

**Determination of hydro geological and Petroleum potential Using high resolution aeromagnetic data, over Matsina and Environs, Part of Nigerian sector of Chad basin, North Eastern Nigeria**

ABSTRACT

*Two dimensional spectral analyses of high resolution aeromagnetic data was carried out over matsina and environs North Eastern Nigeria to determine hydrocarbon and groundwater potentials. The analysis was done using Oasis Montaj version 8.2. The result obtained indicates the existence of two source depths. The shallower magnetic sources varies between 0.35 to 0.55 m and deeper magnetic sources also varies between 0.6 to 2000 m. The shallower magnetic sources could be as a result of basic intrusive within the Borno Basin, while the deeper sources correspond to the basement topography underlying the Chad formation. Comparing the result with what was obtained in Gubio, Maiduguri sub basin, it is apparent that the sedimentary thickness over Matsina and and Environs cannot be a potential site for oil exploration rather for ground water exploration.*

**Keywords:** Aeromagnetic Data, Chad formation, Sub-basin intrusives, Matsina and magnetic sources

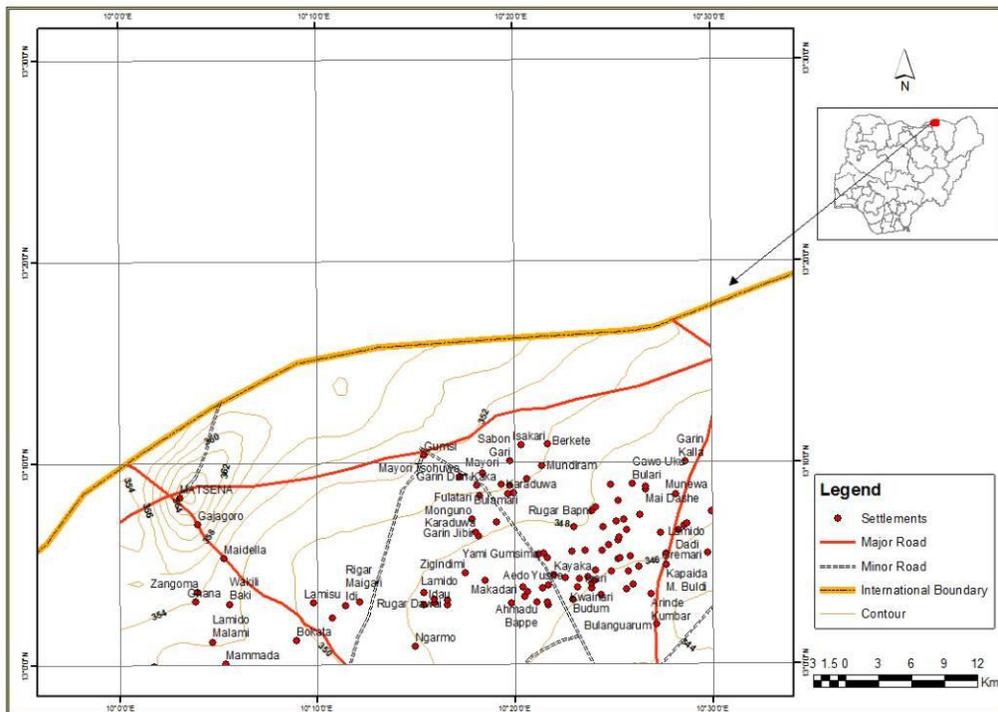
INTRODUCTION

The Nigerian sector of Chad Basin is a plain which slopes gently toward Lake Chad. It is devoid of rock outcrops except for Matsina area which have some exposures of younger granite. Chad Basin is covered with superficial deposits of sand stone and clay. It is the largest sedimentary Basin which covered the Northeastern part of Nigeria. The Chad Basin exhibit flat form of sedimentation in a geotectonic setting related to rifting (Okosun, 1995 ). The Nigerian sector of Chad basin is one-tenth total aerial extent of the Chad Basin which extends to Niger Republic, Chad and Cameroon. It is the part of West Central African Rift System (WCARS) that was formed in response to the Mechanical separation of the African crustal block in Cretaceous period.(Okosun 1995).

The aim of this study is to interpret high resolution aeromagnetic data over MatsIna and environs in order to determine petroleum potential of the study area using Spectral analysis and, to determine groundwater potential in the study area

LOCATION OF THE STUDY AREA

38 The study area ( Matsina) is located in the Northern Part of Yobe state which lies between latitude  $13^{\circ} 00^{\prime}$   
 39 and  $13^{\circ} 30^{\prime}$  N and Longitude  $10^{\circ} 00^{\prime}$  and  $10^{\circ} 30^{\prime}$  E. It consist of many farm land, major road linking  
 40 matsina to Niger republic, Kano and other parts of Yobe State (Fig. 1)



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 43 Fig. 1: GEOLOGY AND REGIONAL GEOLOGICAL SETTING

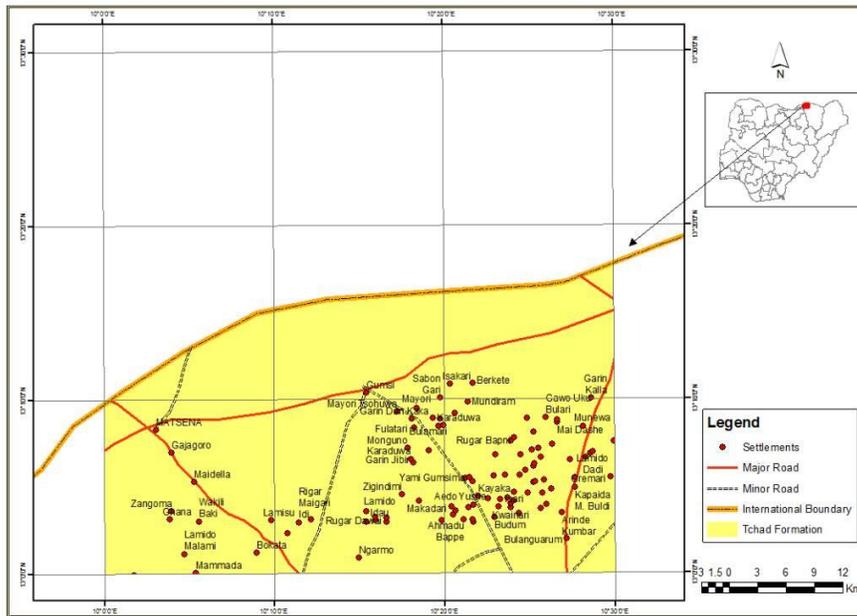
44  
 45 The geology of the study area is mainly chad formation with few outcrops of younger granite in  
 46 Matsina town even though on a small scale which is not captured on the geologic map of  
 47 Nigeria. The geology of the study area mainly Chad formation which consist of Sand and Clays  
 48 (Fig. 2)

49 The Chad Basin lies within the vast area of Central and West Africa at an elevation of  
 50 between 200 m and 500 m above sea level. The basin is centered around Lake Chad  
 51 and occupies an area of approximately 2,500, 000 km<sup>2</sup> extending over parts of the  
 52 Republic of Niger, Chad, Sudan and the northern portions of Cameroon and Nigeria.  
 53 The origin of the Chad Basin has been generally attributed to the rift system that  
 54 developed in the early Cretaceous when the African and South American lithospheric  
 55 plates separated and the Atlantic opened (Cratchley 1960). Pre Santonian Cretaceous  
 56 sediments were deposited within the rift system. The Nigerian sector of the Chad Basin

57 constitutes only about 6.5% of the entire basin and extends 152,000 km<sup>2</sup> of territory in  
58 Borno, Bauchi, Plateau and Kano States. The altitude of the basin ranges from 300 m  
59 within the lake to about 530 m at the western margin, along a distance of about 240 km.  
60 The Basin has developed at the intersection of many rifts, mainly in an extension of the  
61 Benue Trough. Major grabens then developed and sedimentation started. Sedimentary  
62 sequence span from the Paleozoic to Recent accompanied by a number of stratigraphic  
63 gaps. Sediments are mainly continental, sparsely fossiliferous, poorly sorted, and  
64 medium to coarse grained, feldspathic sandstones called the Bima Sandstone. A  
65 transitional calcareous deposit – Gongila Formation that accompanied the onset of  
66 marine incursions into the basin, overlies the Bima Sandstones.

67  
68 The Chad basin is one of the several basins within the West and Central African rift system and  
69 is genetically related to the Benue Trough (Burke, 1973). It consists of several sub-basins spread  
70 around the republics of Niger, Chad, Cameroon and Nigeria (Genik, 1992). Cratchley (1960), has  
71 delineated the Nigerian sector of Chad basin into three sub-basins prosperous for hydrocarbon  
72 exploration. These are centered around Gubio to the SW, Maiduguri to the south and Lake Chad  
73 to the North. The Bima formation is the basal unit in the six sequence found in the basin. The  
74 deposition of this sequence of sand stone, Mudstone and occasional shales of variable  
75 lithologies, textures, colours, and structures (Carter et al 1963) began in Aptian to Albian  
76 (Genik, 1992) unconformably on the basement complex. Up to 5km of this formation have been  
77 encountered in drilled holes, while interpreted seismic sections suggest 7 to 8 km depth extent  
78 (Avbovo et al 1986). Mudstone and shale horizons within shallower parts have total organic  
79 content (TOC) between 0.09 and 0.82 (Geo Engineering International, 1994). Sandstone  
80 horizons with average of 13.74% (Samaila, 2007) suggested to form possible reservoir, while  
81 the shale horizons form possible cap and seal rocks.

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85  
86 Fig. 2. Geological map of the study area  
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89 **MATERIALS AND METHOD**

90 In every research work, the materials selected and the methods to be used must be well understood and  
91 relevant as it affects the outcome. This research work was carried out in three main stages, data  
92 acquisition, data processing and data interpretation

93 **Data acquisition**

94 The aeromagnetic data used for this research work, were obtained as high resolution aeromagnetic data  
95 from Nigerian Geological Survey Agency (NGSA, 2010). The survey was carried at 0.05 seconds  
96 magnetic data recording interval, at 80 m terrain clearance; flight line spacing was 500 meters at 135  
97 degrees flight line trend. Tie line spacing was 500 meters at 225 degrees tie line trend. Cesium vapor 3X  
98 SCINTREX CS3 magnetometer was used for the survey. The data was acquired by Fugro air borne survey  
99 Limited as part of the nationwide geophysical data acquisition. This data was generally plotted using  
100 Universal Transverse Mercator (UTM) projection method WGS 1984.

101 **Data processing**

102 The data used for this research was processed using different filters in aeromagnetic data  
103 interpretation. These filters are commonly applied to remove spikes or noise and smoothed the  
104 grid data while preserving significant features. The goal is to select a filter and parameters that

105 do not introduce artifacts into the data. The magnetic data was further subjected to polynomial  
106 filtering to obtain both regional and residual maps. The residual map obtained was used as an  
107 input data to carry out spectral analysis in order to determine depths to magnetic sources in the  
108 study area (sedimentary thickness). All these processes are carried out in a Geo soft software  
109 Oasis Montaj version 8.2.

110 Several methods have been developed and improved in order to automatically estimate the depths  
111 to magnetic source from gridded magnetic data (Blakely and Simpson 1986; Grauch and Cordell, 1987;  
112 Roest *et al.*, 1993; Stefan and Vijay 1996; Thurston and Smith, 1997; Smith *et al.*, 1998). This method  
113 like the horizontal gradient magnitude (HGM), the 3D analytic signal (AS) or total gradient (TG) and the  
114 local wave number (LW) methods, all have a common approach to automatically determine the depth to  
115 magnetic sources. All the methods are based on the transform of the potential field anomalies into spatial  
116 functions that form gradient peaks and ridges over the sources. These maximum peak values are located  
117 directly above the magnetic source, depending on an assumed geometric model. Again the method can  
118 also use the same function to locate the contacts and estimate the source depths.

119

## 120 **Spectral analysis of aeromagnetic data**

### 121 **(i) Fourier Transformation**

122 It has become a familiar concept to interpret aeromagnetic data with one or two  
123 dimensional spectral analysis consisting of various frequencies which characterize the  
124 anomalies. The amplitude and phase relationship among these frequencies constitute what is  
125 known as a "complex line spectrum". The relationship has been used extensively by several  
126 authors (Spector and Grant, 1970; ; Hahn *et al.*, 1976; Negi *et al.*, 1983; Ofoegbu and Onuoha,  
127 1991, Nur *et al.*, 1994, 2000, Nur *et al.*, 1999, Kasidi and Nur, 2012a and Kasidi and Nur,  
128 2013), Nur *et al.*, 2011, interpreted total intensity magnetic map over Garkida and Environs  
129 statistically in terms of subsurface structures by using two-dimensional power spectral analysis.  
130 Very recently Kasidi and Nur (2012a, 2013), utilized spectral analysis of a simplified  
131 mathematical formula for the interpretation of magnetic data over the Mutum Biyu and  
132 environs, Jalingo and environs Northeastern Nigeria.

133 In this research, the Fourier transform technique was applied to the residual magnetic  
134 data. As the authors mentioned and pointed out, if a residual Magnetic anomaly map of

135 dimensions  $L \times L$ , is digitized at equal intervals, the values can be expressed in terms of double  
 136 Fourier series expansion.

$$137 \quad T(x,y) = \sum_{n=0}^N \sum_{m=-m}^M P_m^n \cos[(2\pi/L)(n_x+m-p)] + Q_m^n \sin[(2\pi/L)(n_x + m_y)] \quad 1$$

138 Where  $L$  = length of the square side,

139  $P_m^n$  and  $Q_m^n$  = Fourier amplitudes and

140  $N, M$  = number of grid points along the X, Y directions.

141 The sum

$$142 \quad P_m^n \cos [(2\pi/L)(n_x+m_y)] + Q_m^n \sin [(2\pi/L)(n_x + m_y)] \quad 2$$

143 Represents a single partial wave having a particular direction and wavelength for which

$$144 \quad (P_m^n)^2 + (Q_m^n)^2 = (C_m^n)^2; \quad 3$$

145  $C_m^n$  Is the amplitude of the partial wave, while the frequency of this wave is given

$$146 \quad f_m^n = (n^2 + m^2)^{1/2} \quad 4$$

147 If the logarithms of such an amplitude spectrum are plotted against the frequency (linear scale)  
 148 one finds series of points which may be represented by one or more straight lines. The line  
 149 segment in the higher frequency range is from the shallow sources and the lower harmonics are  
 150 indicative of sources from deep – seated magnetic bodies. Then the slope of the segment is  
 151 related to depths (Spector and Grant 1970)

152 The Fourier transform of magnetic data digitized in a square grid forms a square matrix  
 153 which can be reduced to a set of average amplitudes depending only on the frequency (Hahn *et*  
 154 *al.*, 1976). These average amplitudes fully represent a spectrum from, which the magnetic  
 155 sources were estimated.

156 The use of Discrete Fourier transformation involves some practical problems, such as the  
 157 problems of aliasing, truncation effect or Gibb's phenomenon and the problems associated with

158 the even and odd symmetries of the real and imaginary parts of the Fourier transformation  
159 (Ofoegbu and Onuoha, 1991).

160 The aliasing effect arises from the ambiguity in the frequency represented by the sampled  
161 data. Frequencies greater than the Nyquist frequency, which tends to impersonate the lower  
162 frequencies are known as the aliasing effect. To avoid or reduce the effect of aliasing,  
163 frequencies, greater than the Nyquist frequency must be removed through the use of an aliasing  
164 filter, which provides high attenuation above the Nyquist frequency. Aliasing can also be  
165 reduced through the use of small sampling intervals such that, the Nyquist frequency is equal to  
166 or greater than the highest frequency component present in the function being analyzed.

167 When a limited portion of an aeromagnetic anomaly map or short profile is subjected to  
168 Fourier analysis, it is difficult to reconstruct the sharp edges of the anomaly with a limited  
169 number of frequencies and this produces what is known as the Gibbs phenomenon. This Gibbs  
170 phenomenon or truncation effect is equivalent to the convolution of the Fourier transform of the  
171 function with that of a rectangular window which is a sine cardinal function. This convolution  
172 introduces ripples at the edges of the function, which manifests itself as spurious oscillations at  
173 the discontinuity. Increasing the length of the window makes the Fourier transform tend towards  
174 a delta function, with subsequent reduction of the ripples at the edges. The truncation effect can  
175 therefore be reduced by selecting a large portion of anomaly or a long profile centered on the  
176 feature of interest. An alternative and more effective approach to reducing the truncation effect is  
177 by the application of cosine taper to the observed data ( Ofoegbu and Onuoha 1991).

178 In this analysis, some practical problems, such as the problems of aliasing, truncation  
179 effect or Gibbs phenomenon and the problems associated with the even and odd symmetries of  
180 the real and imaginary parts of the Fourier transformation (Ofoegbu and Onuoha, 1991) has been  
181 incorporated into the computer program

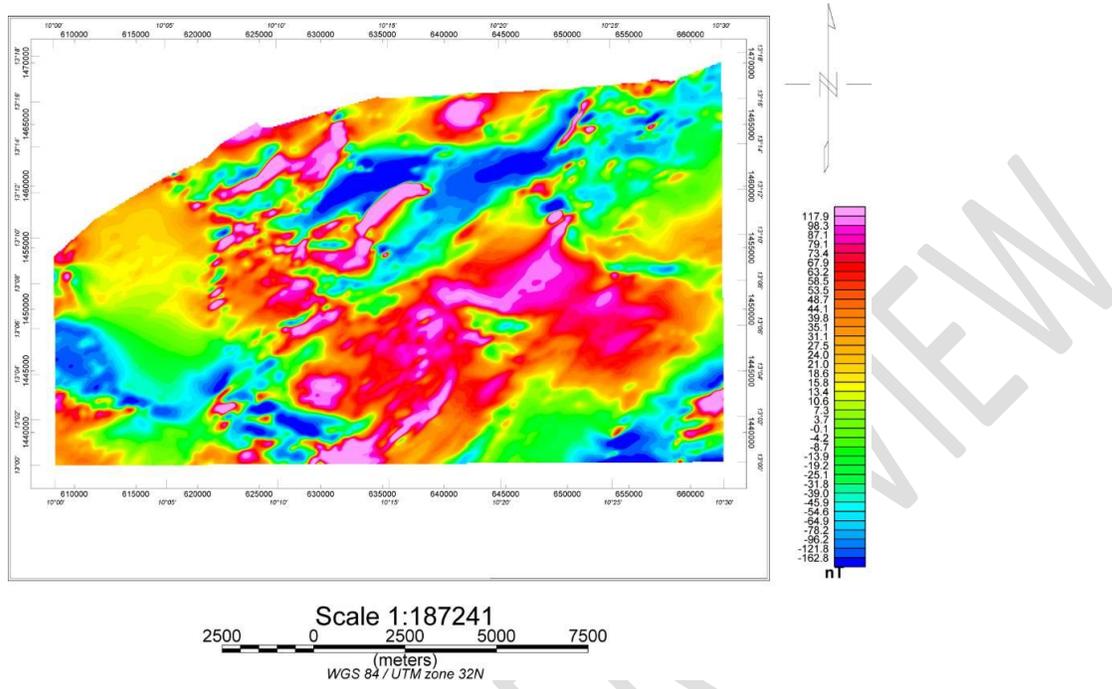
182 In order to carry out spectral analysis, data of the study area were divided into sixteen (16)  
183 blocks, even though the first four blocks have no data because it extend into Niger Republic. So the data  
184 begins in block five

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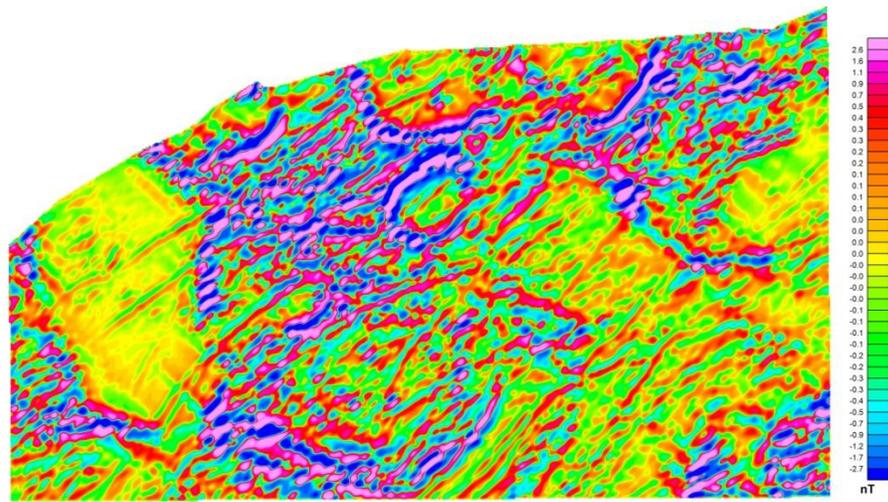
187 RESULTS

188 In this section, the results of the findings are presented inform of figures graphs and tables



189 Fig. 3.Total magnetic Intensity map (TMI) of the study area

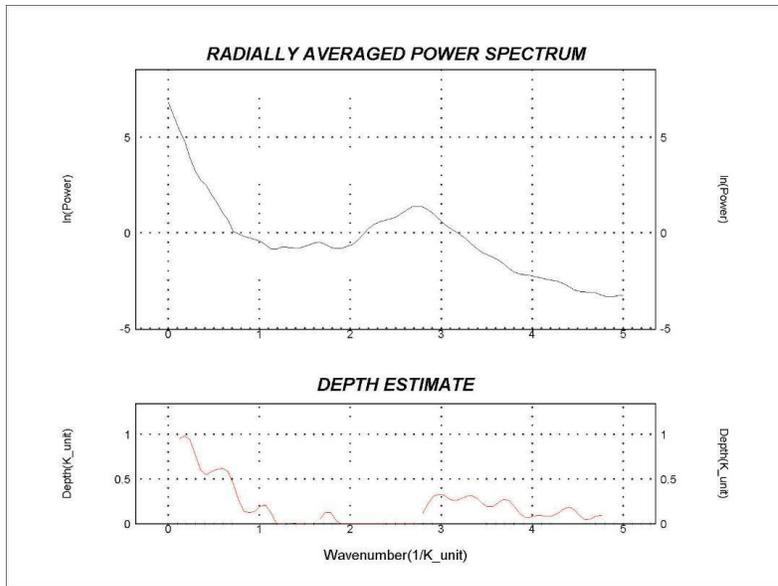
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193 Fig.4. Residual map of the study area

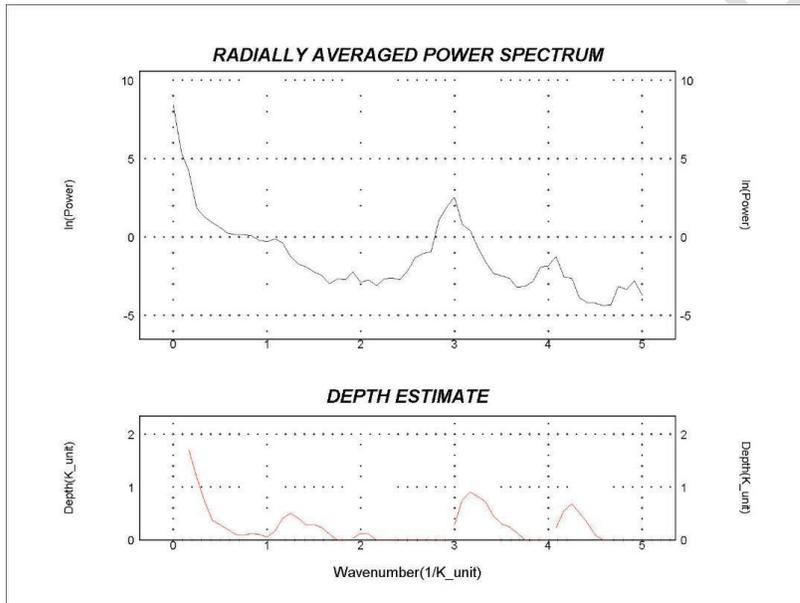
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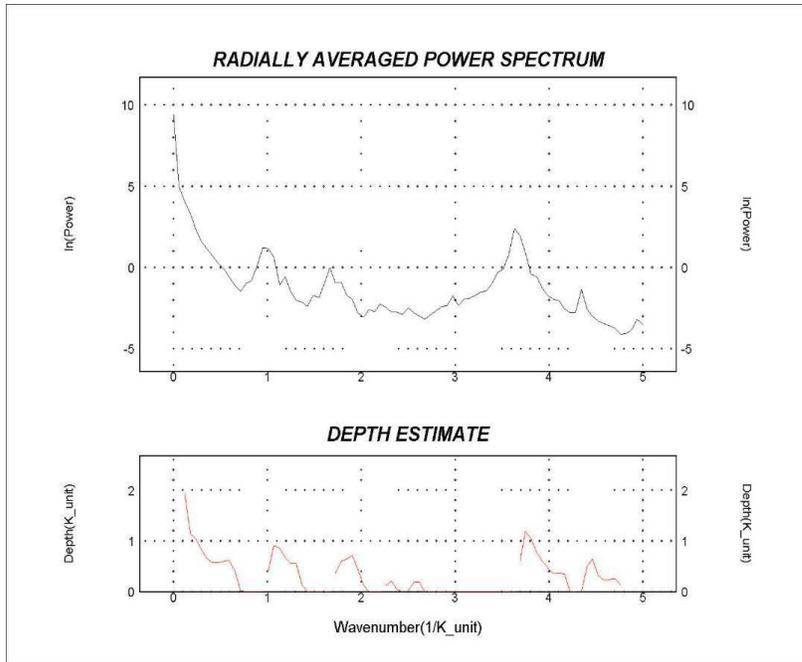


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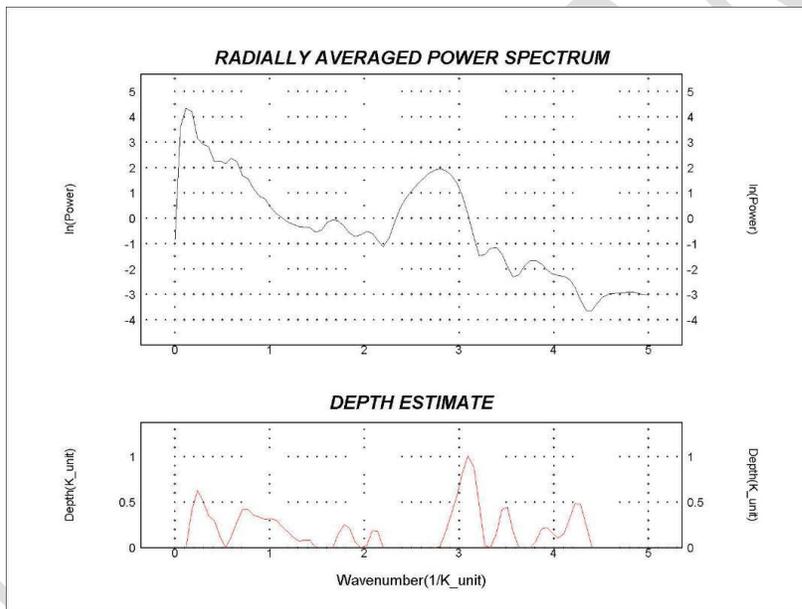


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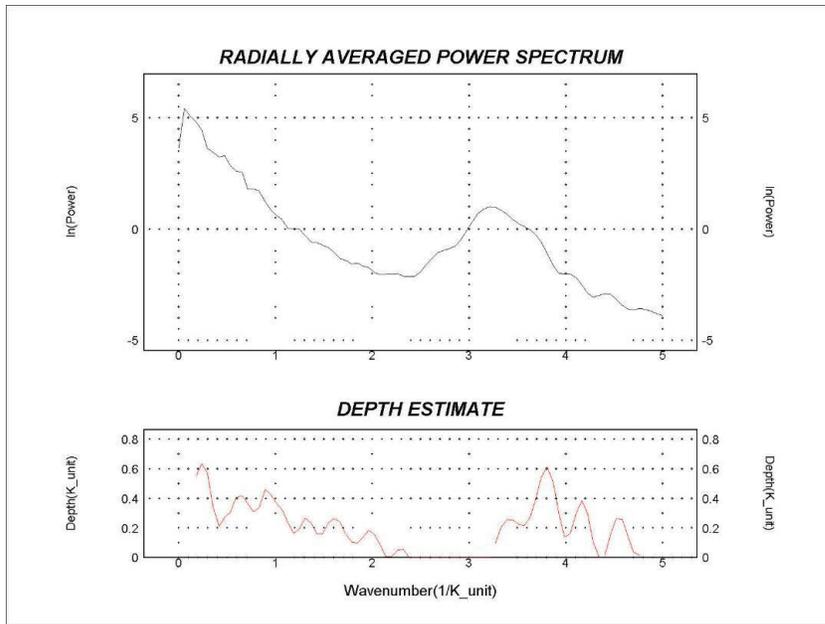
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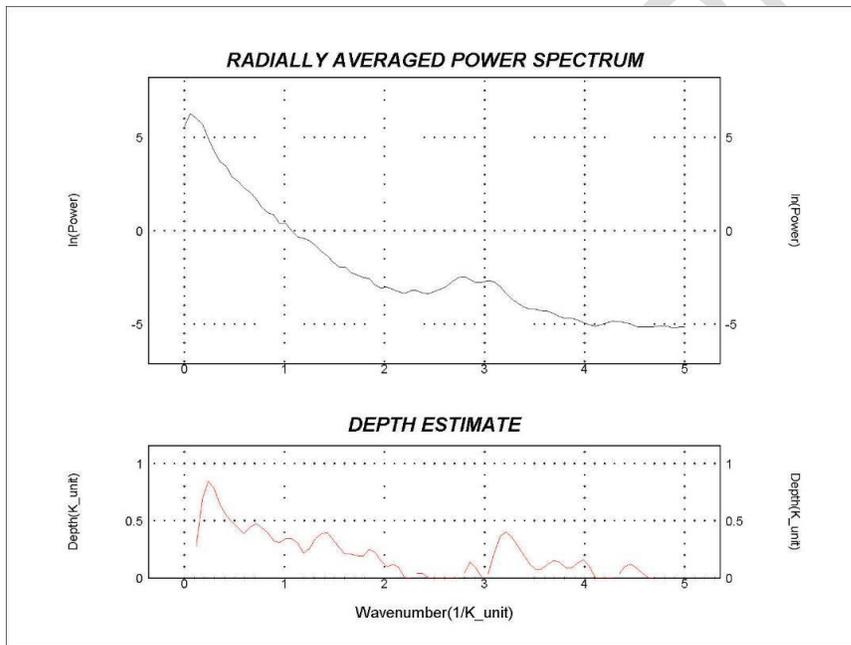
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Block 8



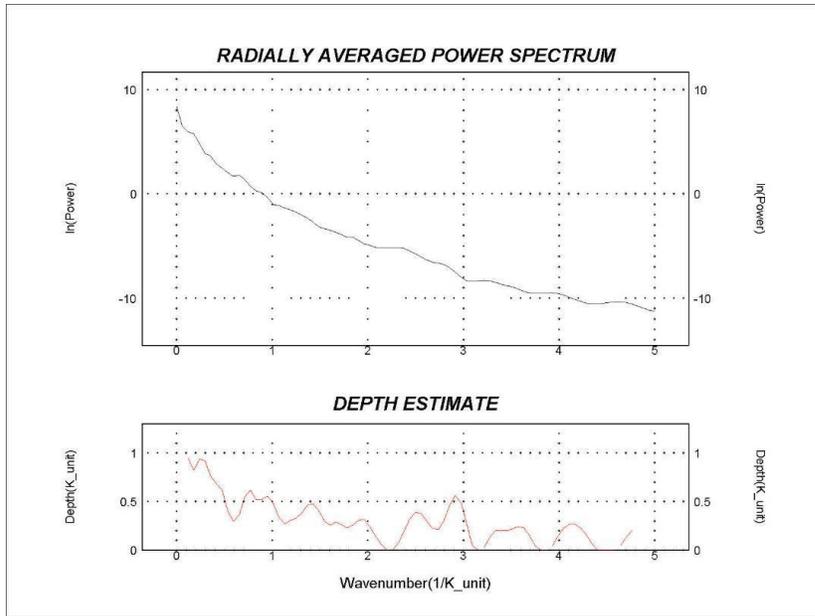
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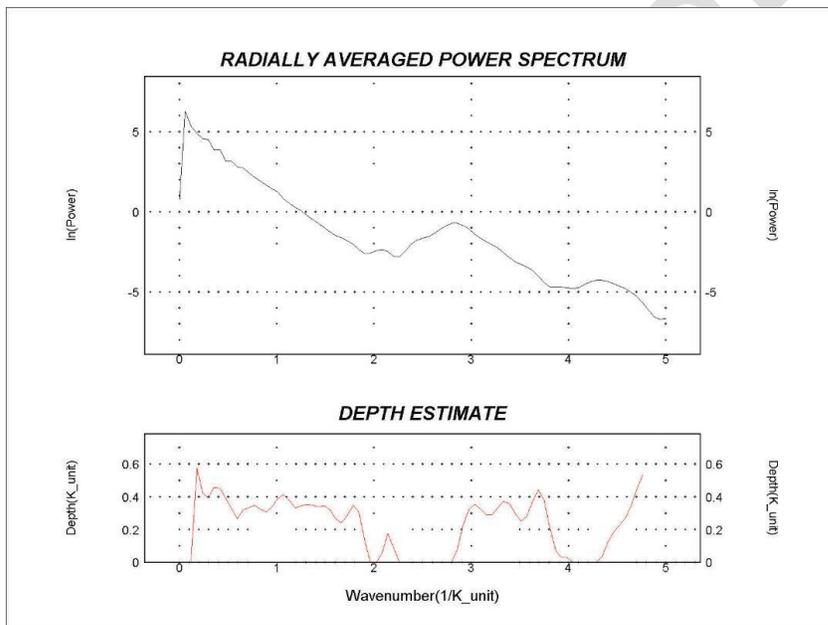
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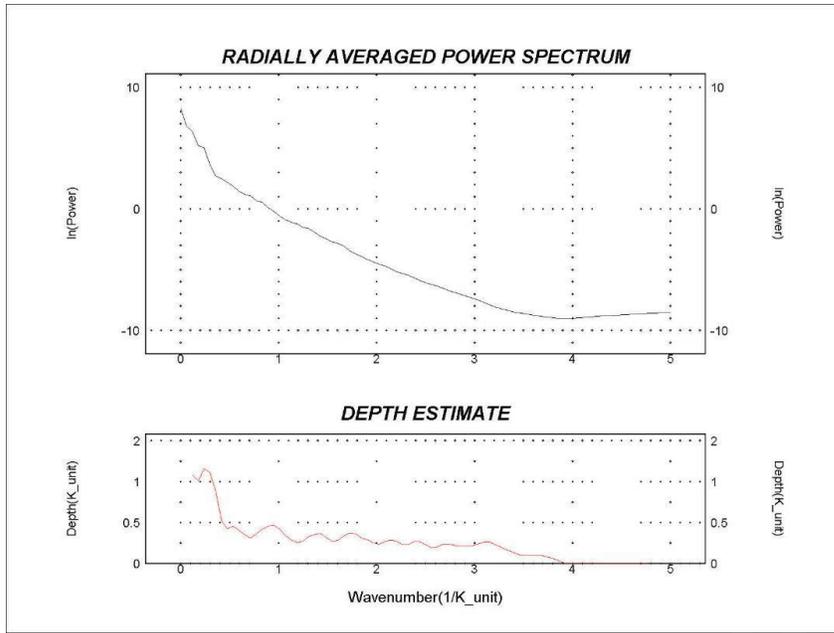
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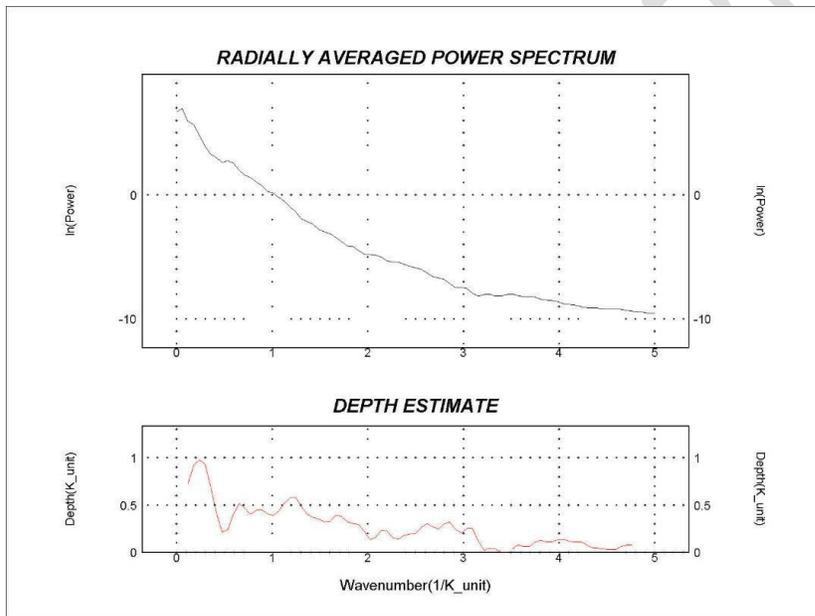
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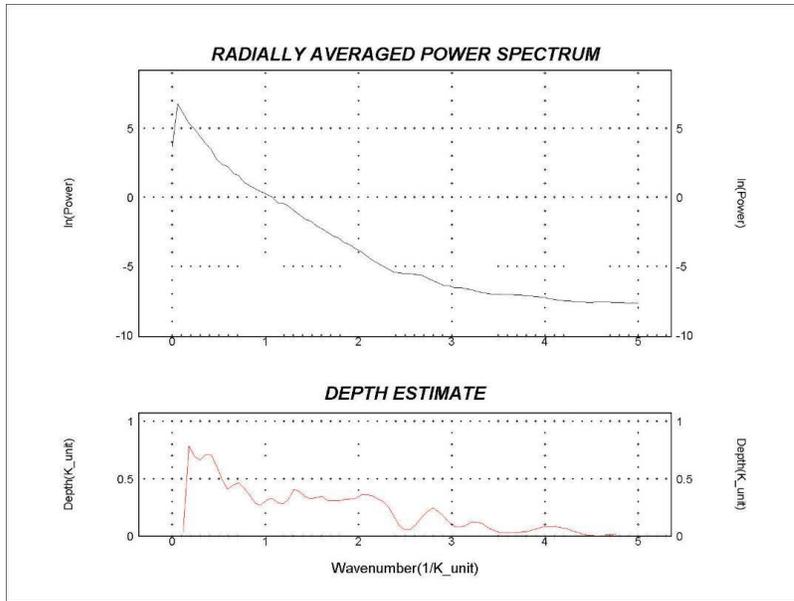
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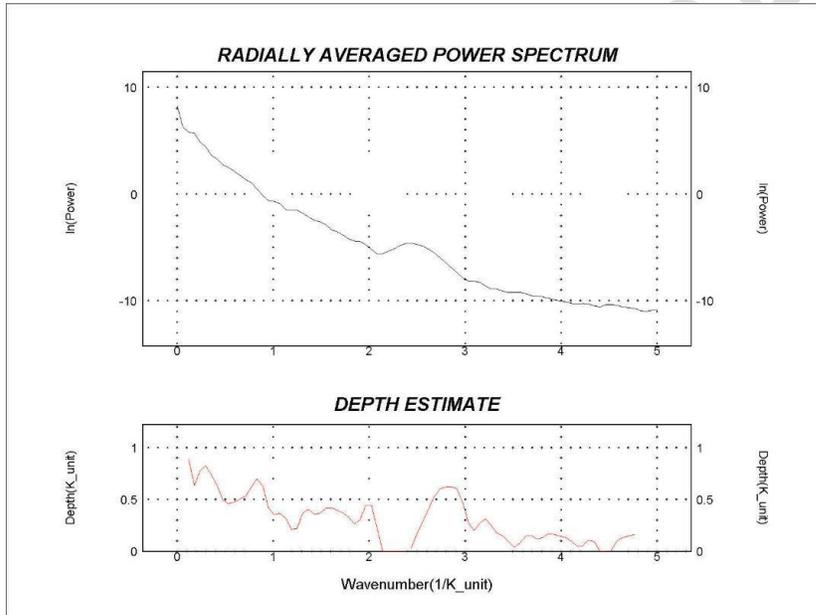
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Block 16

Fig. 5 Radial average Spectrum blocks ( 5-16)

238 Table 1. Average depth to Magnetic sources in Matsina (km)

239	<u>Blocks</u>	<u>Deeper sources (D1)</u>	<u>Shallower sources (D2)</u>
240	Block 5	1.0	0.6
241	Block 6	1.8	0.5
242	Block 7	2.0	0.9
243	Block 8	0.6	0.4
244	Block 9	0.6	0.4
245	Block 10	0.7	0.4
246	Block 11	0.9	0.8
247	Block 12	0.6	0.5
248	Block 13	1.4	0.4
249	Block 14	1.0	0.5
250	Block 15	0.8	0.7
251	<u>Block 16</u>	<u>0.9</u>	<u>0.7</u>

252

## 253 DISCUSSION OF RESULT

254 The total magnetic intensity map (Fig.3) of the study area was used as the input data in the  
 255 software which subjected to regional residual separation to produce residual map (Fig.4) of the  
 256 study area. The residual map shows magnetic anomalies with high and low intensity values  
 257 which ranges between -2.6 to 2.6 nT which spread across the study area as shown by colour  
 258 legend. Like the residual data was used as the input data to perform the spectral analysis in order  
 259 to achieve the aforementioned objectives.

260 The results for determination of depth to magnetic sources for Petroleum and  
 261 Groundwater exploration using spectral analysis of high resolution aeromagnetic data over  
 262 Matsina and Environs North Eastern Nigeria indicates two depth , the shallow sources and the  
 263 deeper sources. The shallow source depths are represented by D2 and deeper sources are  
 264 represented by D1. The shallow source depths range between 0.35 to 0.55 km, while the deeper  
 265 sources range between 0.6 to 2.0 km. These values were obtained from figure 5 i.e the radial  
 266 average spectrum of blocks 5-16, which was displayed on Table 1. The depth to magnetic sources  
 267 normally represent the thickness of sedimentary cover in an area and the sedimentary thickness  
 268 gives the idea for oil and water exploration.

269 The existence of two source depth in Matsina and Environs is as a results of intrusion of younger  
 270 granite given a thin sedimentary cover in the study area compared to what is obtained in other  
 271 parts of Borno Basin. The deeper sources in this area which range between 0.6 to 2.0 km may not

272 support exploration of Petroleum due to the thickness of sedimentary cover compared to areas  
273 like Gubio Basin , Gajigana, Magumeri, which are known to have thick sedimentary cover of  
274 6.0 km and above. This area was affected by upwelling of younger granite during the Pan  
275 African orogeny, the heat due to long time cooling of the magma after intrusion might affect the  
276 hydrocarbon accumulation in the area.

277 Therefore, based on the results, it is considered that all areas of deeper or thick sedimentary  
278 cover which ranges between 0.6 to 2.0 km are potential site for groundwater exploration in the  
279 study area. By implication the study area may no be suitable for hydrocarbon exploration due  
280 thin sedimentary cover and up arching of younger granite.

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