

**ANALYSIS OF QUATERNARY SEDIMENTS OF QUARTZ GRAINS APPLIED TO THE IDENTIFICATION OF THE ENVIRONMENT OF SOME IVORY BEACHES EAST OF ABIDJAN (CÔTE D'IVOIRE)**

**ABSTRACT**

Morphoscopic and exoscopic analysis of quartz grains collected on the ivorian beaches of the gulf of Guinea between Abidjan (Côte d'Ivoire) and Aforenou (Ghana) during topographic surveys between March 2007 and March 2009 on the examination of quartz at the binocular loupe reveals blunting grains shining on all beaches (50% to 70%).

In the Abidjan zone, the majority of the grains are reddish, indicates a ferrous and inherited environment. The exoscopy for the fine and detailed study of the morphology and the surface of the quartz grains was made with the Scanning Electron Microscope (SEM). This approach makes it possible to interpret traces related to events (energy level) or environments (transport, transition and storage environments).

These observations indicate that the quartz after a long transport in a fluvial environment have been reworked in a marine environment. They were finally deposited in a low-energy aquatic continental medium marked by the polishing of the crystalline points and the siliceous corpuscles scarttering the surface of the grains.

The energy of the transport media crossed is high in a turbulent environment through shock marks, as large as many. SEM examination of the samples revealed that beach quartz generally underwent fluvial transport and resumed in an intertidal and / or subtidal environment. They have a continental and marine origin.

*Keywords: morphoscopy; exoscopy; quartz; transport; Côte d'Ivoire*

**1. INTRODUCTION**

The succession of marks of different physico-chemical environments and paleogeographic evolution, sediments are to know for a better understanding of coastal dynamic. The work of several authors [1, 2, 3, 4, 5, 6] have highlighted the importance of the study of sediments, in particular that of quartz grains in shoreline evolution, and the shingle studies of [7] and [8].

In Côte d'Ivoire the work of [9, 10, 11, 12, 13, 14, 15, 16, 17, 18] provided a better understanding of the Quaternary and current deposits of the ivorian coast, as well as the fluvial and marine formation.

When a rock is subject to weathering, it releases a fraction of its constituent minerals in the form of grains. These grains are sometimes trapped and subject to pedogenic influences and in some cases may be released again and carried further [19, 20, 21].

Non-weatherable materials such as quartz undergo progressively during their evolutions more or less significant changes in size, shape and appearance compared to the original characters they had in the parent rock from which they come [22].

The grains found in the sedimentary formations bear on their surface the testimony of their history. It is a multitude of traces inherited from the various influences to which they were subjected during their transport.

47 These traces constitute a real archive which gives information on their conditions of transport,  
48 evolution and deposit [5, 6, 23] and that the examination. The sedimentology of the beach is  
49 one of the first essential steps in any impact study in the coastal environment [7, 8].

50 On this basis, a study of the quartz elements of several sedimentary formations of the Gulf of  
51 Guinea coast from Abidjan to Ghana by exoscopic analysis to characterize the hydrological  
52 environment of deposit of these formations has been realized.

53 The objective of this article is to highlight, through binocular magnifying glass and Scanning  
54 Electron Microscope (SEM) examination of the surface of quartz grains, the sediment  
55 environment that accounts for the mode of transport, of the crossed medium and deposit of the  
56 materials in order to advance in the knowledge of their origin.

57 This study is part of a research project on coastal erosion initiated by the Oceanological  
58 Research Center in collaboration with the laboratory of Marine Geology of the University  
59 Felix Houphouët Boigny will complement existing morphological databases on the coastline  
60 of the West African subregion.

61

62

63

## 64 **2. PRESENTATION OF THE STUDY AREA**

65

66 The study area concerns the segment of coast between Abidjan and Aforenou (border with  
67 Ghana) (Fig.1). Corresponding to the southern zone, situated between 0 and 200 m of altitude  
68 [24], the ivoirian sedimentary basin, formed of coarse sands, medium to very fine, extends  
69 along the atlantic coast. Covering an area of approximately 8000 km<sup>2</sup>, this sedimentary basin  
70 covers approximately 2,5% of the ivoirian territory, doesn't exceed 35 km wide and extends  
71 from Fresco to Axim in Ghana. The coastline of the Abidjan coastline differs from that of the  
72 entire Côte d'Ivoire coast by its morphology.

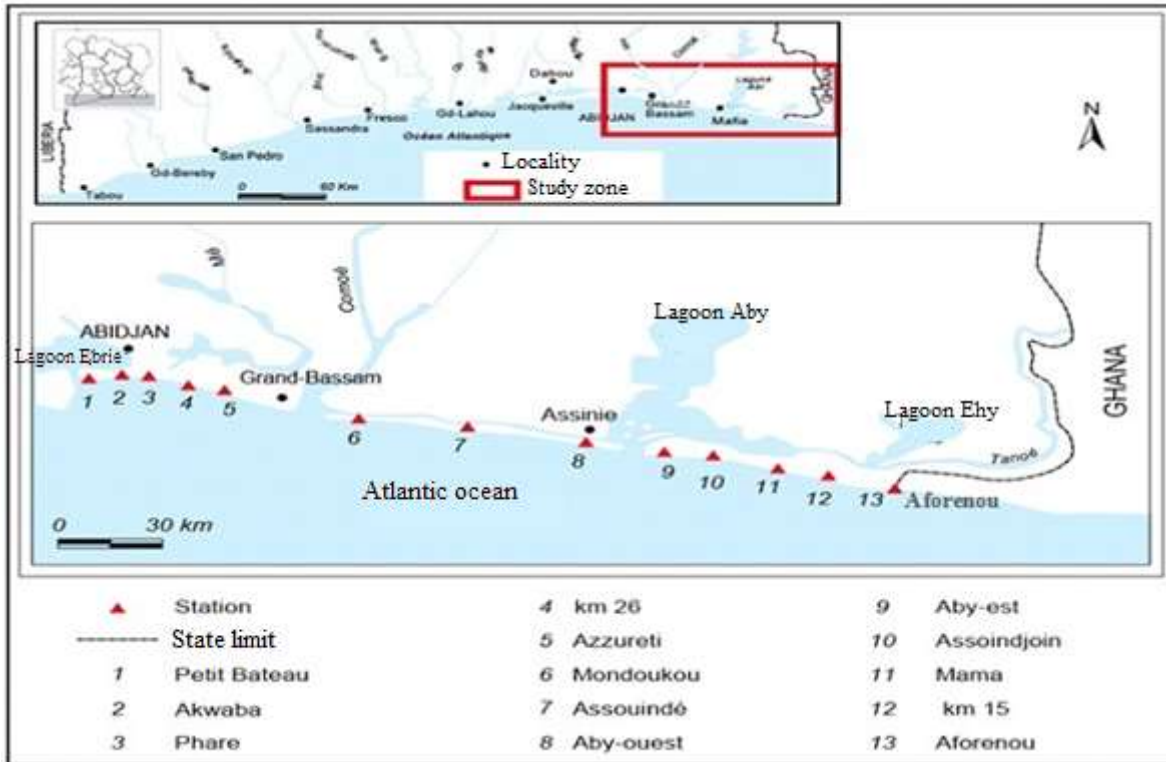
73 Depending on its orientation and its lithological nature, the submarine beach at Abidjan is  
74 influenced by an underwater canyon called "Trou-Sans-Fond" [25]. The coastline draws a bay  
75 before resuming the overall layout rather rectilinear. This configuration follows the outline of  
76 the head of the canyon of "Trou-Sans-Fond".

77

78

79

80



81

82

Fig. 1. Study area and presentation of sampling point

83

### 84 3. MATERIALS AND METHODS

85

86 To carry out this study, several samples were taken (stations 1 to 13) including those of the  
 87 light stations of Port-Bouët in Abidjan (3), Grand-Bassam (6), Assinie (8 and 9 located on the  
 88 and others of the mouth of the Aby lagoon) and Aforenou (13, border with Ghana) were used.  
 89 They were taken manually in the superficial layer of beach profiles on the upper foreshore  
 90 (HE), the mid-foreshore (ME) and the lower foreshore (BE).

91

92

#### 93 3.1 Morphoscopic study of quartz grains

94

95 The analysis of the quartz grains was based on the examination of the 500  $\mu\text{m}$  to 630  $\mu\text{m}$   
 96 fraction [26] on the shape and appearance (wear) of sand grains in order to identify their mode  
 97 of transport.

98 For the morphoscopy of quartz grains, the selection of quartz minerals was carried out using a  
 99 binocular magnifier connected to a screen. To do this the method of [1] was used.

100 The particle size fraction between 80  $\mu\text{m}$  and 100  $\mu\text{m}$  has been neglected because the shaping  
 101 is generally poorly discernible and also requires high magnification [27].

102 The morphoscopic analysis is based on the classification of grains established by [1]. It takes  
 103 into account the shape of the grains and the appearance of their surface. This classification  
 104 thus leads to three main categories of grains:

105 The "No Worn" or "NW", the "Blunt shiny" or "BS" and the "Round-Mast" or "RM". To  
 106 evaluate the percentage of the classes of the grain shape, the visual chart (Diagram of  
 107 evaluation of the relative percentages of the grains) established by [28] was used starting from  
 108 the surface of the thin blades made up.

109

### 3.2 Exoscopic study of quartz grains

The determination of the beach environment in this study is based on the observation of the surface of the quartz grains using the Scanning Electron Microscope (SEM). The sedimentological analysis was done beforehand according to the classical methodology described by [29] and according to the work of [30] on sediment samples taken.

The techniques of [1, 2, 12, 23] allowed the exoscopic study of quartz grains. The study is based on the identification and interpretation of physical, chemical, mechanical and biological factors, which leave on the surface of quartz grains micro-characters of shape and size characteristic of the factors that generated them. These basic characters make it possible not only to accurately determine the deposit medium of a grain of sand, but also to trace its entire geological history, its environment traversed during transport and in some cases, its geographical origin [23].

## 4. RESULTS

### 4.1 Shape and appearance of quartz grains

The morphoscopic analysis of the quartz shows the omnipresence of the blunting blunting grains (Fig.2), which highlights the influence of water transport and a dynamic fluvio-marine process. Matt matte and blunted round grains are found throughout the study area.

The percentage of blunting grains blunting on the coastal zone is as follows: 50% on the beaches of Abidjan-Grand-Bassam, 60% in Assouindé and 70% in Assinie-Aforenou. These quartz grains show shock traces visible by scanning electron microscope (SEM) in the form of cup

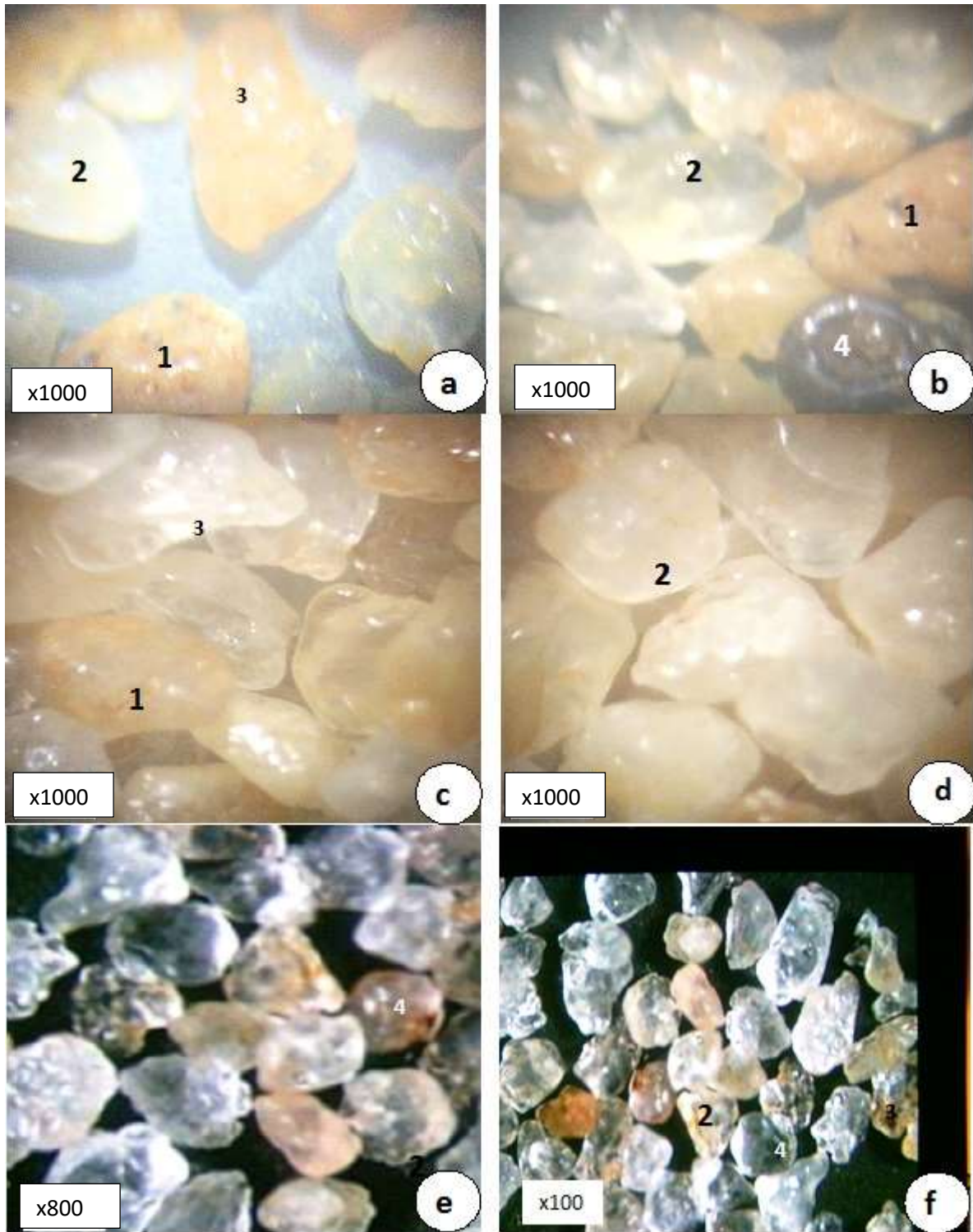
These quartz grains are, testifying to a wind recovery. The presence of a ferruginous coating on a few grains only in Abidjan bay indicates a ferrous environment. Unused grains are present in a small percentage (<2%) of up to 5% immediately east of the inlet channel in the port of Abidjan. The presence of these grains would reflect a nearby source of supply, with sediments little changed.

From Abidjan to Aforenou, there is a tendency for this percentage of unused grains to decrease, although it is not very clear in Port Bouët bay (Abidjan), and at the same time there is an increase in blunted grains. The small proportions of the sub-angular, matte grains appear only in Port Bouët bay, between the Vridi Canal and the Port Bouët lighthouse, just at the entrance to the entrance channel in the port of Abidjan, in upstream of the bay.

**Table 1. Evolution of the morphology of the grains from the morphoscopy of these on the various studied ranges**

	Petit Bateau	Phare de Port Bouët	Km 26	Mondoukou	Assinie France	Aforenou
NW	45%	40%	30%	30%	25%	15%
BS	50%	55%	60%	60%	50%	55%
RM	5%	5%	10%	10%	25%	30%

There is an increase in blunted shiny Quartz (BS) grains as they are transported by coastal drift to the east.



1 = ferrous; 2 = blunted blunting and subarranged; 3 = subangular; 4 = Rounded  
 a-Petit Bateau; b-Phare de Port Bouët ; c-Km 26; d-Mondoukou; e-Assinie France; f-Aforenou ( data [15])

**Fig. 2. Morphoscopy of quartz grains in some areas of study**

156  
 157  
 158  
 159  
 160  
 161  
 162  
 163  
 164  
 165

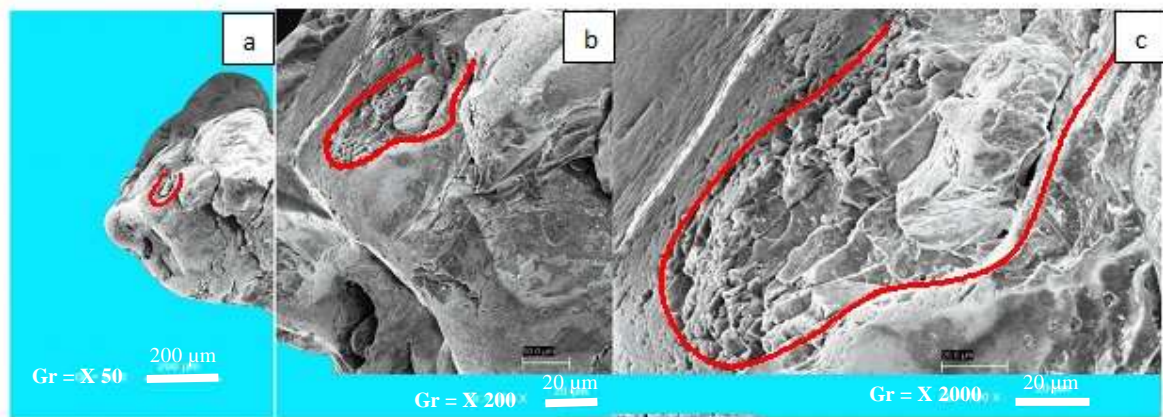


166  
167  
168  
169  
170  
171  
172  
173  
174  
175  
176  
177  
178

## 4.2 Quartz grain environment of Abidjan beaches

### 4.2.1 Appearance of some quartz grains from the beaches of Abidjan

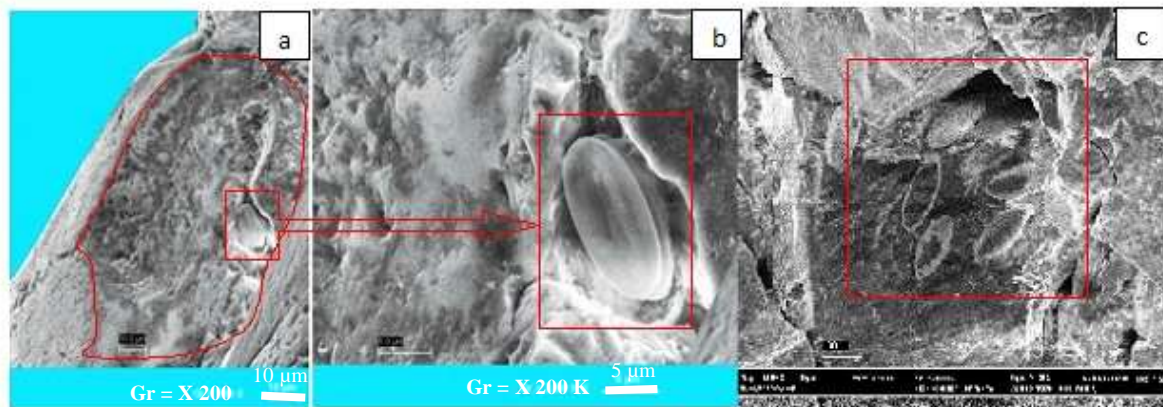
At low magnification of 50 and 200 times (Figs 3a and 3b), shock marks on quartz are numerous and visible. Some polished traces are exploited by dissolving figures on the parts most exposed to mixing (Fig. 3c). The edges are dull and the grains are sub-rounded to rounded. They appear very clearly in Fig. 3c with a white color, milky to the observation. They are distinguished by the presence of conchoidal breaks. Very small sizes smaller than 5  $\mu\text{m}$ , these inclusions are difficult to detect.



179  
180  
181  
182  
183  
184  
185  
186  
187  
188  
189

**Fig. 3. Traces of shock exploited by figures of dissolution on quartz grains of the beaches of Abidjan (data [15])**

Fig. 4a, 4b and 4c show on the surface of the quartz shock-related dips or dissolution in which live diatoms whose size does not exceed 20 microns. At a magnification of 2500 times, appears the complete form of a diatom, unicelleulaire alga which is one of the constituents of the plankton. It is housed at the bottom of a cavity (Fig.4b). On some quartz during corrosion, dissolving diatoms are visible (Fig.4c).



190  
191  
192

**Fig. 4. Diatom housed in the shock traces of some quartz grains (data [15])**

193

## 194 4.2.2 Transport environment, evolution and deposition quartz of the beaches of Abidjan

195

196 The numerous shock marks observed on the quartz grains of this segment of coast reflect a  
197 turbulent environment or strong swell in which they have stayed. They also reflect a high  
198 energy level (Fig.3a and 3b). The rounded or sub-rounded shape of the ridges and the blunted  
199 appearance of the quartz grains (Fig.3c) are characteristic marks of sediments that have  
200 evolved in a coastal marine environment, where the comings and goings between the  
201 intertidal zone and the coastal dunes are very frequent.

202 The progressive rounding of their most prominent ridges indicates a mode of transport where  
203 these grains rubbed against each other. The phyto-plankton (diatoms) observed frequently in  
204 the shock marks on these beaches are proof of a low energy environment and calm deposition.  
205 Indeed, during the low seas, the quartz grains are exposed to the open air. However, their  
206 surface is generally not smooth. The quartz that hosts these phytoplankton, translate  
207 sediments that have passed through an environment, calm low energy. The different types of  
208 trace observed on the grains of this coastal area reflect the different types of energy (high and  
209 low) that shape it.

210

211

## 212 4.3 Quartz grain environment of Grand-Bassam beach

213

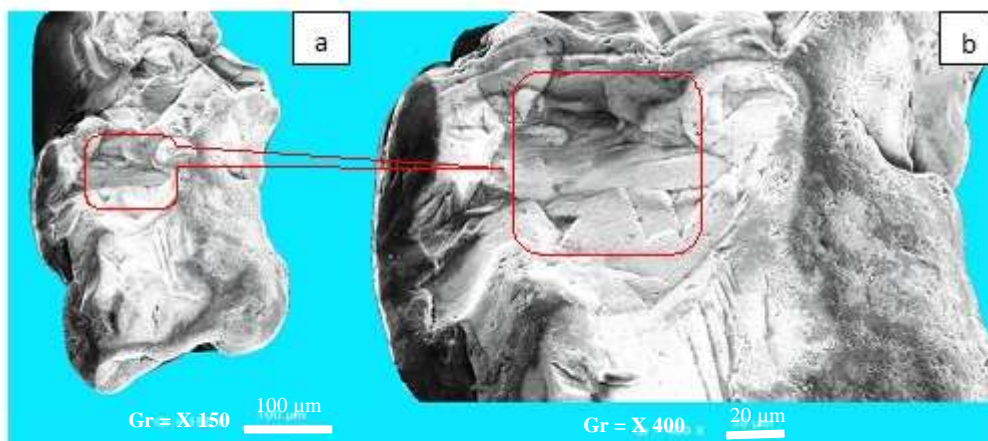
### 214 4.3.1 Appearance of some quartz grains of the beaches of Grand-Bassam

215

216 The quartz grains on this littoral space, generally very rounded, bear numerous traces of  
217 crescent shock, visible even at low magnification in Fig. 5a. These traces sometimes arranged  
218 in steps, appear at a high magnification in Fig. 5b. It is a remodeled quartz whose blunt edges  
219 consist of a group of pyramidal tops with blunt ends (Fig. 5a).

220

221



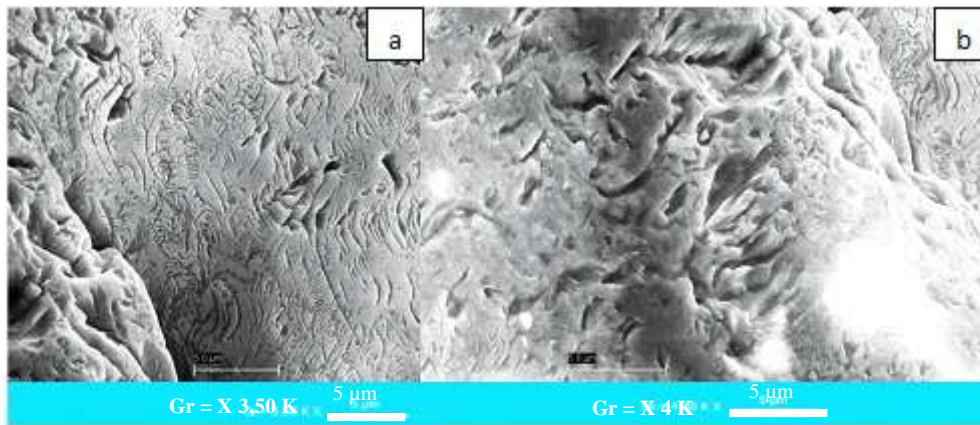
222

223

224 **Fig. 5. Traces of shock and some internal figures on a grain of quartz of Grand-Bassam**  
225 **(data [15])**

226

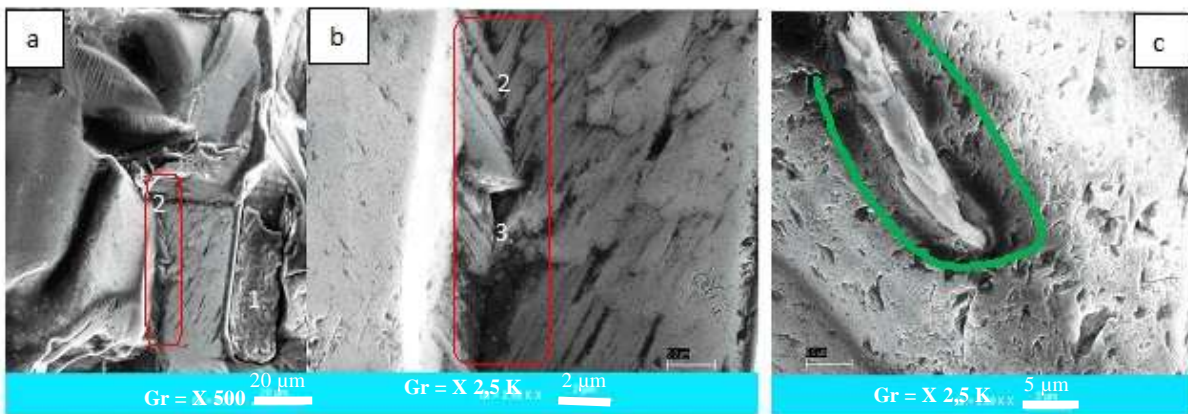
227 The study of the surface of certain quartz grains on this beach shows corroded and flaky  
228 aspects. These imprints, visible in Fig. 6a and 6b, are the marks of an intense chemical surface  
229 dissolution.



230  
231 **Fig. 6. Marks of superficial chemical dissolution on a Grand-Bassam quartz grain (data**  
232 **[15])**

233  
234 The flat surfaces are covered with amorphous silica film, and the general morphology of the  
235 cavities is similar to that of the inclusions (Fig. 7a and 7c). These magmatic vitreous  
236 inclusions totally crystallized in Fig. 7a (1), but not identifiable, coexist with structures in "V"  
237 clearly visible in Fig. 7b (2) (magnification x2500).

238



239  
240 (a) magmatic, (b) glassy, (c) cavities

241 **Fig. 7. Inclusions on the surface of certain quartz grains of the beaches of Grand Bassam**  
242 **(data [15])**

243

244

#### 245 4.3.2 Transport environment, evolution and deposition of quartz of the beaches of 246 Grand-Bassam

247


248 The very rounded shape in general of quartz grains on this littoral area, with a very blunted  
249 appearance, is the imprint of a quartz remodeled in the marine environment (Fig.5a).

250 Some quartz grains of corroded and desquamated appearance (Fig.6a and 6b) are traces of  
251 intense chemical surface dissolution. The numerous structures (in "V", inclusions, vitreous)  
252 that coexist on the surfaces of numerous grains show that this result reflects a low-energy  
253 fluvial or continental environment in which the mineral has been during its transport. The  
254 quartz grains in this littoral zone have, for the most part, had to cross several environments.  
255 They were taken up in an intertidal (aquatic) medium saturated with silica ( river).

256

257



258 **4.4 Environnement des grains de quartz de la plage à l'ouest de l'embouchure de la**  
259 **lagune Aby (Assinie)** 

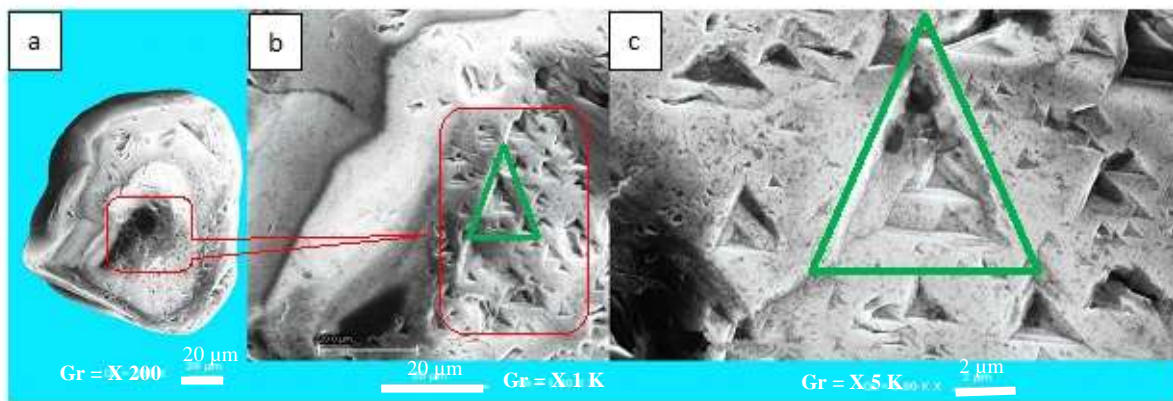
260

261 **4.4.1 Aspect des grains de quartz des plages à l'ouest de l'embouchure de la lagune Aby**  
262 **(Assinie)**

263

265 In this beach, some quartz have smooth faces and faces finely corroded by dissolution figures  
266 visible in Fig. 8a. At a magnification of 1000 times (b) and 5000 times (c), these figures in  
267 triangular form with stair steps have frequent angles of 30 °, 60 ° in the arrangement of the  
268 crests. Parallel lines also confirm the stair arrangement. The preceding figures show shock  
269 traces exploited by the dissolution, on which appear the triangular truncated features of the  
270 latter (Fig. 8b and 8c).

270



271

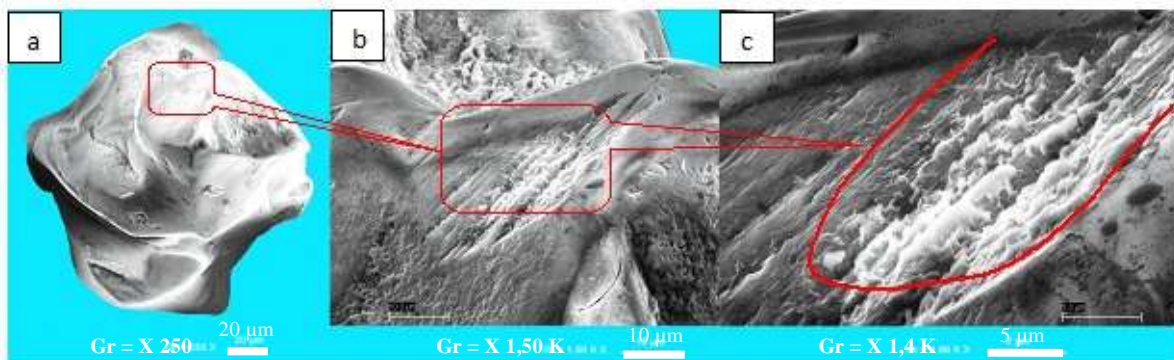
272 **Fig. 8. Traces of dissolution and triangular figures on the beaches west of the mouth of**  
273 **the Aby lagoon in Assinie (data [15])**

274

275

276 The loss of silica by dissolution of the surfaces of some quartz is not negligible although  
277 difficult to evaluate (Fig.9a, x250). At medium magnification (x 1500), the microrelief is  
278 granular, marked with furrows where there is no sharp peaks (Fig.9b). It is a low-energy  
279 aquatic continental recovery: This recovery, which corresponds to the final sedimentation  
280 medium of the sample, results in the precipitation of numerous siliceous globules over the  
281 entire surface of the quartz, including on the top edges (Fig.9c).

282

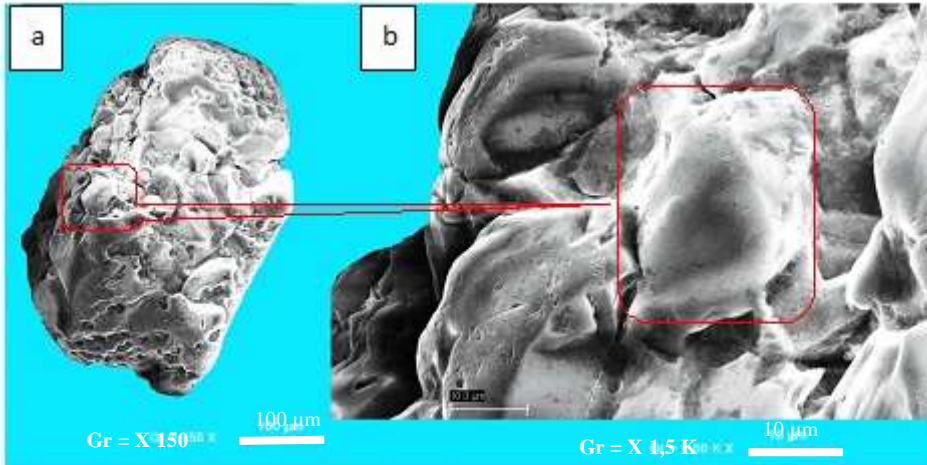


283

284

285 **Fig. 9. Chemical dissolution of the surface of a quartz grain from the beaches west of the**  
286 **mouth of the Aby Lagoon in Assinie (data [15])** 

287 At a magnification of 150 times (Fig. 10a), the surface interspersed with microcracks appears  
 288 to be particles in the course of disintegration, without definite orientation. But the crystalline  
 289 points themselves were in turn dulled by the final fluvial upturn, which scattered all the  
 290 surface of the grains with siliceous globules. This 1500-fold enlarged quartz grain exhibits  
 291 discontinuous microcracks, 1 to 2  $\mu\text{m}$  wide (Fig.10b). They have sinuous and sometimes  
 292 branched outlines. Quartz irregular microrelief, are bumpy, sometimes cavernous, micro-  
 293 fissured; but all the irregularities are blunted.  
 294



295  
 296  
 297  
 298  
 299

**Fig. 10. Dissolution of the entire surface of the quartz grain generating rounded micro-reliefs to the west of the mouth of the Aby lagoon in Assinie (data [15])**

300 **4.4.2 Transport environment, evolution and deposition of quartz beaches to the west of**  
 301 **the mouth of the lagoon Aby (Assinie)**

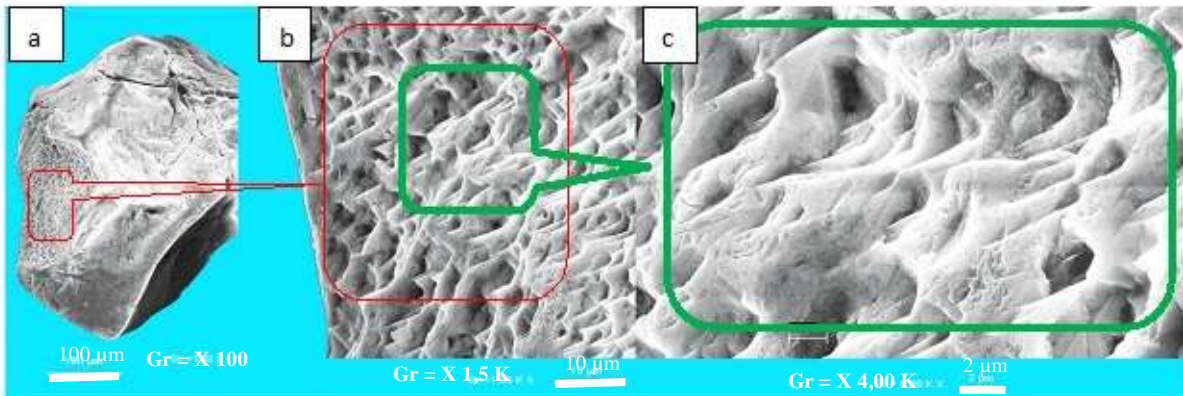
302  
 303 The dissolution phenomena observed on the shock marks of quartz near the mouth are  
 304 frequent and numerous. The hollow surfaces of the grains created by these dissolutions are  
 305 often filled by precipitations of many crystalline siliceous globules (Fig. 9c), with very  
 306 rounded edges and many globules all blunted (Fig.10). But the crystalline points themselves  
 307 were blunted by the final fluvial recovery. These quartz grains were taken up in a calm  
 308 aquatic continental medium of low energy. This recovery, which corresponds to the final  
 309 sedimentation medium of the sample, results in the precipitation of numerous siliceous  
 310 globules over the entire surface of the quartz, including on the vertex of the ridges (Fig.9c).  
 311

312 **4.5 Environment of the quartz grains of the beaches east of the mouth of the Aby lagoon**  
 313 **in Assinie**

314  
 315 **4.5.1 Appearance of quartz grains from the beaches east of the mouth of Aby lagoon in**  
 316 **Assinie**

317  
 318 Quartz have rounded and blunt edges, the faces remain smooth. These edges are affected, as  
 319 in rivers, by polishing gradient shock traces, but another phenomenon appears. Seawater  
 320 gnaws away at shock marks, which are exploited by dissolution figures, especially on the  
 321 ridges. At low magnification (x 100), the microrelief of a healthy quartz flake appears  
 322 irregular with sharp edges, with no apparent preferential orientation (Fig.11a). The face of this  
 323 mineral at high magnification (x 1500) in Fig. 11b shows a pattern oriented along bundles of  
 324 lines or ridges spaced a few microns, arranged in stairs. They are connected in relay on the

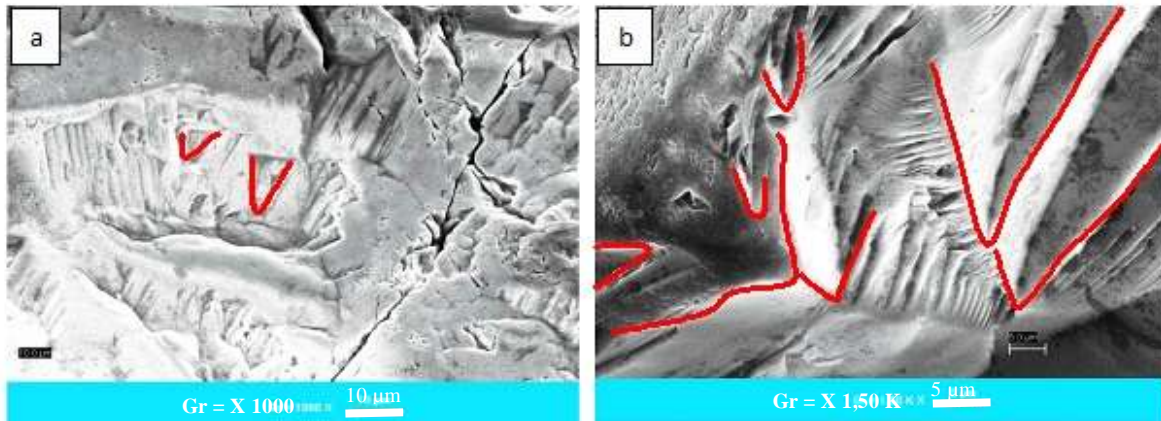
325 main peaks, delimiting somehow the stack of micro-scales detached after the shock. At an  
 326 enlargement of 4000 times in Fig. 11c, there appear fine parallel, tighter streaks spaced 1 to 2  
 327  $\mu\text{m}$  in a direction perpendicular or oblique with respect to the preceding peaks.  
 328



329  
 330

331 **Fig. 11. Chemical dissolution appearing on the tops of a quartz grain from the beaches**  
 332 **east of the mouth of the Aby lagoon in Assinie (data [15])**  
 333

334 The traces in "V" in Figs. 12a and 12b are shapes which, in evolution, precede the figure in  
 335 "pyramid", witnesses of an important evolution. The constant orientation of these figures  
 336 makes it easy to distinguish them from similar-looking figures which, on reworked quartz, are  
 337 caused by shocks or pressures. Occasionally, triangular dissolution figures can be seen at the  
 338 bottom of some shock marks.  
 339

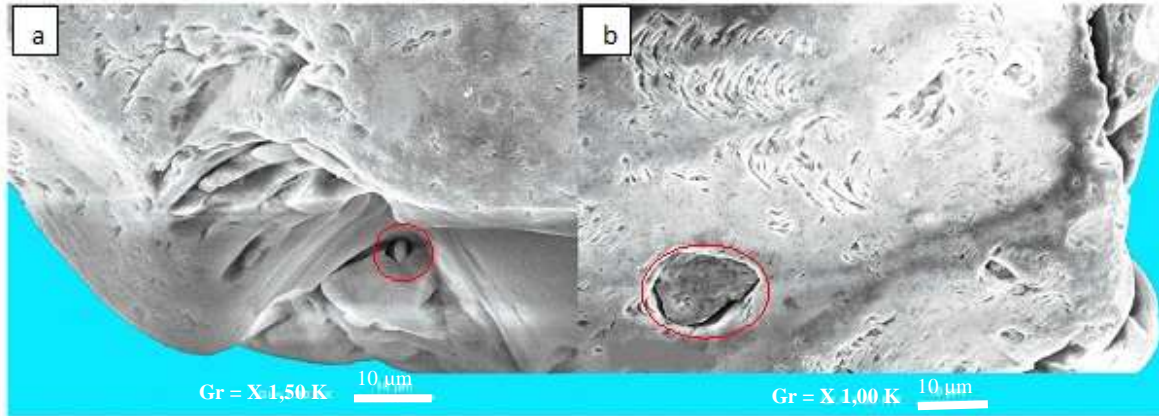


340  
 341

342 **Fig. 12. Appearance of the "V" figures and pressure traces (b) on quartz from the**  
 343 **beaches east of the mouth of the Aby lagoon in Assinie (data [15])**

344 Figs. 13a and 13b show shallow shock cups acquired during transport and that transport is  
 345 prior to the dissolution action. Solid inclusions are represented by minerals having crystallized  
 346 before or at the same time as quartz in granitic magma. The most frequent are tabular or  
 347 rounded crystals (Fig.13b). The quartz grains in this coastal zone come from various fluvial,  
 348 continental, marine and reworked environments.  
 349





**Fig. 13. Traces of shocks and inclusions on the surface of the quartz grains of the beaches east of the Aby lagoon in Assinie (data [15])**

#### **4.5.2 Transport environment, evolution and deposition of quartz beaches east of the mouth of the lagoon Aby (Assinie)**

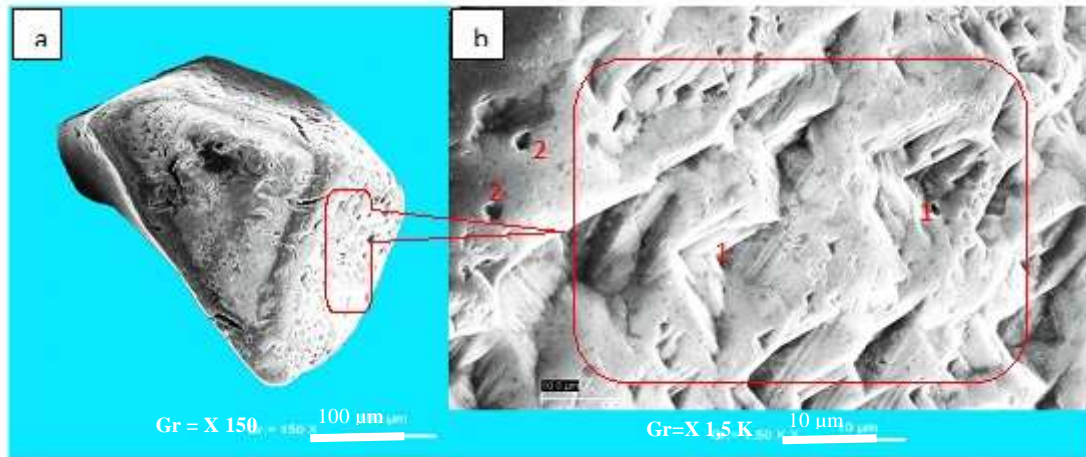
Fig. 13a and 13b show shallow shock cups acquired during transport. Solid inclusions are represented by minerals having crystallized before or at the same time as quartz in granitic magma. The most frequent are tabular or rounded crystals (Fig. 13b). The quartz grains in this coastal zone from various rivers, continental, marine and reworked environments.

However, the appearance of certain sharp edges (Fig.11a) and particular figures "V" and "pyramid" or "stairs" respectively translate turbulent media (Fig.13a), pressure environments (V) and several media different concentrations. These numerous imprints superimposed a few times on the quartz grains coming from the beaches of this zone of the Ivorian littoral of the Atlantic Ocean translate various continental, fluvial, marine and reworked environments.

#### **4.6 Quartz grain environment of Aforenou Beach (Ghana)**

The observations show a very blunt grain with a cavernous and very irregular microrelief. At low magnification (x150), the quartz grain shows a corroded, cavernous surface (Fig. 14a). A magnification of 1500 times the corroded surface (Fig.14b) has cavities of depth and irregular size with parallel lines arranged in stairways, making between them sharp angles intersecting at 60 °, 90 ° and 120 ° approximately (1). The primary points of corrosion on flat surfaces are hollow and circular, rainbow and triangular (2).

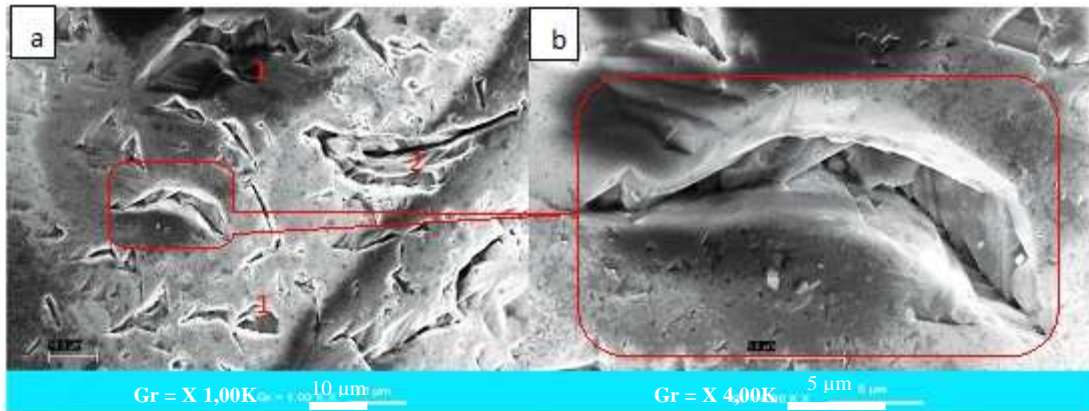




(1-Cavities in stairs whose vertices form angles; 2-Primary points of corrosion)

**Fig. 14. Important corrosion of the quartz grain in the form of cavities in which are visible parallel lines arranged in stairs at Aforenou (Ghana) (data [15])**

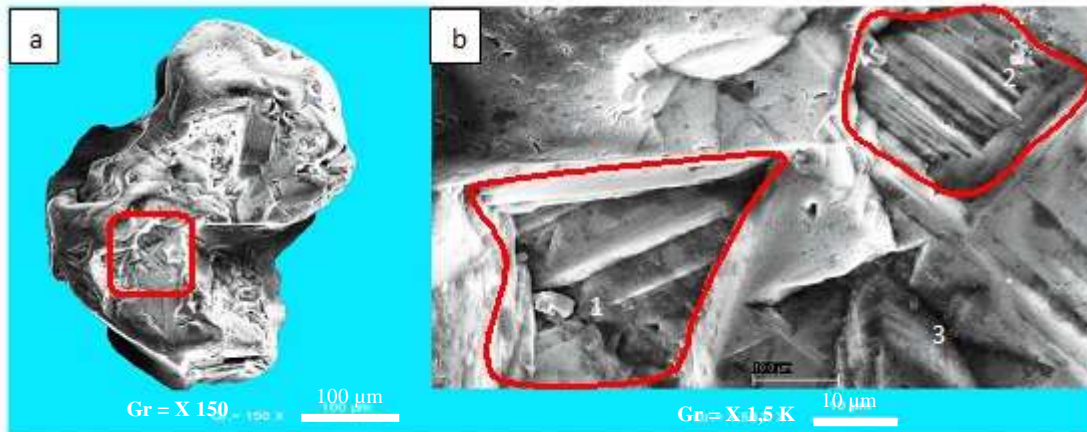
The surface of the quartz grain indicates marks of deep shocks in «nail stroke» of varied orientation (Fig.15a). These quartz grains could come from a wind transport. The edges of these marks are bright to angular visible at a magnification of 4000 times (Fig.15b). However, some of them are half-erased or have a blunt edge due to fluvial wear (Fig.15a (2)) during its course towards the oceanic domain. What [23] calls "polishing gradient". Some triangular markings, in "nail stroke" are still fresh and result from shocks.



(1 and 2 = traces of old shocks and blunted by fluvial wear,data [15])

**Fig. 15. Numerous nail strokes appear in Aforenou (Ghana)**

In Fig. 16a (x150), numerous traces of shock and chemical dissolution appear. At high magnification (x 1500, Fig.16b), there appears a series of parallel streaks. These figures are observable in the grooves along these streaks. Their width is of the order of 0,5 µm and more. Disaggregated particles in the process of desquamation still adhere to this surface (1), (2) and (3).



(1, 2 and 3 = Parallel lines containing disaggregated particles; data [15])

**Fig. 16. Sand grain with many traces of shocks with dissolution figures in Aforenou (Ghana)**

#### 4.6.1 Environment of transport, evolution and deposit of the quartz of the beaches of Aforenou

We observe the presence of very blunted quartz grains on these beaches with corroded surfaces for some and deep cavities (nail claws) some of which in stairs, and quartz sharp edges and angular for others with blunt edges. The appearance of quartz grains on these beaches at the border with Ghana, whose particle size is very fine (less than 200 microns) indicate that they were transported mainly by the wind. They were blunted by fluvial wear by transport and transit through the Aby lagoon (Fig.15a (2)) as it traveled to the oceanic area. This definitive deposition medium has had the effect on certain quartz chemical dissolution.

## 5. DISCUSSION

### 5.1 General situation

Examination of the surface of the quartz grains in the MEB revealed a wide variety of impressions that testify to several environments traversed by them. They are comparable to those described by [2, 3, 31] who grouped them into two large groups. They are in the form of relief or hollow and are the traces of the environments crossed during the displacement or deposit of sands. **They result from mechanical actions or characters of chemical origin.**

Relief figures less frequent than those in hollow on the surfaces of quartz result from the crystallization of silica in an undersaturated marine environment or chemical dissolution. Some that can be offset from each other are running stairs when the offset becomes appreciable. Such figures have been observed and described by many authors [9, 10, 19, 32, 33] who testify that these quartz grains have undergone wind or water transport.

We also note the scattered presence on some of these surfaces of small scales of desquamation, particles in veneer as well as figures in nail clipping and thin elongated cavities. The hollow figures may be generally attributed to the dissolution of the quartz; while the relief figures can be interpreted as figures of crystalline decay [9, 10].

## 440 **5.2 Quartz grain environment on the beaches between Abidjan and Grand Bassam**

441

442 The southeastern coast of the Gulf of Guinea between Abidjan and Aforenou in Ghana,  
443 consists mainly of loose sediments, mostly quartz grains. These sands of neogene age have  
444 various forms. The surface of these grains bears marks, traces or figures that trace the various  
445 environments or events during their transport. The work of [7, 8] reached similar conclusions  
446 by studying pebbles on the beaches of Sorso (Sardinia), Liguria and Corsica.

447 The very rounded quartz grains, some of whose dull edges are found on the coast of the Gulf  
448 of Guinea between Abidjan and Grand-Bassam, bear many crescent marks with triangular  
449 figures in steps or pyramids. These marks indicate a wind transport of the grains before being  
450 taken up in an aquatic environment (Fig.5a and 5b).

451 The SEM observation of 3m deep sands in the Cechi region (Agboville, Côte d'Ivoire) by [19]  
452 also showed frequent pyramid-shaped triangular-based figures indicating a reworked quartz.  
453 In addition, the majority of the quartz grains in these samples have an ochre yellow to reddish  
454 hue. This characteristic suggests a certain pedogenetic influence, a probable index of a  
455 paleosol, a soil that evolved during the Quaternary and Tertiary periods according to [19].

456 The presence of solid inclusions, milky (Fig.3a, 3b and 3c) or vitreous, almost completely  
457 crystallized or mineral enclaves housed at the bottom of the depression of the shock mark  
458 with siliceous deposits located on some faces (Fig. 7 and Fig. 9) confirms that the quartz grain  
459 was taken up in an intertidal medium or marine domain under saturated silica described by  
460 [4].

461 However, for [34] the siliceous deposits indicate a low energy continental or fluvial medium.  
462 This low energy environment is indicated by the presence on the surfaces of the quartz grains  
463 of numerous siliceous deposits, confirmed by phytoplankton that are the diatoms found in the  
464 old dissolution figures or shock figures. The presence of these unicellular algae testifies to a  
465 calm environment.

466 Indeed, the closure of the mouth of the Comoé River located in Grand-Bassam (about 30 km)  
467 allows sediments, mainly quartz grains, to transit through the Ebrié lagoon (low-energy calm  
468 environment), connected to the sea by the Vridi canal (access to the port of Abidjan) before  
469 being evacuated to the Atlantic Ocean.

470

## 471 **5.3 Quartz grain environment in the coastal zone Assinie-Aforenou (Ghana)**

472

473 In the near-coastal environment of Ghana (from Assinie to Aforenou), the quartz grains,  
474 which are mostly rounded, are the expression of the mechanical action of the waves (Fig.8a,  
475 9a and 11a) and (Fig.14a and 16a) or transport over a long distance. The surfaces of the grains  
476 are covered with cavities of variable depth and size, irregular with often parallel lines  
477 arranged in steps (Fig.8b, 8c, 14a, 14b, 16a and 16b). The arrangement of the ridges shows  
478 angles of 30 ° C, 60 ° C, 90 ° C and 120 ° C. These cavities, generally triangular, with  
479 pyramidal points have been reported by [19] on quartz of some ferralitic soils in the region of  
480 Cechi (Agboville, Ivory Coast).

481 Moreover, the works of [9, 35] have shown that they are characteristic of tropical zones.  
482 These samples have a continental origin and come from the catchment area upstream. [23]  
483 emphasizes that these figures constitute an evolution of the aeolian shock croissants exploited  
484 by the dissolution, on which appear the triangular figures in hollow characteristic of these  
485 (Fig.12b). The slightly curved deep striations on which is implanted a succession of "V" in  
486 characteristic form in fishbone (Fig.12b) are traces of strong pressures such as volcanic  
487 quartz.

488 Numerous marks of shocks and very tight nail (Fig.15a and 15b) without preferential  
489 orientation are observed on the surfaces of many quartz grains with sometimes edges blunted

490 by fluvial wear or marine; it is a "polishing gradient". Referring to the analytical work of [19,  
491 36], we attribute these shock marks to wear during a wind transport or a violent torrent [37,  
492 38].

493 The detailed analysis of the surface of certain quartz grains shows that it is rough and covered  
494 with ovoid particles. These are coatings in the form of oval beads on some very rounded and  
495 blunted grains that can be compared by morphological analogy to the amorphous silica film  
496 described by [2]. This form would indicate a fluvial or continental medium of low energy.  
497 Indeed, [34] report that the precipitation of siliceous globules over the entire surface of the  
498 grains, including the vertex of the ridges (Fig.9) indicate a final deposit in a low energy  
499 aquatic continental medium.

500 The steady increase in the percentage of blunted from Abidjan to Ghana (Aforenou) (50% to  
501 70%) shows that the sands undergo wear during their transport by coastal drift to the east. The  
502 swell reshapes the sediments because of its orbital movements that are transformed near the  
503 bottom into alternating currents that can reach a high speed and put in motion the materials  
504 that can be either reworked on site or driven by perpendicular or parallel currents. The  
505 increase in the percentage of blunting blunted grains indicates that there is equipment  
506 transport from Abidjan to Aforenou in Ghana in the direction of the littoral drift (west-east).

507

#### 508 **5.4 Origin of the soft sediments of Ivorian beaches**

509

510 The continental origin of some of the beach sediments is confirmed by the studies of [39].  
511 Indeed, if we refer to the summary work of [39], as part of the Sassandra-Cavally operation  
512 (SasCa) between 1962 and 1968 in the south-west of Côte d'Ivoire, these minerals of  
513 geological interest constitute the bulk of the sediments of the beaches studied. Thus, one more  
514 argument to confirm the continental origin of littoral sedimentary deposits, suggested in the  
515 detailed study of the quartz grain surface of the emerged beaches and the granulometry  
516 [14,15].

517 Studies by [16, 41] on the origin of sediments in the Fresco lagoon indicated mainly  
518 continental inputs (wind, runoff and Bolo and Niouniourou rivers) and oceanic. Some  
519 sediments come from areas close to beaches [14, 17, 18, 42]. This is the case in this study of  
520 the sediments of the beaches of Abidjan, close to the entrance channel in the port of Abidjan  
521 (Vridi Canal).

522

523

#### 524 **6. CONCLUSION**

525

526 The grains of sand have many traces of shock and dissolution figures on their surface. Marine  
527 sand combines bleached and matte round blunted quartz grains, evidence of the plurality of  
528 coastal wear agents. The slight increase in the percentage of blunted blunted grains from  
529 Abidjan to Aforenou in Ghana indicates the drift (West-East) in this part of the Gulf of  
530 Guinea.

531 The grains released during the rock weathering process were transported and finally  
532 sedimented in a low-energy aquatic continental medium (Ebrié lagoon and Aby lagoon with  
533 their respective estuaries) marked by the polishing of the crystalline points and the siliceous  
534 globules. sometimes scattering all the surface of the grains. The main attacker of quartz  
535 crystals is thus transport (air, fluvial or marine), during which they are rubbed against each  
536 other. The most exposed parts are the grain edges, which are affected by polishing gradient  
537 shock traces, recent for the angular and old for the blunt ones. These traces are even larger  
538 and more numerous than the transport energy is high.



539 The SEM examination indicates that some of the quartz have fluvial shaping marks, the  
540 unworn quartz carries abundant marks of chemical dissolution. The still varied, better sorted  
541 and largely dull-glistening grains of these samples reveal the impact of hydraulic transport on  
542 the original material over a long distance driven by the eastward drift of the original arena.  
543 During this study, it was impossible to find two identical grains of sand. They all carry a  
544 multitude of specific information concerning their origin (continental, fluvial), and each  
545 episode of their existence. The SEM examination of quartz mineral samples revealed that the  
546 sediment present on the beach of the study area are quartz that have generally undergone  
547 violent air transport, before fluvial transport and taken up in an intertidal environment and /or  
548 infratidal.  
549 They result from a mixture of sands brought by the continent (Comoé and Tanoé rivers) and  
550 the sea (reworking).

551

552

## 553 REFERENCE

554

555 1. Strawberries A. Distinction of marine and fluvial pebbles, Bull. Soc. Geol. France, 5th  
556 series, XV. 1947 375-404.

557

558 2. The Ribault L. Presence of an amorphous silica film on the surface of quartz crystals of  
559 sandy formations. Cah. ORSTOM, slv. Geol. IV. 1971; 53-65.

560

561 3. The Ribault L, Tastet JP. Contribution of quartz exoscopy to the determination of the origin  
562 of coastal quaternary deposits of Côte d'Ivoire. Proceedings of the 1978 International  
563 Symposium on Coastal Evolution in the Quaternary, I.G.C.P., Project 61, Sao Paulo (Brazil)  
564 1979; 573-587.

565

566 4. Legigan P, Turon JL, Weber O. Evolution of coastal deposits during a climatic cycle on the  
567 North Atlantic coast, Bull. Inst. Geol. Bassin d'Aquitaine, Bordeaux. 1986; (39): 135-147.

568

569 5. Barusseau JP, Kounkou GL. Origin of surface textures on quartz grains in sediments of the  
570 upper Cayar deep-sea fan (Senegal-Gambia abyssal plain) Oceanologica Acta, Marine  
571 Sedimentology Research Laboratory, University of Perpignan, Cedex, France. 1989; 12 (2):  
572 117-129.

573

574 6. Mitsova H, Overton M, Harmon RS. Geospatial analysis of a coastal sand dune field  
575 evolution: Jockey Ridge, North Carolina Elsevier, Geomorphology. 2005; 72: 204-221.

576

577 7. Ozer A. Study of variations of shape and size of pebbles on the beach of Sorso (Northern  
578 Sardinia); Bulletin of the Geographical Society of Liège. 1978; 14: 117-126.

579

580 8. Ozer A, Omhaire AL. Contribution of morphometry of pebbles to the knowledge of coastal  
581 transport. Example in Sardinia, Liguria and Corsica. Bulletin of the Royal Society of  
582 Sciences. 1988; 4-5: 429-440.

583

584 9. Leneuf N. Microscopic aspects of the quartz grain surface of the Continental Terminal of  
585 Côte d'Ivoire, Inst. of Earth Sciences, University of Dijon, Cah. ORSTOM, ser. Geol., IV.  
586 1972; (1): 53-65.

587

- 588 10. Leneuf N. Stereoscopic observations on the quartz corrosion patterns in certain superficial  
589 formations. Cah. ORSTOM ser. Pedol., XI. 1973; (1): 43-51.  
590  
591
- 592 11. Martin L. Morphology, sedimentology, Quaternary paleogeography of the Ivorian  
593 continental shelf. Works and document, 61, ORSTOM, Paris. 1977; 265.  
594
- 595 12. The Ribault L. A method for determining the history of sand grains: exoscopy, Bulletin of  
596 the South West Anthropology Society, Volume IX, No. 2. 1984; 123-137.  
597
- 598 13. Tastet JP. The Ivorian coastline: geology, morphology and dynamics. Ann. Univ. Abidjan,  
599 Ser. C, 21 B. 1985; 189-218.  
600  
601
- 602 14. Konan KE. Contribution to the study of morphological and sedimentological evolution of  
603 the littoral between Grand Bassam and Assouindé, memory DEA, University Cocody. 2004 ;  
604 86.  
605
- 606 15. Konan KE. Morphodynamic study and sensitivity to the exceptional events of the Ivorian  
607 sandy coastline east of Abidjan (Abidjan-Aforenou). PhD Thesis, Félix Houphouët Boigny  
608 University of Abidjan (Ivory Coast). 2012; 224.  
609  
610
- 611 16. Yao NA, Adopo KL, Amani EM, Konan KB, Toure M, World S, Aka K. Bathymetric,  
612 sedimentological and environmental study of superficial sand deposits of the Fresco lagoon  
613 (western zone of the Ivorian coast), Journal of Asian Scientific Research. 2013; 3 (3): 308-  
614 320.  
615
- 616 17. N'doufou GHC, Abe J, S Bamba, Hauhouot C, Aka K. Effects of the opening of the Vridi  
617 Canal on littoral sedimentary stocks between Abidjan and Jacquville (Ivory Coast), Paralia  
618 Review. 2015 8. 1.16.  
619
- 620 18. N'doufou GHC, Konan KE, AE Boga, World S, Abe J. Origin and evolution of littoral  
621 quartz between Fresco and Jacquville, Ivory Coast, Africa Science. 2018; (1): 171-180  
622
- 623 19. Flagellot JC. Morphoscopic and exoscopic aspects of quartz in some ferralitic soils of the  
624 Cechi region (Ivory Coast), Cah. ORSTOM., Ser. Pédol.1981; 18: 111-121.  
625
- 626 20. Mhammdi N, Ahab M, Hamoumi N, Azza A. The titaniferous sands of the Azemmour  
627 coast and the Oum Er-Rbia estuary (Moroccan Atlantic coast): sedimentology and  
628 exploitation potential, Bulletin de l' Scientific Institute, Earth Sciences Section. 2005; 27: 83-  
629 91.  
630
- 631 21. Mahaney CW, Dohmb JM, Costa P, Krinsley DH. Tsunamis on Mars: Martian sediment,  
632 Planetary and Space Science. 2010; 58: 1823-1831.  
633
- 634 22. Bellahbib N, Rezqi H., Oujidi M, Bengamra S. Granulometric and mineralogical study of  
635 the superficial sediments of the littoral of Saïdia and the estuary of Moulouya (north-east of  
636 Morocco), Larhyss Journal. 2015 ; 24: 19-40.  
637

- 638 23. Ribault L. Exoscopy of quartz, Masson et Cie editors, Paris. 1977; 150.  
639
- 640 24. Arnaud JC. The relief of Ivory Coast, Atlas of Ivory Coast, Ed. JA. 1978; 6-7  
641
- 642 25. Tastet JP, Caillon L, Simon B. Coastal sedimentary dynamics in Abidjan: impact of  
643 developments. Contribution to the understanding of the phenomena of erosion and  
644 sedimentation. Rapp. Min., Marine, Abidjan. 1985; 39.  
645
- 646 26. Parpenoff A, Pomerol C, Tourenq J. Minerals in grains. Methods of study and  
647 determination. Masson Ed. (Paris). 1970; 578.  
648
- 649 27. Pinot JP. Common sedimentological manipulations. 1994; 118.  
650
- 651 28. Cailleux A, Tricart J. Initiation to the study of sands and pebbles. Documentation Center  
652 Univ. edict. Paris, 1, ed., 5, Place de la Sorbonne. 1959; 379.  
653
- 654 29. Saaidi E. Sedimentology Treaty. Ellipses Edition. 1991; 393.  
655
- 656 30. Folk RL, Ward WC. Brazos River bar: a study in the significance of grain size parameters.  
657 J. Sedim. Petrol. 1957; 27 (1): 3-26.  
658
- 659 31. Deicha G, Sella C. Investigation of inter-granular cavities by scanning electron  
660 fractography. CR som. Soc. Geol. En.3. 1971; 179-181.  
661
- 662 32. Muller D. Contribution to the study of the differentiation of nodular horizons of  
663 Congolese ferralitic soils on granito-gneiss. 3rd cycle thesis, Univ. Paris VII. 1979; 118.  
664
- 665 33. Muller D, Bocquiere GN, Ahon D, Package H. Analysis of the mineralogical and  
666 structural differentiations of a ferralitic soil with nodular horizons of the Congo. Cah.  
667 ORSTOM, ser. Pedol., XVIII. nineteen eighty one ; 2: 87-109.  
668
- 669 34. Glocchiatti R, The Ribault L, Rodrigo LA. Endoscopy and exoscopy of quartz grains of  
670 the Pliocene and Quaternary formations of La Paz (Bolivia), Cah. ORSTOM, ser. Geol., X.  
671 1978; (1): 127-143.  
672
- 673 35. Leprun JC. The ferruginous cuirasses of the crystalline countries of dry West Africa.  
674 Genesis, transformation, degradation. Thesis Sci. Geol. Same. Strasbourg. 1979; (58): 224.  
675
- 676 36. The Ribault L. Exoscopy, method and application, notes and memoirs, Comp. Franc. Petr.  
677 1975; (12): 232.  
678
- 679 37. Costa PJM, Andrade C, Dawson AG, Mahaney WC, Freitas MC, Paris R, Taborda R.  
680 Microtextural characteristics of quartz grains transported and deposited by tsunamis and  
681 storms, Elsevier, Sedimentary Geology. 2012 ; 275 :55-69.  
682
- 683 38. Costa PJM, Andrade C, Mahaney WC, Marques DSF, Freire P, Freitas MC, Janardo C,  
684 Oliveira MA, Silva T, Lopes V. Aeolian microtextures in silica spheres induced in a wind  
685 tunnel experiment: Comparison with aeolian quartz, Elsevier, Geomorphology. 2013 ;120-  
686 129.  
687

688 39. Papon A. Geology and mineralization of southwestern Côte d'Ivoire. Summary of the  
689 work of SASCA 1962-1968. Bull. Direction of Mines and Geology of Ivory Coast. # 6. 1973;  
690 285.  
691  
692 40. Konan KE, Abe J, Akka K, Neumeier U, Nyssen J, Ozer A. Impacts of exceptional swells  
693 on the Ivorian coast of the Gulf of Guinea; rev. Geomorphology: Relief, Process,  
694 Environment. 2016; 22 (1): 105-120.  
695  
696 41. Issola Y, Kouassi AM, Dongui BK, Adingra AA, Biemi J. Heavy metal concentration of  
697 sediments in a tropical coastal lagoon: Fresco lagoon (Ivory Coast), Journal of Applied  
698 Biosciences. 2009; 18: 1009-1018.  
699  
700 42. Degbe CGE. Geomorphology and coastal erosion in the Gulf of Guinea. International  
701 Chair in Mathematical Physics and Applications (CIPMA - UNESCO Chair) Master of  
702 Science in Physical Oceanography, Univ Abom. Cala. (Benign). 2009; 100.  
703  
704  
705  
706  
707  
708  
709

UNDER PEER REVIEW