

**SYNOPTIC ANALYSIS OF A PERIOD WITH ABOVE-NORMAL PRECIPITATION
DURING THE DRY SEASON IN SOUTHEASTERN BRAZIL**

ABSTRACT: This study presents an analysis of a period with positive precipitation anomalies and anomalous severe weather activity in parts of the Southeastern Brazil during the dry season (austral winter). The objective of this work is to identify the synoptic pattern associated with the severe episodes on August 2018. The analysis was based on observational data and the CFSR (Climate Forecast System Reanalysis). Standardized anomalies of the main meteorological variables were used in the analysis. An anomalous trough in the mid and upper troposphere and the associated low-level flow from the Amazon Basin to Southeastern Brazil provided high amount of precipitable water, and were the main responsible for the rainfall markedly above the climatology along with severe hail and strong winds. The atmospheric environment was conditionally unstable due to a stationary front. These type of analyzes should be taken into consideration at this time of year to improve weather forecasts and minimize impacts in such circumstances.

KEY-WORDS: Brazil, Extreme Event, Precipitation, Synoptic.

1. INTRODUCTION

Brazil has a wide variety of climates, especially due to its extensive territory (about 8.5 million km²), topography, vegetation and dynamics of air masses and sea currents (Mendonça and Danni-Oliveira, 2007; Zavattini and Boin, 2013). The Southeastern Brazil is the most economically developed region of the country. The disorderly population growth has been contributing to many environmental disasters occurring in the region, associated mainly with high rainfall volumes (Coelho Neto *et al.*, 2009; Ávila *et al.*, 2016), caused by phenomena such frontal systems, South Atlantic Convergence Zones, troughs in the middle troposphere and mesoscale convective complexes.

33 The tropical climate predominates in most of Southeastern Brazil, which is
34 characterized by rainy summers and dry winters (Torres and Machado, 2011;
35 Alvares *et al.*, 2013). This region presents high interannual variation of precipitation
36 (Cavalcanti *et al.*, 2009) and extreme weather events are common between October
37 and March (austral spring and summer; Grimm, 2011). However, anomalous
38 atmospheric configurations between April and September (austral autumn and
39 winter) can also lead to severe weather episodes, although they are rarer in
40 comparison to the warm season.

41 In the first half of August 2018, in the middle of the dry season in
42 Southeastern Brazil, part of the region recorded above-average precipitation in
43 several cities. Some cities of the States of São Paulo and Minas Gerais registered
44 precipitation nearly three times the August monthly mean in just three consecutive
45 days. Severe storms also occurred during this period, in particular between August
46 1st and 10th and were responsible for high winds and hail.

47 In this context, the objective of this study is to perform a synoptic analysis of
48 the period and the causes of precipitation above average in the region during the
49 peak of the dry season. The importance of this study lies in the fact that severe
50 weather events during the dry period, when high atmospheric instability is unlikely,
51 can cause greater impacts to the affected areas, especially to agriculture and
52 general population. In addition, events like this may compromise climatic
53 predictions for the period, since accumulated rainfall in the order of 40 mm to 80
54 mm, or even greater than 100 mm, can be recorded in only a few days or hours of
55 the dry season. Thus, when the conditions favoring precipitation stall over a given
56 region for a few days, several locations can easily surpass the monthly mean
57 precipitation.

58 Therefore, the identification of atypical synoptic-scale patterns can be used
59 to forecast such events and take preventive actions, minimizing impacts (Dolif and
60 Nobre, 2012). Although these extreme events are not seen in a seasonal climate
61 forecast, according to Coelho and Costa (2010), moving toward a more practical
62 purpose of seasonal climate forecasts to uphold decision making depends on more
63 valuable synergy among climate scientists and decision makers.

64

65 2. METHODOLOGY

66 2.1. Study Area

67 The analyzes of this study focus on the Southeastern Brazil (Figure 1), in
68 particular the States of São Paulo and Minas Gerais. According to the Brazilian
69 Institute of Geography and Statistics (IBGE), this area of the country has a
70 population of approximately 70 million inhabitants, 42% of the entire Brazilian
71 population. In addition, about 90% of the population of this portion of Brazil lives in
72 urban areas. The Southeastern region concentrates important elements of the
73 economy, such as agriculture, energy, livestock, petrochemical and steel industries
74 and tourism. In this scenario, many people and sectors are vulnerable to natural
75 disasters.



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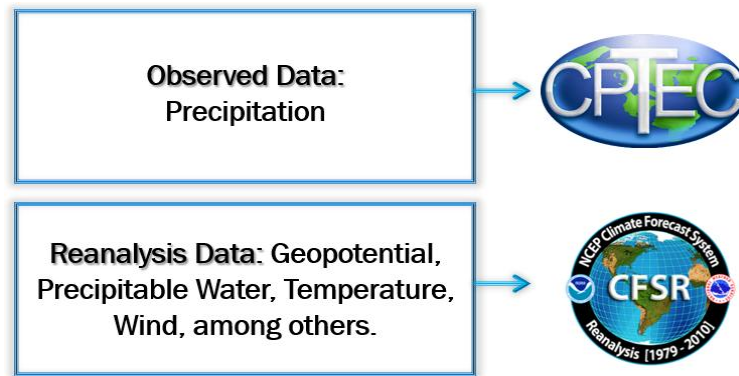
77 **Figure 1** – Brazil with the Southeastern region highlighted in red. The respective States are
78 Espírito Santo (ES), Minas Gerais (MG), Rio de Janeiro (RJ) e São Paulo (SP).

79

80 2.2. Data

81 Rainfall data were extracted from the Center for Weather Forecasting and
82 Climate Studies / National Institute for Space Research (CPTEC/INPE) website –
83 <http://clima1.cptec.inpe.br/monitoramentobrasil/pt> (Figure 2). Data used include
84 information from the National Institute of Meteorology (INMET – SYNOP and
85 automatic meteorological stations), as well as the records of state meteorological
86 centers. Data are interpolated in a grid of 1.0° x 1.0° resolution using the objective

87 analysis method of Cressman (1959). The anomaly maps represent the point-to-
 88 point difference between the variable recorded in the selected month and year (in
 89 this research, precipitation) and the climatology (historical average of the period
 90 from 1979 to 1995). Severe weather reports were obtained from a database
 91 maintained by the third author of this research.



92
 93 **Figure 2** – Data sources used in this work.

94 The standardized anomalies of some atmospheric variables were calculated
 95 with the Climate Forecast System Reanalysis Version 2 (CFSR; Saha *et al.*, 2014)
 96 data (Figure 2), from the National Center for Environmental Prediction (NCEP). The
 97 CFS is a model representing the global interaction between Earth’s oceans, land
 98 and atmosphere. The CFS reanalysis is generated with a T382 model (~ 38 km
 99 horizontal resolution, interpolated to 0.5° x 0.5°), 27 vertical levels between 1000
 100 and 100 hPa and available at the synoptic times (0000, 0600, 1200 and 1800
 101 UTC). The atmospheric fields analyzed were geopotential, precipitable water,
 102 temperature, wind, among others. In order to investigate a particular case of a
 103 severe storm during the analyzed period, some variables, instability indices (Craven
 104 and Brooks, 2004) and satellite images were also examined in the Southeastern
 105 region of Brazil.

106 To produce the standardized anomalies from the CFSR data, a calculation is
 107 made using the mean and standard deviation from the CFSR of each variable on all
 108 days of the year during the 32-year reanalysis series (1979 – 2010). Then, the
 109 CFSR data for the studied period is transformed into an anomaly (the difference
 110 between the observed value and the climatological mean) that is proportional to the
 111 standard deviation of the variable at each location and at each time, as:

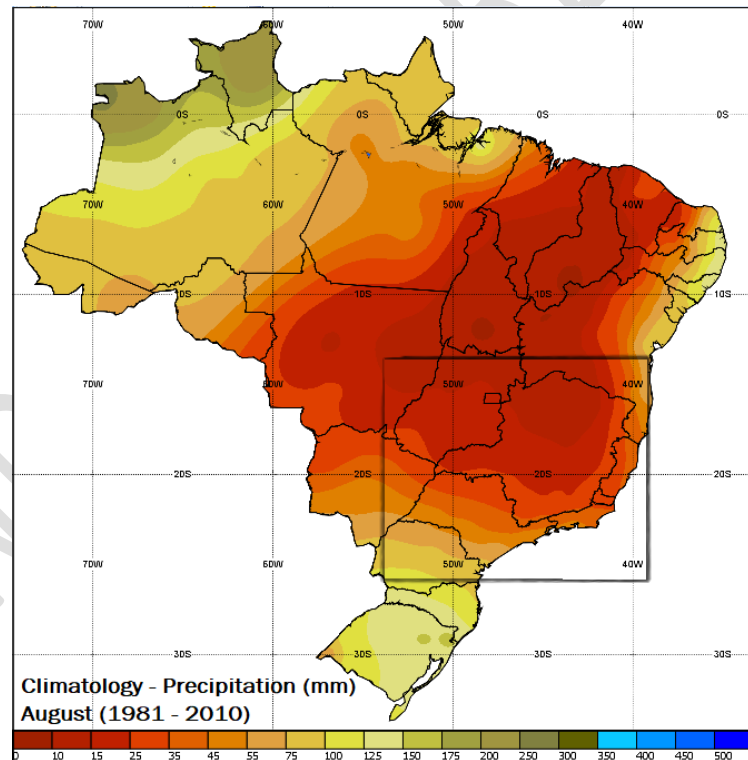
$$112 \text{ Standard Anomalie} = \frac{\text{Observed Value} - \text{Climatological Mean}}{\text{Standard Deviation}} . \quad (1)$$

113 Thereby, a standard deviation of +2, for instance, means the variable has a value
114 that is 2 standard deviations above the mean for that time of the year in that region.

115

116 3. RESULTS

117 Figure 3 presents the average August precipitation for the period from 1981
118 to 2010 over Brazil, according to INMET. In Southeastern Brazil, the August average
119 rainfall varies from just 10 mm in the north of the region to 50 mm in the south.
120 This situation is characteristic of the peak of the dry season in this area, which
121 extends from April to September. At this time of year, the South Atlantic Subtropical
122 Anticyclone (SASA) reaches its westernmost position over the continent, greatly
123 reducing the precipitation. Rainfall events mostly occur when frontal systems and
124 subtropical and extratropical cyclones are strong enough to displace the SASA
125 (Reboita *et al.*, 2010).



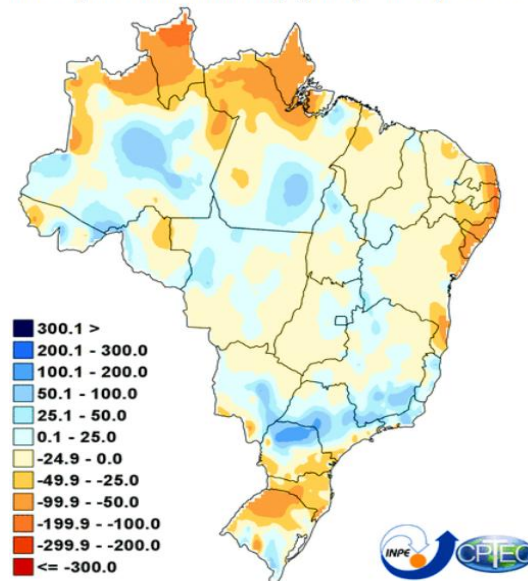
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127 **Figure 3** – Mean precipitation for August from 1981 to 2010 over Brazil. The gray box
128 delimits the Southeastern region of the country (Source: INMET).

129

130 The rainfall anomalies for August 2018 (Figure 4) show positive precipitation
131 anomalies appear especially in the State of Paraná (PR, Southern Brazil) and part of
132 Southeastern Brazil with values up to 200 mm above the climatology (about four
133 times the average). These rainfall volumes were associated with the most
134 significant rain episodes that occurred in the first 10 days of the month. However,
135 the common cold fronts that pass through this area at this time of year usually do
136 not have enough moisture to cause such expressive precipitation (Siqueira and
137 Machado, 2004).

Precipitation Anomaly (mm) - AUG/2018



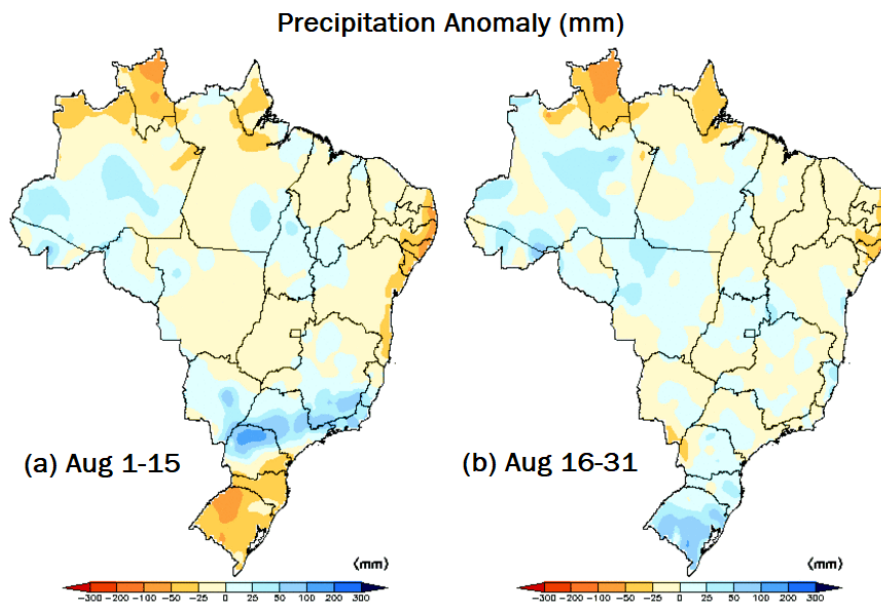
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139 **Figure 4** - Rainfall anomaly (mm) on August 2018 over Brazil (Source: CPTEC/INPE).

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141 Figure 5 elucidates the differences between the first and second half of
142 August, regarding the anomalies of the precipitation field, which leads us to focus
143 our analysis in the period from August 1st to 10th, when most rainfall was
144 concentrated.

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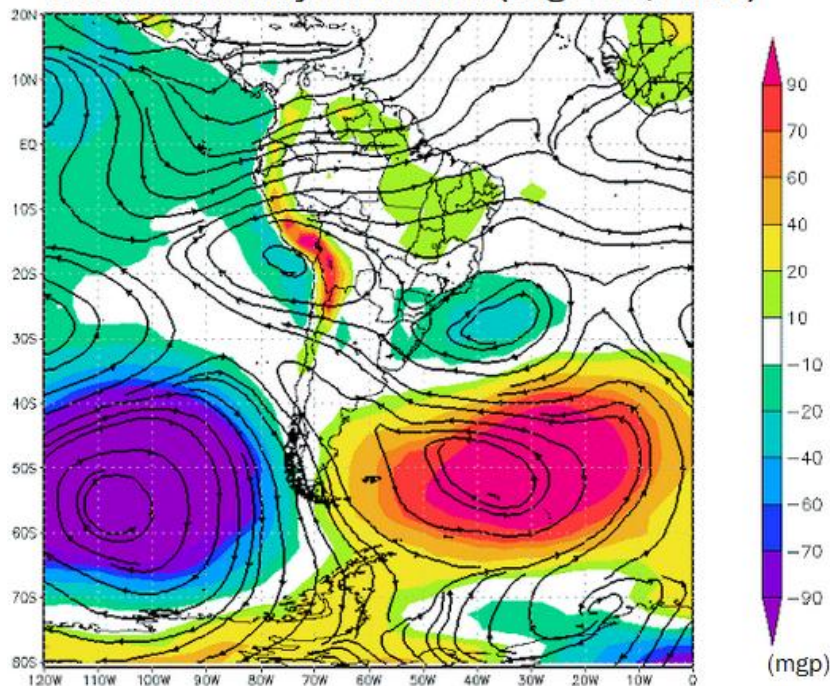
147 **Figure 5** – Precipitation anomalies over Brazil in the (a) first half and (b) second half of
 148 August 2018 (Source: CPTEC/INPE).

149

150 The 850-hPa geopotential height anomalies (Figure 6) were in phase with
 151 cyclonic circulation at 250 hPa along the coast of Southeastern Brazil in August
 152 2018, which relates to the passage of troughs and frontal systems over the region.
 153 Souza and Ambrizzi (2004) emphasize that zonal westerly flow occurs over most
 154 subtropical South America during this period. The inland penetration of the SASA is
 155 recurrent in winter and prevents the meridional displacement of frontal systems
 156 moving towards the tropical areas of South America (Ito, 1999). Also, this cyclonic
 157 upper-level circulation, in association with an anomalous 850-hPa anticyclone
 158 farther south, forms a dipole blocking system in the Atlantic Ocean. The typical
 159 winter configuration resumed after the advance of cold air over the region at the
 160 end of the first half of August.

161

Geopotential anomaly at 850 hPa and Stream function anomaly at 250 hPa (Aug 1-15, 2018)



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163 **Figure 6** – 850-hPa geopotential height (shaded, m gp) and 250-hPa stream function
164 anomalies during the first half of August 2018 (Source: CPTEC/INPE).

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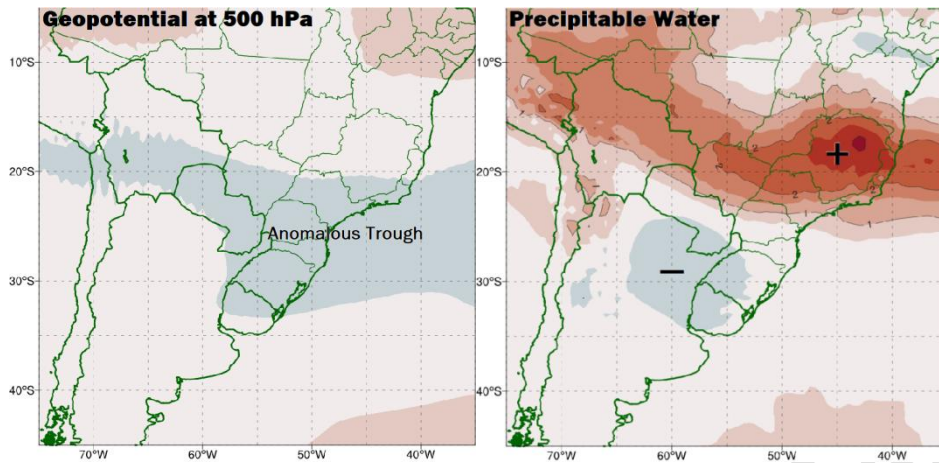
166 The standardized anomalies (Figure 7a) depict an anomalous 50-hPa trough
167 over south-central Brazil, including part of the Southeastern region. Above-average
168 precipitable water over the north-central São Paulo and the entire Minas Gerais
169 (Figure 7b) associated with the SASA's circulation and the constant transport of
170 moisture from the Amazon Basin by the South American Low-Level Jet (SALLJ;
171 Oliveira *et al.*, 2018). According to flow patterns and moisture transport values in
172 the lower atmosphere, the SALLJ transports atmospheric moisture from the trade
173 winds flowing over the Amazon, which can obtain a greater amount of moisture due
174 to evapotranspiration in the forest region. Then, this flux suffers a change of
175 direction due to the topographical block and runs parallel to the Andes towards the
176 Southern/Southeastern Brazil (Marengo *et al.*, 2004). The anomalous northwesterly
177 flow during this period is evidenced by the standardized anomalies of 850-hPa
178 zonal (Figure 7c) and meridional (Figure 7d) wind, which respond to the anomalous
179 low-pressure over the Atlantic Ocean (Figure 6).

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(a)

(b)

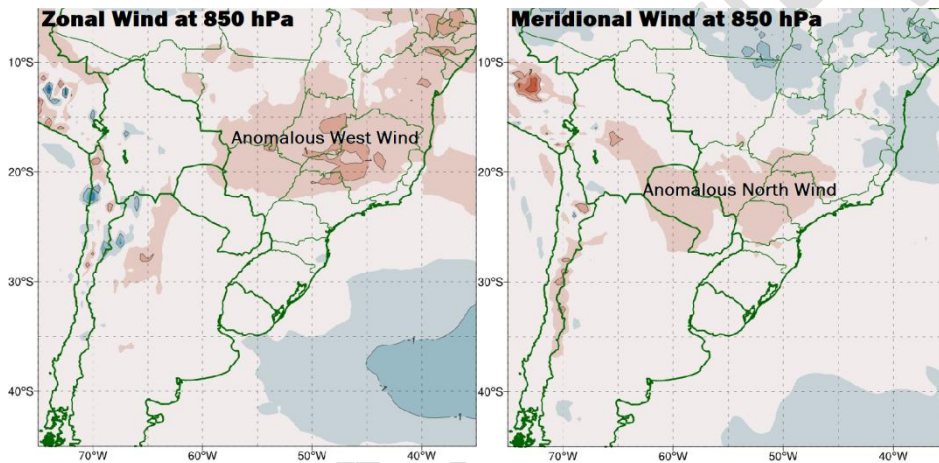


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(c)

(d)

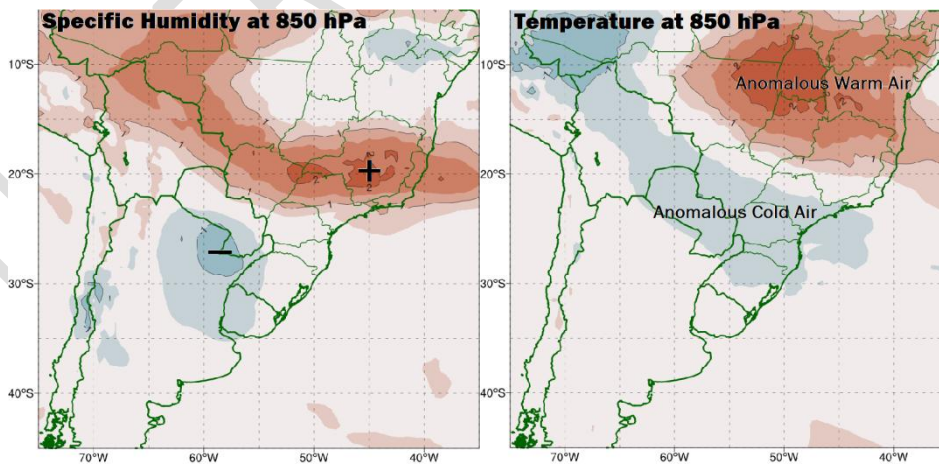


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(e)

(f)



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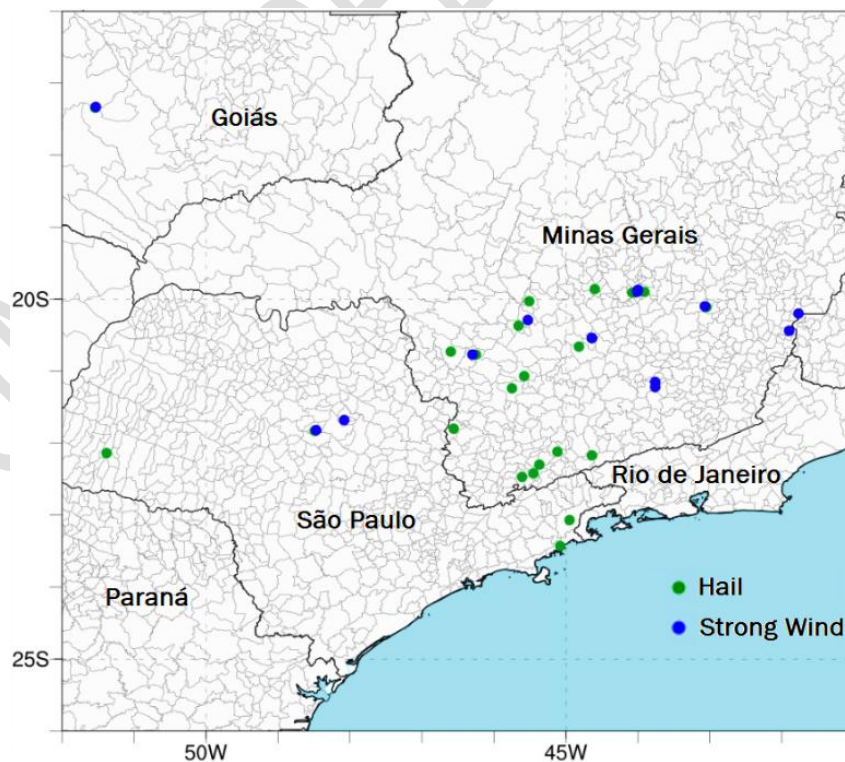


188 **Figure 7** – Standardized anomalies of (a) geopotential at 500 hPa, (b) precipitable water,
189 (c) specific humidity at 850 hPa, (d) temperature at 850 hPa, (e) zonal wind at 850 hPa and
190 (f) meridional wind at 850 hPa between August 1-10, 2018 over center of South America.

191 The standardized anomalies of 850-hPa specific humidity (Figure 7e) follow
192 the precipitable water anomalies and are positive from Northern Brazil to the
193 Southeastern. Above-average 850-hPa temperatures are observed over the
194 northern portion of Southeastern Brazil and below-average in the central-southern
195 sector of the country (Figure 7f), with enhanced baroclinicity in the region where the
196 anomalous precipitation occurred.

197 One way to analyze the intraseasonal variability of SALLJ is by the frequency
198 of rainfall extremes and their connection to the jet events. Several researchers have
199 shown that such intense precipitation episodes associated with the SALLJ are much
200 more common during the austral summer (e.g., Liebmann *et al.*, 2004; Weykamp
201 and Ambrizzi, 2006), since there is sufficient warm and moist air for the
202 development of severe storms, differently to what occurs in the austral winter.

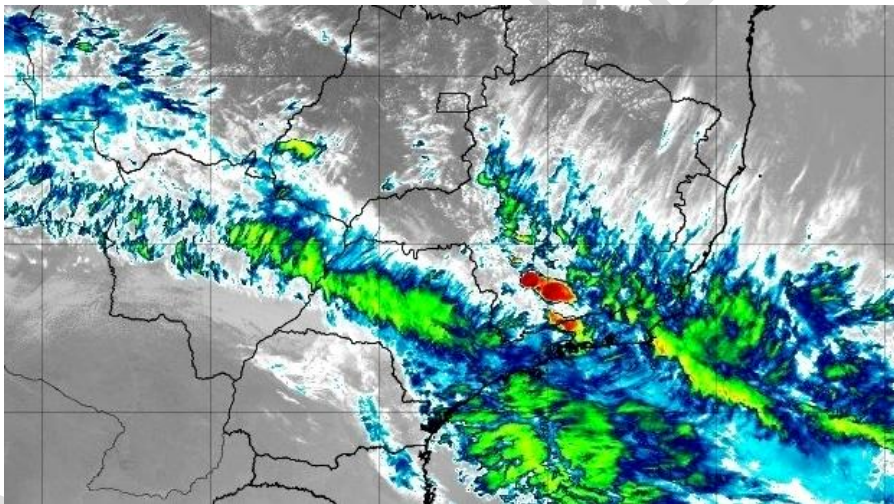
203 The reports of severe hail and wind between August 1st and 10th, mainly in
204 the south of the State of Minas Gerais, northeast of São Paulo and Paraíba Valley,
205 are shown in Figure 8. According to the National Institute of Meteorology, the hail
206 occurred in several cities of Minas Gerais on August 6, 2018, between afternoon
207 and evening, was the only report in this month since 1910.



208

209 **Figure 8** – Severe weather reports over Southeastern Brazil between August 1st and 10th
210 (2018) with green dots representing hail and blue dots strong winds.

211 The severe weather event of August 6, 2018 in southern Minas Gerais is
212 shown in the satellite image at 1900 UTC of Figure 9, where it is noted the presence
213 of a stationary front in the Southeastern Brazil. The atmospheric environment was
214 conditionally unstable in this region at 1800 UTC this day, especially due to the high
215 850-hPa θ_e (Figure 10a) associated with a warm and moist lower troposphere. The
216 moisture convergence is noticed in the 850-hPa θ_e field between São Paulo and
217 Minas Gerais states. The instability is further evidenced by the most-unstable
218 convective available potential energy (MUCAPE) values greater than 2000 J/kg in
219 this area (Figure 10b), which is very atypical for this time of the year. The 0-6-km
220 wind shear is also high (more than 15 m/s) in association with the strong upper-
221 level jet nearly parallel to the stationary front during the event. The northwesterly flow
222 from the SALLJ for several consecutive days contributed to the high MUCAPE values,
223 thus allowing gradual warming and increased humidity of the air mass in the north
224 of the frontal system.

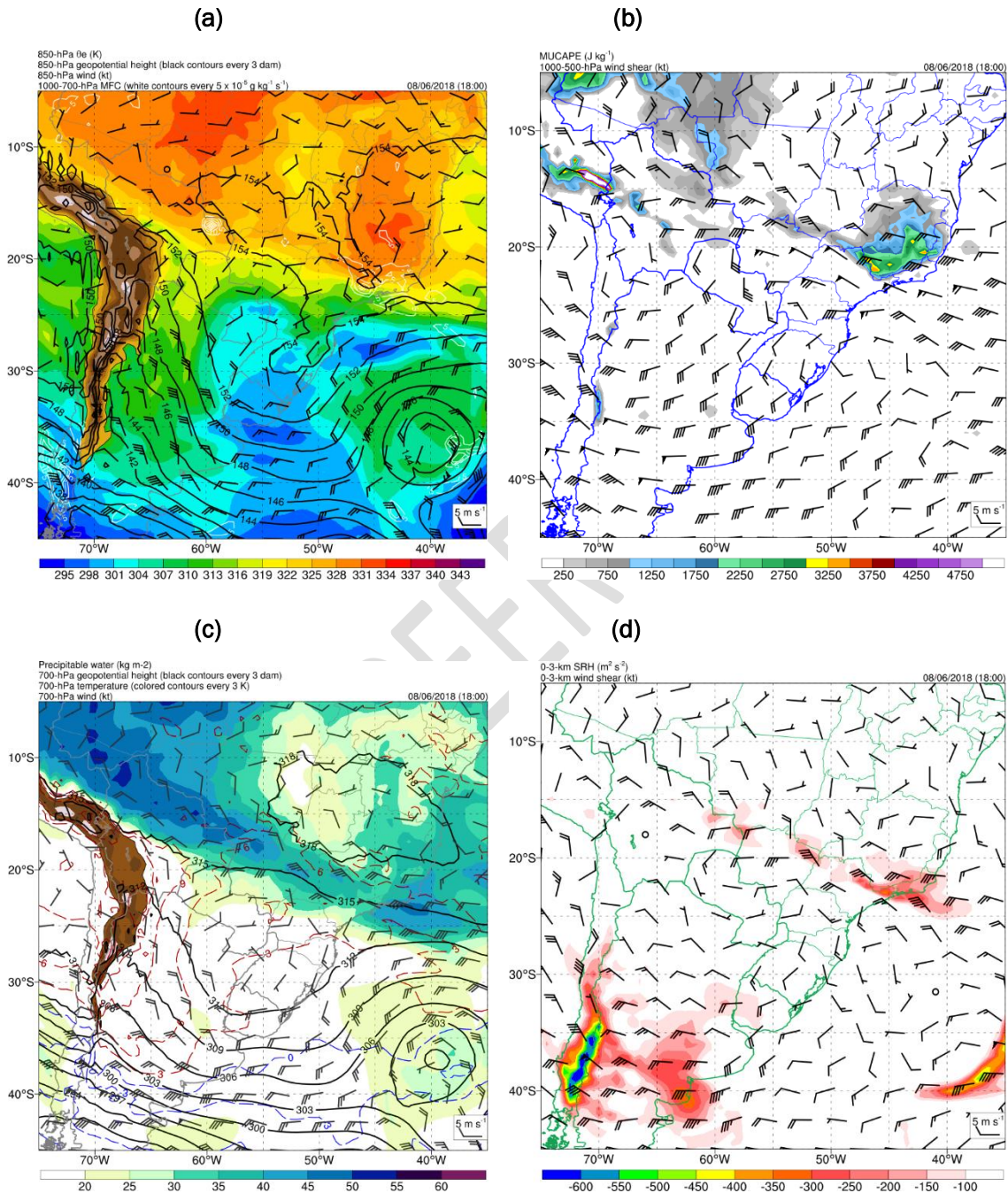


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226 **Figure 9** – Infrared channel enhanced image of the Geostationary Operational
227 Environmental Satellite – GOES 16 over the Southeastern Brazil on August 6, 2018 at
228 1900 UTC (Source: CPTEC/INPE).

229
230 Convective events are strongly modulated by the total water vapor found at
231 700 hPa due to the thermodynamic instability, which is clear by the expressive
232 precipitable water available between the States of Minas Gerais and São Paulo
233 (Figure 10c). The 0-3-km storm-relative helicity (SRH) is greater than 200 m²/s²
234 along and north of the stationary front (Figure 10d), which is sufficient for the

235 development of strong atmospheric instabilities. The stationary front in association
 236 with the anomalous baroclinicity between MG and SP was related to high wind
 237 shear in the region. This factor linked to the air mass with high MUCAPE generates a
 238 favorable environmental for severe storms.

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243 **Figure 10** – (a) 850-hPa θ_e (shaded, K), geopotential heights (black contours, mgp) and
 244 winds, and MFC (white contours every $2 \times 10^{-5} \text{ g kg}^{-1} \text{ s}^{-1}$), (b) MUCAPE (shaded, J/kg) and
 245 wind shear between 1000 and 500 hPa, (c) PW (shaded, mm), 700-hPa geopotential
 246 height, temperature and wind, and (d) 0-3-km SRH and wind shear on August 6, 2018 at
 247 1800 UTC.

248 **4. CONCLUSIONS**

249 An atypical atmospheric pattern occurred in the first half of August 2018 and
250 culminated in above-normal precipitation and severe weather in areas of the
251 Southeastern Brazil (middle of the dry season). In short, this event occurred
252 specially as a function of an anomalous synoptic-scale stationary trough over south-
253 central Brazil. Coupled with this, there was the displacement of troughs in the mid
254 troposphere combined with an expressive positive precipitable water anomaly.
255 Moreover, the flow at low-levels was favorable to the moisture transport and
256 convergence directed to Southeastern Brazil. The SALLJ performance is more active
257 in the austral summer months due to the predominant absence of SASA on the
258 continent, thus allowing the jet to be established. In austral winter, this scenario is
259 much more unusual, but still requires careful monitoring.

260 It should be noted that this kind of analysis permeates all the time scales
261 that deserve attention in the management and operation environment of several
262 sectors. In terms of climatology, it was observed that rainfall volumes were
263 established above the climatology and also beyond what was predicted on the
264 seasonal scale. In addition, this synoptic-scale atmospheric pattern can present its
265 first evidence about 7 to 10 days in advance, thus alerting that a phenomenon with
266 potential for anomalous severity has a high probability of occurrence. Finally, once
267 the configuration is established, there is continuous monitoring of the
268 meteorological variables and atmospheric instability indices, a process carried out
269 with particular caution in nowcasting. In this sense, it is essential to emphasize the
270 relevance that short-term meteorological phenomena can exert on the historical
271 averages and seasonal forecasting.

272 From the synoptic point of view, this atmospheric configuration in the dry
273 season has expressive intensity with respect to extreme weather episodes, despite
274 the low frequency at this time of year (austral winter). This type of investigation has
275 great importance, so that such severe events could be better identified in weather
276 forecasting. Therefore, more detailed studies are needed to examine other possible
277 dynamic and thermodynamic mechanisms involved in these meteorological
278 situations. In this way, it will be also possible to investigate its climatology and
279 trends over the last decades.

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- 282 Alvares, C. A.; Stape, J. L.; Sentelhas, P. C.; Gonçalves, J. L. M.; Sparovek, G.
283 Köppen's Climate Classification Map for Brazil. *Meteorologische Zeitschrift*, v. 22, n.
284 6, p. 711-728, 2013.
- 285 Ávila, A.; Justino, F.; Wilson, A.; Bromwich, D.; Amorim, M. Recent Precipitation
286 Trends, Flash Floods and Landslides in Southern Brazil. *Environmental Research*
287 *Letters*, v. 11, n. 11, p. 1-13, 2016.
- 288 Cavalcanti, I. F. A.; Ferreira, N. J.; Justi da Silva, M. G. A.; Silva Dias, M. A. F. *Tempo e*
289 *Clima no Brasil*. Oficina de Textos Publisher, 463 pp., 2009.
- 290 Coelho, C. A. S.; Costa, S. M. S. Challenges for Integrating Seasonal Climate
291 Forecasts in User Applications. *Current Opinion in Environmental Sustainability*, v. 2,
292 n. 5-6, p. 317-325, 2010.
- 293 Coelho Neto, A. L.; Avelar, A. S.; Lacerda, W. A. Landslides and Disasters in
294 Southeastern and Southern Brazil. *Developments in Earth Surface Processes*, v. 13,
295 p. 223-243, 2009.
- 296 Craven, J. P.; Brooks, H. E. Baseline Climatology of Sounding Derived Parameters
297 Associated with Deep Moist Convection. *National weather Digest*, v. 28, p. 13-24,
298 2004.
- 299 Cressman, G. W. An Operational Objective Analysis System. *Monthly Weather*
300 *Review*, v. 87, p. 367-374, 1959.
- 301 Dolif, G.; Nobre, C. Improving Extreme Precipitation Forecasts in Rio de Janeiro,
302 Brazil: are Synoptic Patterns Efficient for Distinguishing Ordinary from Heavy Rainfall
303 Episodes? *Atmospheric Science Letters*, v. 13, p. 216-222, 2012.
- 304 Grimm, A. M. Interannual Climate Variability in South America: Impacts on Seasonal
305 Precipitation, Extreme Events, and Possible Effects of Climate Change. *Stochastic*
306 *Environmental Research and Risk Assessment*, v. 5, n. 4, p. 537-554, 2011.
- 307 Ito, E. R. K. Um Estudo Climatológico do Anticiclone Subtropical do Atlântico Sul e
308 sua Possível Influência em Sistemas Frontais. Masters Dissertation, São Paulo
309 University, 1999.

310 Liebmann, B.; Kiladis, G. N.; Vera, C. S.; Saulo, A. C.; Carvalho, L. M. V. Subseasonal
311 Variations of Rainfall in South America in the Vicinity of the Low-Level Jet East of the
312 Andes and Comparison to those in the South Atlantic Convergence Zone. *Journal of*
313 *Climate*, v. 17, n. 19, p. 3829-3842, 2004.

314 Marengo, J. A.; Soares, W. R.; Saulo, C.; Nicolini, M. Climatology of the Low-Level Jet
315 East of the Andes as Derived from the NCEP-NCAR Reanalyses: Characteristics and
316 Temporal Variability. *Journal of Climate*, v. 17, n. 12, p. 2261-2280, 2004.

317 Mendonça, F.; Danni-Oliveira, I. M. *Climatologia – Noções Básicas e Climas do*
318 *Brasil*. Oficina de Textos Publisher, 206 pp., 2007.

319 Oliveira, M. I.; Nascimento, E. L.; Kannenberg, C. A New Look at the Identification of
320 Low-Level Jets in South America. *Monthly Weather Review*, v. 146, n. 7, p. 2315-
321 2334, 2018.

322 Reboita, M. S.; Gan, M. A.; Rocha, R. P.; Ambrizzi, T. Precipitation Regimes in South
323 America: A Bibliography Review. *Brazilian Journal of Meteorology*, v. 25, n. 2, p. 185-
324 204, 2010.

325 Saha, S.; Moorthi, S.; Wu, X.; Wang, J.; *et al.* The NCEP Climate Forecast System
326 Version 2. *Journal of Climate*, v. 27, n. 6, p. 2185-2208, 2014.

327 Siqueira, J. R.; Machado, L. A. T. Influence of the Frontal Systems on the Day-to-Day
328 Convection Variability over South America. *Journal of Climate*, v. 17, n. 9, p. 1754-
329 1766, 2004.

330 Souza, E. B.; Ambrizzi, T. Pentad Precipitation Climatology over Brazil and the
331 Associated Atmospheric Mechanisms. *Climanálise Journal*, n. 1, 2004.

332 Torres, F. T. P.; Machado, P. J. O. *Introdução à Climatologia – Série Textos Básicos*
333 *de Geografia*. Cengage Learning Publisher, 280 pp., 2011.

334 Weykamp, F. V.; Ambrizzi, T. The Role of the Low-Level Jet East of the Andes in
335 Extreme Rainfall Events over Southern South America. *Proceedings of the*
336 *International Conference on Southern Hemisphere Meteorology and Oceanography*,
337 p. 1231-1234, 2006.

338 Zavattini, J. A.; Boim, M. N. *Climatologia Geográfica – Teoria e Prática de Pesquisa*.
339 *Alínia Publisher*, 152 pp., 2013.