurs in January. The average in this month is 7 mm. The greatest amount of precipitation occurs in September, with an average of 265 mm. The temperatures are highest on average in March, at around 27.9 °C [17]. The lowest average temperatures in the year occur in August, when it is around 23.8 °C. The variation in the precipitation between the driest and wettest months is 258 mm. The variation in temperatures throughout the year is 4.1 °C.

The climate is tropical in Jalingo. The temperature here averages 27.3 °C. About 1053 mm of precipitation falls annually. The driest month is January. There is 0 mm of precipitation in January. In August, the precipitation reaches its peak, with an average of 224 mm. With an average of 30.6 °C, April is the warmest month [14]. At 25.2 °C on average, December is the coldest month of the year. The precipitation varies 224 mm between the driest month and the wettest month. The variation in annual temperature is around 5.4 °C.

The climate in Sokoto is referred to as a local steppe climate. There is little rainfall throughout the year. In Sokoto, the average annual temperature is 28.4 °C [18]. The average annual rainfall is 629 mm. The least amount of rainfall occurs in January. The average in this month is 0 mm. The greatest amount of precipitation occurs in August, with an average of 211 mm. The temperatures are highest on average in April, at around 33.2 °C. The lowest average temperatures in the year occur in January, when it is around 24.5 °C. The variation in the precipitation between the driest and wettest months is 211 mm. The variation in temperatures throughout the year is 8.7 °C.

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163 4.0 Data Source and Analysis

Surface meteorological data, including pressure, temperature and relative humidity, was downloaded from Modern-Era Retrospective analysis for Research and Application (MERRA - 2). The data project was stimulated by the recognition that various aspects of the hydrologic cycle represented in previous generations of re-analyses were not adequate for climate and weather studies. MERRA – 2 offered an improvement upon the water cycle as a contribution to the science community and reanalysis research [19].

MERRA uses a three-dimensional variation (3D-Var) analysis algorithm based on the Grid-point Statistical Interpolation scheme. Like other current re-analyses, it makes extensive use of satellite radiance information, including data from hyper-spectral instruments such as the Atmospheric Infrared Sounder (AIRS) on Aqua [20]. MERRA was processed in three separate streams, each spun-up in two stages: Stream 1 for 1 January 1979 to 31 December 1992, followed by Stream 2 for 1 January 1993 to 31 December 2000, and then continues with Stream 3 for 1 January 2001 to the present.

176 The MERRA meteorological data for six locations defined by different climatic conditions over 177 Nigeria, namely Yenagoa, Abakaliki, Ibadan, Anyigba, Jalingo and Sokoto, were downloaded for two 178 periods of 20 years, 1979 – 1998 and 1999 – 2018, on monthly averages. The dataset of atmospheric 179 temperature, T (K), atmospheric pressure, P (hPa) and relative humidity, Rh (%) were extracted using 180 text import wizard and appropriate delimited options to prepare excel spreadsheet format for data manipulations. Radio refractivity (N) was computed from a set of T, Rh and P on monthly basis within 181 182 the period of 1979 – 1998 using ITU-R equation 3.0 with a view to deduce monthly averages for 20 183 years. Monthly values of t ($^{\circ}$ C), Rh and P were also separately estimated for the period covering 1979 – 1998 as shown in tables 1 - 3. 184

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Figure 1: Map showing the Six Locations of the Study Areas over Nigeria

Multiple linear regression was developed for each location to estimate radio refractivity directly from meteorological parameters. This was carried out using scattered plots of N and t (0 C); N and Rh; N and Rh, t (⁰C) with the aid of Origin graphing software to obtain a linear relationship between N and other meteorological parameters for different climatic zones across Nigeria. The linear expressions obtained for N from different regression analysis in the six climatic locations (as a function of t, Rh and t, Rh) and the ITU-R equation were applied on new set of monthly averages of meteorological data within a period of 1999 - 2018. Values of N obtained from linear regressions from each location were compared with values of N estimated from ITU-R equations. Correlation between radio refractivity values evaluated from linear expressions and ITU-R equation were computed to determine the degree of agreement and suitability of the new empirical relationships.

Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	299.46	297.01	298.45	299.28	299.00	297.92
2	300.06	300.32	299.12	300.44	299.87	298.30
3	300.19	300.16	299.82	302.05	303.53	302.55
4	300.13	300.05	299.62	300.75	303.06	306.18
5	299.85	299.83	299.28	299.88	300.32	306.32
6	299.03	298.79	298.58	299.12	299.86	303.22
7	298.13	298.42	297.63	298.10	298.18	300.68
8	297.84	298.34	297.38	297.93	297.75	299.38
9	298.21	298.73	297.91	298.15	298.46	300.22
10	299.04	298.90	298.63	298.94	299.49	299.95
11	299.38	299.87	298.58	298.78	299.53	298.77
12	299.67	298.65	298.16	298.47	299.01	297.25
Monthly Mean	299.25	299.09	298.59	299.32	299.84	300.89

221	Table 2: Atmospheric Pressure (hPa) at ground level averaged monthly over a period of 1979 – 1998
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Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	1008.29	1001.57	989.87	985.05	985.05	979.13
2	1007.86	999.28	989.53	984.69	984.69	979.52
3	1007.43	1000.41	988.88	983.80	983.8	976.51
4	1007.77	1000.42	989.25	984.28	984.28	975.40
5	1009.13	1001.31	990.66	985.68	985.68	976.39
6	1010.04	1003.09	991.51	986.48	986.48	977.88
7	1011.85	1003.61	993.03	988.04	988.04	979.54
8	1010.93	1003.60	992.10	987.08	987.08	978.78
9	1010.58	1001.75	991.83	986.91	986.91	978.88
10	1009.39	1001.49	990.83	985.81	985.81	978.54
11	1008.50	1000.43	990.10	985.19	985.19	978.68
12	1008.30	1001.80	990.07	985.28	985.28	979.60
Monthly Mean	1009.17	1001.56	990.64	985.69	976.01	978.24

227 Table 3: Relative humidity (%) at 2 m above ground averaged monthly over a period of 1979 – 1998)

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Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	82.31	54.61	76.42	58.85	38.77	18.94
2	80.13	75.93	76.77	45.73	21.27	11.03
3	86.59	80.79	82.67	64.66	26.1	13.15
4	86.85	81.80	85.16	73.63	55.25	20.31
5	87.69	82.89	87.26	81.23	76.07	42.54
6	89.43	86.18	87.35	82.80	76.1	56.97
7	89.20	86.25	88.52	84.86	83.91	70.96
8	89.67	85.61	88.49	85.79	85.97	79.72
9	89.67	85.76	87.97	85.81	83.24	74.31
10	88.04	85.96	87.02	82.62	70.67	46.15
11	85.66	79.24	83.09	74.25	49.28	24.15
12	81.55	64.03	76.31	62.47	35.25	24.78
Monthly Mean	86.40	79.09	83.92	73.56	58.49	40.25

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233 5.0 Results and Discussion

234 5.1 Derivation of Linear Meteorological Models for Surface Refractivity Estimation

235 The surface radio refractivity values estimated from existing ITU-R equation (quoted in equation 236 3) for six locations across Nigeria were presented in table 4. The results showed monthly averages of 237 surface refractivity a period of 1979 – 1998. The spatial trend showed a significant decrease in monthly 238 average of surface refractivity from southern location to the northern locations with the highest value of 239 383.41 N-Unit in Yenagoa and the least value of 314.81 N-Unit in Sokoto. Monthly distributions of 240 refractivity values showed that high values were attained from March to October for Yenagoa, Abakaliki 241 and Ibadan while high values were attained from May to October for Anyigba and Jalingo. Sokoto 242 attained its high values from June to September.

The distributions were greatly influenced by wide variations in meteorological parameters most especially temperature (T) and relative humidity (Rh) along the latitudes across Nigeria. Temperature values reduced on monthly average in the south from 299.25 K in Yenagoa to 298. 59 K at Ibadan while it increased in the north from 299.32 K at Anyigba to 300.89 K at Sokoto. Atmospheric pressure reduced south – north on monthly average from 1009.17 hPa at Yenagoa to 978.24 hPa at Sokoto. Also, relative humidity on monthly average reduced south – north from 86.40 % at Yenagoa to 40.25 % at Sokoto.

Linear regression analysis between the N values obtained from ITU-R in table 4 and meteorological parameters measured within the same period were carried out to obtain new expressions in terms of t (0 C) and Rh for the six locations identified over Nigeria. Linear expressions 6a – 6c were obtained for Yenagoa as possible empirical models for estimating surface refractivity expressed in terms of relative humidity and temperature only (Eq. 6a); in terms of temperature only (Eq. 6b) and in terms of relative humidity only (Eq. 6c).

The regression analysis was repeated for other locations namely Abakaliki, Ibadan, Anyigba, Jalingo and Sokoto. The resulting linear expressions are presented subsequently. Linear expressions for Abakaliki were presented in equations 7a – 7c; expressions for Ibadan were presented in equations 8a – 8c; expressions for Anyigba were presented in equations 9a – 9c; expressions for Jalingo were presented in expressions for Anyigba were presented in equations 9a – 9c; expressions for Jalingo were presented

260 Table 4: Surface Refractivity (N-Units) estimated from ITU-R Equation (N_{itu-r}) averaged over a period 261 of 1979- 1998

	1					
Months	Yenagoa	Abakaliki	Ibadan	Anyigba	Jalingo	Sokoto
1	378.55	360.38	360.62	338.47	307.18	279.92
2	378.44	351.45	364.15	322.90	283.39	269.60
3	388.58	379.52	375.96	358.10	294.80	272.44
4	388.71	380.79	378.56	366.16	344.21	287.97
5	388.66	383.41	380.02	373.36	365.51	333.31
6	386.68	380.93	376.51	371.80	363.46	348.79
7	381.86	377.73	373.42	369.75	366.35	360.53
8	380.69	377.50	371.84	369.86	366.65	366.80
9	382.59	379.08	373.85	370.97	366.61	363.17
10	384.63	379.98	376.16	370.44	353.74	320.61
11	382.94	370.70	370.35	357.95	323.21	287.38
12	378.57	355.28	359.21	340.65	302.34	287.15
Monthly Mean	383.41	373.06	371.72	359.20	336.45	314.81

The expressions showed linear dependence of surface refractivity on temperature and relative humidity. It was observed that the expressions had a pattern of highly varied numerical constants which makes the expressions significantly dependent on weather characteristics at specific location. The linear meteorological models are listed as follows:

274	$N_{Rh,t(Yenagoa)} = 335.32 + 0.33Rh + 0.75t$	6a
275	$N_{t(Yenagoa)} = 344.07 + 1.50t$	6b
276	$N_{Rh(Yenagoa)} = 326.57 + 0.66Rh$	6c
277		
278	$N_{Rh,t(Abakaliki)} = 251.0 + 0.69Rh + 1.78t$	7a
279	$N_{t(Abakaliki)} = 280.27 + 3.55t$	7b
280	$N_{Rh(Abakaliki)} = 261.73 + 1.37Rh$	7c
281		
282	$N_{Rh,t(Ibadan)} = 285.68 + 0.6Rh + 1.39t$	8a
283	$N_{t(Ibadan)} = 300.25 + 2.79t$	8b
284	$N_{Rh(Ibadan)} = 271.1 + 1.2Rh$	8c
285		
286	$N_{Rh,t(Anyigba)} = 358.18 + 0.66Rh + 1.67t$	9a
287	$N_{t(Anyigba)} = 446.8 - 3.33t$	9b
288	$N_{Rh(Anyigba)} = 269.56 + 1.22Rh$	9c
289		
290	$N_{Rh,t(Jalingo)} = 381.84 + 0.67Rh - 3.15t$	10a
291	$N_{t(Jalingo)} = 505.25 - 6.29t$	10b

292	$N_{Rh(Jalingo)} = 258.43 + 1.33Rh$	10c
293		
294	$N_{Rh,t(Sokoto)} = 381.84 + 0.67Rh - 3.15t$	11a
295	$N_{t(Sokoto)} = 505.25 - 6.29t$	11b
296	$N_{Rh(Sokoto)} = 258.43 + 1.33Rh$	11c

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5.2 Performance of Linear Meteorological Model in Estimation of Surface Refractivity

The results of monthly estimates of surface refractivity from new set of meteorological data for a period of 1999 - 2018 using equations 3, 6, 7, 8, 9, 10 and 11 are shown in tables 5 - 10. The values of surface refractivity averaged on monthly basis within the period of 20 years showed that surface refractivity had highest linear dependence on relative humidity in five out of the six locations. The five locations include Yenagoa, Abakaliki, Anyigba, Jalingo and Sokoto. Surface refractivity estimated from ITU-R equation was compared with values from linear meteorological models in equations 6 - 11 to determine the suitability of the expressions.

The mean values of surface refractivity from different linear meteorological models for the six 306 307 locations are presented in table 11. In Yenagoa, ITU-R had a monthly average value of 381.87 N-Units 308 while N_{Rh} had the closest value of 383.16 N-Units. In Abakaliki, ITU-R had a monthly average value of 370.75 N-Units while N_{Rh} had the closest value of 369.35 N-Units. In Ibadan, ITU-R had a monthly 309 310 average value of 370.02 N-Units while Nt had the closest value of 370.89 N-Units. In Anyigba, ITU-R had 311 a monthly value of 358.74 N-Units while N_{Rh} had the closest value of 359.58 N-Units. In Jalingo, ITU-R 312 had a monthly value of 338.77 N-Units while N_{Rh} had the closest value of 337.09 N-Units. In Sokoto, 313 ITU-R had a monthly value of 319.09 N-Units while N_{Rh} had the closest value of 319.61 N-Units. The 314 temporal and spatial difference in surface refractivity values estimated from linear meteorological 315 models with highest dependence is given as ±1.12 N-Units when compared with values from ITU-R.

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321 Table 5: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Yenagoa for 322 a period of 1999 – 2018

Months	N _{itu-r}	N _{Rh,t(Yenagoa)}	N _{t(Yenagoa)}	N _{Rh(Yenagoa)}
1	363.77	379.11	382.64	375.58
2	383.28	383.13	384.86	381.41
3	386.42	383.99	384.23	383.75
4	386.54	384.03	384.17	383.90
5	386.89	384.23	383.72	384.75
6	385.46	384.21	382.62	385.81
7	382.87	383.80	381.99	385.60
8	380.17	383.30	381.54	385.05
9	383.83	384.07	382.19	385.96
10	385.49	384.17	383.01	385.32
11	386.98	384.06	384.39	383.73
12	370.75	380.45	383.85	377.05
Monthly Mean	381.87	383.21	383.27	383.16

Table 6: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Abakaliki for <mark>a period of 1999 – 2018</mark>

Months	N _{itu-r}	N _{Rh,t(Abakaliki)}	N _{t(Abakaliki)}	N _{Rh(Abakaliki)}
1	330.03	331.42	365.51	335.82
2	371.35	352.02	377.26	365.02
3	378.02	355.09	376.69	371.68
4	378.92	355.59	376.30	373.07
5	379.55	355.95	375.52	374.56
6	379.05	356.37	371.82	379.07
7	377.31	355.76	370.51	379.16
8	376.03	355.18	370.23	378.29
9	377.81	355.97	371.61	378.49
10	378.93	356.41	372.22	378.77
11	374.23	353.50	375.66	369.56
12	347.73	340.84	371.33	348.72
Monthly Mean	370.75	352.01	372.89	369.35

Table 7: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Ibadan for a period of 1999 – 2018

Months	N _{itu-r}	N _{Rh,t(Ibadan)}	N _{t(Ibadan)}	N _{Rh(Ibadan)}
1	342.73	359.37	366.68	352.28
2	371.04	371.00	373.52	368.73
3	375.60	372.78	373.65	372.15
4	377.65	373.71	373.49	374.18
5	377.85	374.16	372.12	376.45
6	375.43	373.55	370.33	377.01
7	373.73	373.06	369.41	376.94
8	371.97	372.69	368.21	377.40
9	375.22	373.81	369.80	378.06
10	376.35	373.69	371.59	376.03
11	371.36	371.24	372.71	370.03
12	351.26	362.97	369.16	357.01
Monthly Mean	370.02	371.00	370.89	371.36

342Table 8: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Anyigba for a343period of 1999 – 2018

Months	N _{itu-r}	N _{Rh,t(Anyigba)}	N _{t(Anyigba)}	N _{Rh(Anyigba)}
1	310.60	346.72	367.68	321.72
2	354.19	355.95	354.96	350.57
3	363.88	359.09	354.09	357.18
4	364.82	359.42	354.03	357.86
5	368.71	364.15	356.72	364.10
6	370.20	370.81	361.79	371.71
7	369.46	372.34	363.35	373.10
8	369.19	372.73	363.75	373.44
9	370.60	372.72	363.02	374.10
10	371.87	370.76	360.95	372.39
11	361.19	363.48	359.02	360.73
12	330.12	353.96	364.45	338.10
Monthly Mean	358.74	363.51	360.32	359.58

Table 9: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Jalingo for a period of 1999 – 2018

Months	N _{itu-r}	N _{Rh,t(Jalingo)}	$N_{t(Jalingo)}$	N _{Rh(Jalingo)}	
1	283.17	321.98	355.74	288.24	
2	307.86	312.42	319.95	304.85	
3	322.58	312.11	310.83	313.29	
4	341.74	320.35	312.27	328.20	
5	359.84	341.17	329.70	352.21	
6	362.29	348.49	337.31	359.18	
7	365.60	357.57	345.92	368.63	
8	364.25	358.98	348.25	369.13	
9	366.15	357.76	345.61	369.34	
10	362.14	348.04	336.68	358.91	
11	327.98	330.85	334.67	326.79	
12	301.62	328.06	349.64	306.36	
Monthly Mean	338.77	336.48	335.55	337.09	

355 Table 10: Monthly Estimates of Surface Refractivity from Linear Meteorological Models at Sokoto for a 356 period of 1999 – 2018

Months	N _{itu-r}	N _{Rh,t(Sokoto)}	N _{t(Sokoto)}	N _{Rh(Sokoto)}	
1	276.13	289.67	300.32	278.81	
2	283.63	300.08	314.14	285.76	
3	273.52	297.91	320.64	274.88	
4	290.79	306.81	326.21	287.09	
5	331.74	321.98	326.31	317.33	
6	347.40	329.10	320.51	337.40	
7	355.37	334.67	315.33	353.74	
8	363.85	341.26	311.48	370.79	
9	365.93	342.13	311.97	372.04	
10	353.01	334.84	313.03	356.39	
11	298.71	306.43	307.20	305.44	
12	289.03	299.13	302.38	295.66	
Monthly Mean	319.09	317.00	314.13	319.61	

Table 11: Mean Surface Refractivity Values from ITU-R and Linear Meteorological Models for a period of 1999 - 2018

Stations	N _{itu-r}	N _{Rh,t}	Nt	N _{Rh}	
Yenagoa	381.87	383.21	383.27	383.16	
Abakaliki	370.75	352.01	372.89	369.35	
Ibadan	370.02	371.00	370.89	371.36	
Anyigba	358.74	363.51	360.32	359.58	
Jalingo	338.77	336.48	335.55	337.09	
Sokoto	319.09	317.00	314.13	319.61	

368 5.3 Validation and Comparison of Linear Meteorological Models with ITU-R Equation

The results of surface refractivity estimated from different linear meteorological models were compared in scattered plots to establish levels of disparity in the values obtained. The graphs are presented in figures 2 – 7 showing the line of best fit. Surface refractivity values from N_{itu-r} was compared with values obtained from N_{Rh} , N_t and $N_{Rh,t}$ being the contributions of relative humidity, temperature and combination of relative humidity and temperature to linear meteorological models. This was to determine the correlation coefficient and the standard deviation of the surface refractivity from ITU-R results at six locations - 3 from the south and 3 from the north - over Nigeria and the results are shown in table 12.

- 377 In Yenagoa, contributions of Rh and t (N_{Rh.t}) had the best linear correlation of 0.98 and standard 378 deviation of ±0.32 while temperature contribution (Nt) had the least correlation of 0.23 and standard 379 deviation of ±1.09. The contributions of relative humidity (N_{Rh}) had a dominance on temperature with 380 correlation of 0.88 and standard deviation of ±1.73. Results from Abakaliki and Ibadan were similar to 381 that of Yenagoa, correlation of N_{Rh.t} with N_{itu-r} was 0.99 at Abakaliki and Ibadan while the standard deviation was ±0.44 and ±0.49 respectively. Similarly, N_{Rh} had a dominance over N_t having correlation 382 of 0.97 and 0.95 at Abakaliki and Ibadan while N_{t} had correlation of 0.62 and 0.65 respectively. High 383 384 relative humidity and low temperature values observed in the 3 southern stations - Yenagoa, Abakaliki 385 and Ibadan – significantly enhanced the suitability of N_{Rht} linear meteorological model for estimation of surface refractivity values comparable to N_{itu-r}. 386
- 387 The remaining 3 stations in the north – Anyigba, Jalingo and Sokoto – had divergent trend from 388 the southern stations. Surface refractivity values obtained from the contributions of temperature (N_t) 389 was extremely low and mostly negative. Results from Nt at Anyigba, Jalingo and Sokoto had correlation -390 0.46, -2.99 and 0.17 while the standard deviation was ±4.21, ±15.53 and ±8.64 respectively. Contrary to the observations in the southern stations, in the northern stations, relative humidity contributions (N_{Bh}) 391 392 had the best linear correlation having 0.96 at Anyigba, 0.98 at Jalingo and Sokoto. Standard deviation was given as ±4.71 at Anyigba, ±6.39 at Jalingo and ±7.17 at Sokoto. The trend in the north was as result 393 394 of low relative humidity and high temperature.
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Figure 4: Correlation between ITU-R and Linear Meteorological Models Values for Ibadan















- Figure 6: Correlation between ITU-R and Linear Meteorological Models Values for Jalingo



Figure 7: Correlation between ITU-R and Linear Meteorological Models Values for Sokoto
 Figure 7: Correlation between ITU-R and Linear Meteorological Models Values for Sokoto
 Figure 7: Correlation between ITU-R and Linear Meteorological Models Values for Sokoto
 Table 12: Correlation Coefficient (r) and Standard Deviation (o) between Surface Refractivity Values estimated from ITU-R and Linear Meteorological Models

	N _{Rh,t}		Nt		N _{Rh}	
Stations	r	σ	r	σ	r	σ
Yenagoa	0.98	0.32	0.23	1.09	0.88	1.73
Abakaliki	0.99	0.44	0.63	2.83	0.97	3.52
Ibadan	0.99	0.49	0.65	1.84	0.95	2.78
Anyigba	0.86	4.64	-0.46	4.21	0.96	4.71
Jalingo	0.81	10.83	-2.99	15.53	0.98	6.39
Sokoto	0.99	2.91	0.17	8.64	0.98	7.17

438 **6.0 Conclusion**

439 The problem of quantifying surface refractivity as a function of atmospheric variables in various 440 climatic situations is very complex because the three meteorological variables involved are neither 441 perfectly correlated nor totally uncorrelated. In this study, new sets of linear expressions were obtained 442 in terms of meteorological parameters from six locations across Nigeria as alternate means of estimating 443 surface refractivity. The linear expressions were validated using data obtained for a period of 20 years 444 and compared with results from existing equations. Linear expressions as a function of relative humidity 445 and temperature were best suitable for stations in the south with minimum correlation of 0.98 while 446 linear expressions as a function of relative humidity only were best suitable for the stations in the north 447 with minimum correlation of 0.96.

The sensitivity of the parameters showed that surface refractivity is much more likely to vary due to moisture changes in the atmosphere than temperature changes. In the northern Nigeria, relative humidity changes near the surface have more than twice the impact on refractivity as temperature changes. This is partly a result of the high variability of humidity and partly due to the inherent sensitivity of refractivity to moisture in the additional wet term, which is also a maximum in wet months.

453 At low temperatures, surface refractivity varies very little with relative humidity because the 454 saturation vapour pressure is low: as air temperatures increase, the saturation vapour pressure rises, 455 expanding the range of possible refractivity values. Hence, estimation of surface refractivity from linear 456 meteorological models is mainly a function of low temperature and high relative humidity in the south 457 and a function of relative humidity in the north.

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460 **7.0 References**

- 1. Amajama, J., and Eshiet, M. (2016). Impact of weather components on (UHF) radio signal. International Journal of Engineering Research and General Science (IJERGS), 4(3), 474-480.
- Joseph, A. (2015) Mathematical Relationships between Radio Refractivity and Its Meteorological
 Components with A New Linear Mathematical Equation to Determine Radio Refractivity.
 International Journal of Innovative Science, Engineering & Technology (IJISET), Vol. 2 Issue 12,
 December, 2015
 - 3. Jari, L., and Ismo, H. (2015). Effect of temperature and humidity on radio signal strength in outdoor wireless sensor. In Proceedings of the federated conference on computer science and information systems (Vol. 5, pp. 1247-1255).
 - Enyenihi H. J., Simeon Ozuomba, Kalu Constance (2019): Development of Model for Estimation of Radio Refractivity from Meteorological Parameters. Universal Journal of Engineering Science 7(1): 20-26, DOI: 10.13189/ujes.2019.070103
 - 5. ITU-R (2013): "Propagation data and prediction methods required for the design of terrestrial line-of-sight systems," Geneva, July 2013, pp. 530
- 4806. Stephen J. S., Hedley Hansen and Derek Abbott (2014): Prediction of Surface Refractivity481Gradient Distributions, from Weather Station Surface Data. Conference Paper,482https://www.researchgate.net/publication/272093235; DOI: 10.1109/EuCAP.2014.6901750
- 484
 Ayantunji, B. G. and Okeke, P. N. (2011): Diurnal and Seasonal Variation of Surface Refractivity
 485 over Nigeria. Progress in Electromagnetics Research B, Vol. 30, 201 222, 2011

- 487 8. ITU-R (2003): The radio refractive index: Its formula and refractivity data," pp. 453 459
- 488

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- 489
 9. Odjugo, P. A. O. (2013): Analysis of climate change awareness in Nigeria. Scientific Research and
 490 Essays, Vol. 8 (26), 1203 1211.
- 492 10. Nwagbara, M. O., Ijeoma, M. A., Chima, G. N. (2009): Climate change and flood in Northern
 493 Nigeria. An examination of Rainfall Trends over the region". In R. N. C. Anyadike, I. M. Madu and
 494 C. K Ajaero, Eds. Climate change and the Nigeria Environment, pp. 525 –538.
- 495 11. Maria D. S., Diego, B. and Giovanni, S. (2013): Measuring the effect of climate change on agriculture: A literature review of analytical models. Journal of Development and Agricultural
 497 Economic Vol. 5 (2), pp. 499 509.
- 499 12. James, F. (2009): Climate change: How o report the story of the century.
 500 http/www.scidev.met/ent. Practical guide/ climate. Change -how-to-report-the- storyof-the
 501 cent.html.
 - 13. Abu B (2007): Sea level rise and the Niger Delta of Nigeria. Journal of Wetland, 3(1): 44-52
- 50514. Intergovernmental Panel on Climate (IPCC) (2007): Climate change 2007. The fourth assessment506report (AR4). Synthesis report for policy makers. Http://www.ipcc.ch/pdf/assessment-507report/ar4/syr/ar4_syr_spm.pdf. (Access 10th August, 2009).
- 509 15. Jagtap, S (2007): Managing vulnerability to extreme weather and climate events:
 510 Implications for agriculture and food security in Africa. Proceedings of the Inter-national
 511 Conference on Climate Change and Economic Sustainability held at Nnamdi Azikiwe
 512 University, Enugu, Nigeria. 12-14, June 2007.
 - 16. Ayoade, J. O (2004): Climate Change. Ibadan. Vantage Publishers, pp. 45-66
- 516 17. Nwafor J. C., (2007): Global climate change: The driver of multiple causes of flood intensity
 517 in Sub-Saharan Africa. Paper presented at the International Conference on Climate Change
 518 and Economic Sustain-ability held at Nnamdi Azikiwe University, Enugu, Nigeria, 12-14
 519 June 2007.
 - Ayuba H. K, Maryah U. M, and Gwary D. M (2007): Climate change impact on plant species composition in six semi-arid rangelands of Northern Nigeria. Nigerian Geographical Journal 5(1): 35-42
- 525 19. Colarco, P. R., A. da Silva, M. Chin, and T. Diehl, (2010): Global Aerosol Distributions in the NASA
 526 GEOS-4 Model and comparisons to Satellite and Ground-based Aerosol Optical Depth. Journal of
 527 Geophysical Research, 115, DOI: 10.1029/2009JD012820
- 52920. Bacmeister, J. T., Suarez, M. J. and Robertson, F. R. (2006): Rain Re-evaporation, Boundary-530Layer/Convection Interactions and Pacific Rainfall Patterns in an AGCM, Journal of Atmospheric531Science, 63, pp. 3383-3403.